Research Article

Trends in Extreme Precipitation Indices in Iran: 1951–2007

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We investigate trends in extreme precipitation in Iran for 1951–2007 using the recently released APHRODITE daily rainfall time series. We find that seven different indices of extreme precipitation all show an upward trend through the study period. The seven different precipitation indices include annual precipitation total, number of days above a certain threshold, maximum precipitation received over a certain period of time, maximum one-day precipitation, and number of days with precipitation above the 90th percentile. A principal components analysis reveals one eigenvector explaining much of the variance in the seven indices and reveals that this component exhibits a strong upward trend for the whole of Iran. On a regional level, we find that the upward trend in extreme precipitation has a strong southwest-to-northeast gradient across the country for all the indices. We repeated all the analyses for 42 stations across the country to compare with the results from the gridded data; trends in extreme rainfall generated from the station data compare favorably with the results from the APHRODITE daily rainfall time series thereby reinforcing the robustness of our conclusions.

1. Introduction

An extreme event is generally defined as the occurrence of a weather or climate event above or below a threshold value near the upper or lower ends of the range of observed values for a specific variable. Some of the main findings from the latest report on climate change and a special report on extreme events from the Intergovernmental Panel on Climate Change indicate a greater consensus within scientific literature about a likely increase in the frequency and intensity of heavy precipitation events over land areas since 1950, with a likely increase in the frequency of heavy precipitation or proportion of heavy precipitation in the 21st century. Additionally, there is medium confidence about the anthropogenic influence on the intensification of extreme precipitation on the global scale [1, 2]. In a comprehensive study examining the spatial patterns of precipitation extremes, Alexander et al. [3] reported a significant increase in precipitation extremes, with less spatially coherent patterns compared with trends in extreme temperatures. In view of the widespread impacts of the extreme events on human, ecological, and/or physical systems, there is increased focus on the long-term trends of such events across the globe. The majority of studies based on observational data indicate a general increase in extreme heavy precipitation events, which is attributed to anthropogenic forcing caused by increased levels of moisture in the atmosphere and warmer temperatures overall [4].

Indeed, enhanced levels of water vapor due to warmer oceans in the lower latitudes were found by Trenberth et al. [5]. This has been further confirmed by increasing trends in specific humidity at the global level since 1970 [6, 7]. Additionally, results from CMIP3 and CMIP5 simulations also show an increase in the globally averaged 20-year return values of annual maximum 24-hour precipitation amounts at approximately 6 to 7% for each °C of global average warming [8, 9]. Several studies have highlighted the significant role of large scale circulation patterns on the positive trends in extreme precipitation events also [10, 11].
There are significant variations in long-term trends in precipitation regionally, with most of the continents, except North America and Europe, showing medium confidence in increasing trends in the frequency and intensity of observed heavy precipitation events [12]. Additionally, the majority of the land areas show an increase in extreme events during the summer season, except Europe which experienced increasing trends during the winter season [12]. Some of the specific regional level studies for extreme precipitation events show likely increases in North America [13–15] and South America [14, 16] for the entire 20th century. In the case of Europe and the Mediterranean region, the trends were mixed with most of the increase observed during the winter season and decreasing trends in the summer season precipitation [17–19]. Most of Asia and Oceania experienced mixed trends, with a greater proportion of the region showing a positive trend in extreme precipitation events [20–24]. Additionally, in the case of Africa, there were no clear significant trends [12, 25]. However, on the other hand there are a limited number of studies examining the trends in extreme precipitation patterns in the Middle East, likely due to the lack of long-term data, where Iran, the main focus of this study, is located. One of the few studies includes an analysis of station level precipitation across Iran in which Alijani et al. [26] found more than 20% of the land area exposed to the risk of extreme rainfall. Specifically the hot, dry southern coast and the western slopes of the main north–south range, the Zagros Mountains, experienced the most frequent heavy rainfall events. Furthermore, Rahimzadeh et al. [27] examined extreme temperatures and precipitation across 27 synoptic stations across Iran. The results of their analysis indicated marked negative trends in cool days and cool nights, diurnal temperature range (DTR), and positive trends in warm days and tropical nights. Recently, with the advent of gridded data developed from station level measurements at high spatial resolutions, it has become possible to conduct detailed analysis of extreme precipitation events over these regions. Accordingly, in the present study we have analyzed trends in extreme precipitation for Iran which represents a location with precipitation coming largely from cool-season cyclonic events and limited warm-season convective precipitation.

