The Influence of Injected Fluids on Microscopic Pore Structures in the Intersalt Dolomitic Shale Oil Reservoirs

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Given their low porosity and permeability, intersalt dolomitic shale oil reservoirs need to be developed via large-scale hydraulic fracturing to achieve economic effects. However, various lithologies and salt materials make these reservoirs vulnerable to salt mineral dissolution and recrystallization and salt plugging during development. This study investigated the variation rules of pore structures, porosity, and permeability of intersalt dolomitic shale oil reservoirs under the influence of hydraulic fracturing fluid. Correspondingly, a series of experiments (i.e., high-temperature and high-pressure soaking experiments, focused ion beam scanning helium ion microscope analyses, and porosity and pulse permeability tests) is performed on the Qian 3°10 rhythmic intersalt dolomitic shales. Results show that distilled water dissolves the salt crystals inside the matrix pores to improve the reservoir permeability. However, the distilled water-rock interaction will cause the massive migration of salt minerals. By contrast, the supercritical CO₂ can disperse salt particles, dredge the channels, and enlarge pores by expansion, but it has an overall weak capability of changing the pore structure and matrix permeability. To simulate the supercritical CO₂ composite fracturing, the mixed solution of supercritical CO₂ and distilled water favors the salt dissolution effect in the water-based fracturing fluid and recovery enhancement by CO₂. This solution can remarkably improve the reservoir porosity and permeability and avoid massive salt mineral migration and salt crystallization damage. This study is theoretically and practically important to the effective and enhanced development of intersalt dolomitic shale oil reservoirs.

1. Introduction

In recent years, shale oil has become a key area for unconventional oil and gas exploration and development after shale gas. China’s continental basins have developed a number of lacustrine mud shale strata with wide distribution, high abundance of organic matter, large thickness, and huge potential for shale oil resource, and they are considered a strategic option for the sustainable development of old oilfields in the east [1]. Among them, intersalt dolomitic shale oil reservoirs in the Qianjiang Depression of the Jianghan Basin have become an important target for China’s shale oil technology breakthrough due to its numerous oil-bearing strata, superior hydrocarbon source conditions, and good oil-bearing properties.

Intersalt dolomitic shale oil reservoirs in the Qianjiang Depression have become so effective that they require volumetric reconstruction through hydraulic fracturing. The production data after hydraulic fracturing show that the initial increase of production is good and that the daily production of a single well can reach 69.2 tons. However, after fracturing, the very high salinity of the fracturing fluid (up to 26 × 10⁴ mg L⁻¹) is ubiquitous, and salt particle crystals can be seen in the flowback fluid. Moreover, the stable production time is short, usually one to four months. After this period, the production drops rapidly. Given the particularity of the intersalt dolomitic shale oil reservoir, the reservoir contains soluble salts, and a thick salt layer develops at the top of the reservoir. After the water-based fracturing fluid enters the reservoir, it dissolves the soluble salt minerals inside the rock,
causing the migration and recrystallization of the salt minerals, which forms the "salt blockage" in the formation and wellbore. The seepage capacity near the wellbore gradually decreases, and the production decreases rapidly several months after the fracturing.

CO₂ fracturing is one of the waterless fracturing techniques that has unique production-increasing effects, such as increasing fracture complexity, reducing damage, improving crude oil flow, and increasing formation energy [2–6]. In the early 1980s, CO₂ fracturing has been used in oil/gas field practices and has achieved good production-increasing effects in many wells [2, 7, 8]. For intersalt dolomitic shale oil reservoirs, using CO₂ fracturing can avoid the dissolution and recrystallization of salt minerals caused by water-based fracturing fluids.

In terms of physical properties, CO₂ is easy to transform and can exist in gaseous, liquid, or supercritical fluid forms, which are primarily affected by pressure and temperature. When the temperature and pressure exceed 31.1°C and 7.38 MPa, respectively, CO₂ exists in a supercritical state. In this state, the intermolecular force of CO₂ is small, the surface tension is zero, and the fluidity is so strong that molecules can enter any space larger than the supercritical CO₂ molecule [9]. As CO₂ is easily soluble in formation water and forms carbonic acid, it can dissolve minerals such as feldspar and carbonate in rock under reservoir pressure and temperature. Therefore, the physical properties of the reservoir and the mechanical properties of the rock are changed [8, 10–14]. However, the change processes of microscopic pore structures and the porosity/permeability of shale reservoirs during supercritical CO₂ fracturing remain unknown. In addition, due to the differences in physical properties of shale reservoirs, the influence of supercritical CO₂ on the physical properties of the reservoir is different.