2. Study Area: Iran

Iran is located between 25° and 40°N and 45° and 60°E and is a mountainous country bordering the Gulf of Oman, the Persian Gulf, and the Caspian Sea (Figure 1). The total area of Iran is 1.648 × 10^6 km^2 which represents 0.32 percent of the Earth’s surface. Overall, sixty percent of Iran is covered
by mountains, with the central part of the country consisting of two dry deserts: the Dasht-e-Kavir and the Dasht-e-Lut. The Alborz range in the north, close to the Caspian Sea, extends in an east–west direction with a maximum elevation of approximately 5000 m. The Zagros Mountains are aligned in a northwest-to-southeast direction and reach a maximum elevation of approximately 3500 m. These two ranges play a significant role in determining the nonuniform spatial and temporal distribution of precipitation across the entire country [28]. For instance, the high ranges of the Alborz Mountains in the north and Zagros Mountains in the west inhibit much of the moisture available from adjacent water bodies from reaching the interior of the country. Thus, the interior parts of the country receive much less precipitation. Most of the interior slopes of the Zagros Mountains experience a rain shadow effect with annual rainfall much less than their western counterparts. More than half of the country receives less than 200 mm of precipitation, with some regions that get less than 50 mm annually [26].

Over the past decade, important papers have been published focusing on trends in precipitation across Iran (e.g., [29, 30]). The analysis of station level precipitation data revealed a decreasing trend in annual rainfall at 67% of the stations, while an increasing trend was observed in the 24 hr maximum rainfall at 50% of the stations by Modarres and Sarhadi [29]. Recently, Tabari and Talaei [31] studied temporal trends in the annual rainfall time series in the west, south, and southwest of Iran during 1966–2005. The results revealed no visible rainfall trends in the region for their study period. Soltani et al. [32] investigated annual and monthly trends in rainfall amount, number of rainy days, and maximum rainfall in 24 h based on the data collected at 33 synoptic stations in Iran. The results indicated that there are no significant linear trends in monthly rainfall at most of the synoptic stations. Someè et al. [33] investigated the spatiotemporal trends and variability of precipitation data from 28 synoptic stations in Iran on the annual and seasonal timescales for the period 1967–2006. Their results revealed negative trends in annual precipitation at 22 sites (79%), but only three sites had a statistically significant negative trend in precipitation. However, a subsequent study by Raziei et al. [34] using the gridded APHRODITE dataset found an upward trend in daily annual precipitation in most of Iran. Talaei [35] analyzed the annual, seasonal, and monthly rainfall time series at seven rain gauge stations in Hamadan Province located in the west of Iran for a 40-year period (from October 1969 to September 2009). Their results showed no clear rainfall trends for this Province of Iran. Alijani et al. [26] analyzed daily rainfall concentration and intensity over Iran using 90 stations over varying time periods. The results showed that daily precipitation tends to be irregular and intense across much of Iran and that a disproportionately large share of the annual rainfall comes from a small number of high-intensity-to-extreme rainfall events. Using meteorological data from the Urmia synoptic station, Delju et al. [36] analyzed climate variability and change in the Urmia Lake Basin in the northwest of Iran. They found that mean precipitation has decreased by 9.2% during 1964–2005.

3. “APHRODITE” Dataset

The Asian Precipitation–Highly Resolved Observational Data Integration towards the Evaluation of Water Resources (APHRODITE) dataset is a long-term daily gridded precipitation dataset for Asia, which is based on a dense network of rain gauges. The spatial resolution of this dataset is 0.5° latitude by 0.5° longitude. The data used in the construction of this gridded dataset are from three sources which include GTS-based data (the global summary of the day), data precompiled by other projects or organizations, and APHRODITE’s own collection [37]. In this paper the version APHRO_V1101 of APHRODITE was used. In case of Iran station level rain gauge data were provided by the Iran Meteorological Organization. Several quality control measures were conducted on the raw dataset, such as checking for conversion between units of millimeters and inches, discrepancies between two or more databases containing the same measurements, and comparing locations with national boundaries and elevation. Next, the station level data were interpolated using a modified version of the Shepard [38] algorithm, which takes into consideration local elevation differences and horizontal distances. The interpolation technique used in APHRODITE products takes into consideration the ratio of daily precipitation to daily climatology. Further detailed information about the interpolation techniques and additional detailed information about the creation of the APHRODITE dataset are available from Yatagai et al. [37, 39, 40]. This dataset has been successfully used to study different aspects of precipitation patterns in Iran by [30, 41]. It is important to note that the APHRODITE dataset is limited in its estimation of precipitation for high altitude regions in Asia [42–44]. However, the assessment of this dataset for estimation of precipitation patterns across Iran indicates higher levels of accuracy and efficiency [45]. More specifically, a recent study by Ghajarnia et al. [46], consisting of a comparative evaluation of various gridded datasets over Urmia Basin, Iran, revealed that the APHRODITE dataset was able to detect 75% of the rainfall events, and 39% of its rainfall estimations were no rain observations.

4. Station Precipitation Data

Daily rainfall data from 42 stations operated by the Meteorological Organization of Iran were used to evaluate extreme rainfall trend throughout the country (Figure 2). These stations were selected because they have the longest daily rainfall records and are fairly evenly spread throughout the country. Quality controls are applied to the data by the meteorological organization before releasing them to the user’s community. There are many weather stations in Iran, but we limited our study to only 42 stations that had data for the period from 1970 to 2009. 18 stations were used in this study with no missing data whatsoever. Another 11 stations had less than 1% of the daily data missing; however, 5 stations had less than 10% of the daily data missing. Eight stations had more than 10% of the data missing. We used several methods to address the missing data issue including substituting the missing value with the climatological average for that day and/or simply...
eliminating stations with more than 10% missing values. Our decisions did not appear to influence the final results in any meaningful way.

5. Analyses and Results

As seen in Figure 3, 618 of the $0.5 \times 0.5$ grid points in the APHRODITE dataset fall within the political boundary of Iran. Therefore, our initial matrix of daily precipitation consisted of 20818 rows, one for each day from January 1, 1951, to December 31, 2007, and 618 columns, one for each grid point in Iran. The average monthly precipitation for the entire area is shown in Figure 4. We analyzed seven different popular indices (for a similar analysis, see [22]) of extreme precipitation for each grid point and year including the following:

1. Annual precipitation total (AnnP).
2. Number of days with precipitation $\geq$ 10 mm (ND10mm).
3. Percent of annual precipitation from daily events $\geq$ 10 mm (%Ann10mm).
4. Number of days with precipitation $\geq$ 20 mm (ND20mm).
5. Number of days with precipitation $\geq$ the 90th percentile of the distribution (ND90%).
6. Maximum precipitation received over five consecutive days (Max5day).
7. Maximum one-day precipitation (Max1day).

The calculation of these indices by year and grid point resulted in a new matrix for each index of 57 rows, one for each year from 1951 to 2007, and 618 columns, one for each gridpoint. In order to compute a countrywide time series for each of the seven indices, the time series at each grid cell was converted to $z$-scores (mean of 0, standard deviation of 1), and the $z$-scores were then averaged across the 618 points. This resulted in a final matrix of 57 rows, one for each year, and seven columns, one for each extreme precipitation index. We conducted all analyses with and without an adjustment for leap years and found no meaningful changes in our results.

Recognizing that various statistical techniques used in this study assume normality (a Gaussian distribution) in these time series, we calculated the standardized coefficients of skewness, $z_1$, and kurtosis, $z_2$, calculated as

$$z_1 = \frac{\sum_{i=1}^{N} (x_i - \bar{X})^3 / N}{\left(\sum_{i=1}^{N} (x_i - \bar{X})^2 / N\right)^{3/2}}$$

$$z_2 = \frac{\left[\sum_{i=1}^{N} (x_i - \bar{X})^4 / N\right] \left[\sum_{i=1}^{N} (x_i - \bar{X})^2 / N\right]^{-2}}{24/N^{1/2}} - 3$$

where the resulting $z$ values are compared against a $t$-value deemed appropriate for a selected level of confidence (e.g., for $N = 57$, $t = 2.66$ for the 0.99 level of confidence). If the absolute value of $z_1$ or $z_2$ exceeds the selected value...
Table 1: Normality test results.

<table>
<thead>
<tr>
<th>Index</th>
<th>$z_1$</th>
<th>$z_2$</th>
<th>K-S</th>
</tr>
</thead>
<tbody>
<tr>
<td>AnnP</td>
<td>0.707</td>
<td>-1.196</td>
<td>0.082</td>
</tr>
<tr>
<td>ND10mm</td>
<td>0.865</td>
<td>-1.439</td>
<td>0.096</td>
</tr>
<tr>
<td>%Ann10mm</td>
<td>2.510</td>
<td>0.338</td>
<td>0.132</td>
</tr>
<tr>
<td>ND20mm</td>
<td>2.827</td>
<td>0.118</td>
<td>0.180</td>
</tr>
<tr>
<td>ND90%</td>
<td>0.583</td>
<td>-1.191</td>
<td>0.087</td>
</tr>
<tr>
<td>Max5day</td>
<td>1.234</td>
<td>-0.750</td>
<td>0.086</td>
</tr>
<tr>
<td>High1day</td>
<td>2.015</td>
<td>-0.186</td>
<td>0.115</td>
</tr>
<tr>
<td>Component 1</td>
<td>0.081</td>
<td>-1.126</td>
<td>0.075</td>
</tr>
</tbody>
</table>

$z_1$ and $z_2$ are standardized indices of skewness and kurtosis; K-S is the Kolmogorov-Smirnov test statistic.

of $t$, a significant deviation from the normal curve is confirmed. Otherwise, no statistically significant deviation from a normal distribution is determined (the null hypothesis that the samples came from a normal distribution cannot be rejected). We also used the Kolmogorov-Smirnov one-sample test in which the variable is tested against another variable defined as having a normal distribution. It is similar to a $t$-test determining whether two variables were drawn from different populations. If the Kolmogorov-Smirnov test is statistically significant, we rejected the hypothesis that the observed data follow the normal distribution.

As seen in Table 1, the only time series with a significant deviation from the normal distribution was the ND20mm index as judged by the Kolmogorov-Smirnov and skewness tests. The deviation was not severe and it could be corrected using a square root transformation in which the sign is maintained and the square root is taken of the absolute value of the $z$-score. All analyses were conducted with and without this transformation and no meaningful differences were observed.

Given the general lack of deviations from normality, we calculated the Pearson product-moment correlation coefficients among the seven time series (Table 2). All intercorrelation coefficients were significant at the $<0.01$ level of confidence and many of the correlation coefficients were above +0.90. Given the high correlation among the different indices, a principal components analysis (PCA) was conducted to explain the predominant trends in all seven indices analyzed in the present study. The results of PCA to the matrix of seven indices revealed one component explaining 80.9% of the variance in the matrix. The loadings (Table 3) were all positive and ranged from +0.77 for AnnP to +0.99 for ND10mm. The strength of this one component suggests that it does adequately capture a robust dimension in the precipitation data related to all of the measures of extreme precipitation.

There was no significant deviation from normality in the component scores (Table 1) and as seen in Figure 5, the scores show a distinctive upward trend. Using the year of record as the predictor variable, a simple regression shows that the trend is upward and highly statistically significant ($p < 0.01$) for component 1 and for five of the seven indices; AnnP and ND90% had upward trends that were not significant at the 0.01 level of confidence.

As a final procedure with the APHRODITE dataset, we calculated the linear regression between ND10mm and year of record at each of the 618 grid points and plotted the $r$ values (Figure 6). A strong northeast-to-southwest gradient appeared; a simple first-order polynomial interpolation (basically a plane) explained 83% of the spatial variance in the data. The lower values were generally concentrated in the northern border of the study area. The more sophisticated universal kriging method, which is useful for estimating local trend, was used to produce the pattern shown in Figure 6 [47]. It explained 95% of the spatial variance in the correlation coefficient values. Spatial variance explained by an interpolated surface is determined as $1 - (\text{RMSE}^2/\text{SD}^2)$, where RMSE is the root mean square error of the interpolation and SD is the standard deviation of the variable being mapped. We repeated the trend analysis using the Mann-Kendall Rank Statistic which also determines the strength, sign, and significance of
Table 2: Pearson product-moment correlation coefficients amount the indices of extreme precipitation.

<table>
<thead>
<tr>
<th>Index</th>
<th>AnnP</th>
<th>ND10mm</th>
<th>%Ann10mm</th>
<th>ND20mm</th>
<th>ND90%</th>
<th>Max5day</th>
<th>High1day</th>
</tr>
</thead>
<tbody>
<tr>
<td>AnnP</td>
<td>1.00</td>
<td>0.78</td>
<td>0.39</td>
<td>0.53</td>
<td>0.99</td>
<td>0.66</td>
<td>0.57</td>
</tr>
<tr>
<td>ND10mm</td>
<td>1.00</td>
<td>1.00</td>
<td>0.87</td>
<td>0.92</td>
<td>0.81</td>
<td>0.92</td>
<td>0.93</td>
</tr>
<tr>
<td>%Ann10mm</td>
<td>1.00</td>
<td>1.00</td>
<td>1.00</td>
<td>0.95</td>
<td>0.45</td>
<td>0.85</td>
<td>0.96</td>
</tr>
<tr>
<td>ND20mm</td>
<td>1.00</td>
<td>0.56</td>
<td>0.89</td>
<td>1.00</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>ND90%</td>
<td></td>
<td>1.00</td>
<td>0.69</td>
<td>0.60</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Max5day</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>1.00</td>
<td>0.90</td>
<td></td>
</tr>
<tr>
<td>High1day</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>1.00</td>
<td></td>
</tr>
</tbody>
</table>

Table 3: Principal component loadings and trend results for extreme precipitation indices.

<table>
<thead>
<tr>
<th>Index</th>
<th>Loading</th>
<th>B</th>
<th>r</th>
</tr>
</thead>
<tbody>
<tr>
<td>AnnP</td>
<td>0.765</td>
<td>1.41</td>
<td>0.19</td>
</tr>
<tr>
<td>ND10mm</td>
<td>0.991</td>
<td>4.76</td>
<td>0.54</td>
</tr>
<tr>
<td>%Ann10mm</td>
<td>0.885</td>
<td>6.06</td>
<td>0.63</td>
</tr>
<tr>
<td>ND20mm</td>
<td>0.936</td>
<td>5.81</td>
<td>0.62</td>
</tr>
<tr>
<td>ND90%</td>
<td>0.796</td>
<td>1.87</td>
<td>0.24</td>
</tr>
<tr>
<td>Max5day</td>
<td>0.944</td>
<td>4.20</td>
<td>0.49</td>
</tr>
<tr>
<td>High1day</td>
<td>0.952</td>
<td>6.16</td>
<td>0.64</td>
</tr>
<tr>
<td>Component 1</td>
<td>4.82</td>
<td>0.54</td>
<td></td>
</tr>
</tbody>
</table>

B values are the standardized regression coefficients (beta weights) and r values are the Pearson product-moment coefficients between the variable and year of record.

Conclusions

We analyzed gridded daily precipitation data from APHRODITE across Iran over the period 1951–2007 and found a strong positive trend in extreme precipitation events. The upward trend appeared in seven different indices and was especially strong in a composite variable of extreme precipitation developed through a principal components analysis. On a regional basis, the trend upward in extreme precipitation was highest in the southwest and least in the northern portions of the country. We found very similar results when we conducted the same analyses based on daily precipitation data from 42 stations across Iran. Our results suggest that the APHRODITE gridded daily precipitation data can appear to be relevant for applications such as trend analysis of extreme events.

Raziei et al. [34] found that precipitation tends to decrease during the warm seasons (spring and summer) and increase of 0.97 with the time series generated based on all stations. The variations in the trends in the spatial patterns of extreme weather events are mainly driven by orography in the form of the mountains in the northern and western interiors [34].
during cold seasons (autumn and winter) in most of Iran. This would imply less precipitation occurrences during the warm season and an intensification of the seasonality and dryness over the country. Iran usually receives the largest proportion of its rainfall during cold seasons. The findings of the present study indicate a significant upward trend has occurred in extreme indices in southwest of Iran (meaning that the portion of extreme occurrences that contributed to the annual rainfall has increased significantly) along the Zagros Mountains where the most important and the biggest rivers of this region originate (e.g., Karoon River, Dez River). Other parts of the country including the southeast and northwest regions showed no trend over our study period. Increases in both intensification of the seasonality and dryness over Iran, along with the increase in extreme indices of rainfall, would be very hazardous for water availability in an arid region like Iran, which receives only 250 mm annually. Consistent with our findings, Alijani et al. [26] indicated that days with rain totals above the 90th percentile account for a disproportionate percentage of the nation’s total rainfall, and even very rainy areas are at risk of extreme rainfall and associated hazards. Our results, along with the findings of many other studies, suggest that water access will continue to be a challenge in Iran moving forward.

**Conflict of Interests**

The authors declare that there is no conflict of interests regarding the publication of this paper.

**References**