To investigate the influence of water-based fracturing fluid and supercritical CO₂ on the physical properties of an intersalt shale reservoir, a series of high-temperature and high-pressure shale immersion experiments was performed, followed by microscopic pore structure observation and porosity and pulse permeability tests. The pore structure, porosity, and permeability change of intersalt shale under different fluids and the damage law of salt crystallization on physical properties of shale are studied. The results are expected to provide theoretical support for the correct formulation of measures for stable production and increased production in intersalt shale oil reservoirs.

2. Overview of the Region

The Jianghan Basin is an inland salt lake basin. Under the closed, high-salinity, and strong-evaporation environment of the Qianjiang Formation of Paleogene in the Qianjiang Depression, thousands of meters of thick salt-bearing strata are deposited, and hundreds of salt cyclothem (a sedimentary cycle is a rhythm) have developed. Each cyclothem consists of upper and lower salt rocks and a set of carbon-rich laminated argillaceous dolomite, dolomitic mudstone, and calcium-glauberite mudstone (or calcium glauberite rock filled with cloud mudstone) strata in between. Each salt cyclothem generally has an approximate thickness of 5 to 12 m and can sometimes reach up to 20 m. The formation is composed of hydrocarbon source and reservoir layers, which are blocked by upper and lower salt rock to form an intersalt dolomitic shale oil reservoir.

The intersalt dolomitic shale oil reservoir is a high-quality hydrocarbon source layer integrally, which is composed mainly of argillaceous dolomite facies followed by dolomitic mudstone [15]. Based on fluid compartment theory, the formation water in the intersalt shale reservoir is primary water with no external water interference [16, 17]. At the same time, salt minerals accumulate in the cracks and pore channels for over a long period of time under high-temperature and high-pressure conditions, resulting in no formation water, low reservoir permeability, and good sealing. On-site production data also indicate that wells without water injection measures have no produced water after production.

The object of the study is the intersalt dolomitic shale of Qian 34-10 cyclothem in the Qianjiang Depression of the Jianghan Basin. Figure 1 presents the results of mineral composition analysis. The intersalt strata are primarily composed of mud, salt, and carbonate minerals. The upper part of the cyclothem is carbon-rich laminated argillaceous dolomite facies. The dolomite content is 12.8% to 78.2% with an average of 50.7%. The average total content of carbonate minerals in argillaceous dolomite is 65.2%. The middle part is carbon-rich laminated dolomite or calcareous mudstone facies. The argillaceous mineral content is relatively high, mainly feldspar minerals, which account for approximately 12.5% to 48.7% with an average of 44.5%, followed by carbonate minerals, with an average of 40.2%. The bottom of the cyclothem is mainly composed of glauberite mudstone facies, which consist of argillaceous, salt (mainly sodium chloride and glauberite), and carbonate minerals with indistinguishable contents. In general, the intersalt dolomite shale oil reservoir has a low clay mineral content (mainly chlorite and illite with an average content of 10.3%) that shows characteristics of low clay minerals, low quartz, high sodium feldspar, and high carbon. In addition, the 34-10 cyclothem of the intersalt dolomite shale oil reservoir has overall salt-bearing properties. The salt content of the upper and middle argillaceous dolomite and dolomite mudstones is relatively low, ranging from 2.5% to 9.1%. The glauberite mudstone near the bottom of the cyclothem layer has a relatively high salt content of up to 46.6%.

The physical properties of the different lithofacies of Qian 34-10 cyclothem in the intersalt dolomite shale reservoir have obvious differences (Table 1). Based on the lithology statistics of several typical wells in the exploration area, the porosity distribution of mud dolomite is mostly between 6% and 15%, and the permeability of around 45% of the sample is below 0.5 × 10⁻³ μm². The porosity of dolomite mudstone is less than 25%, with 50% of the samples concentrated in 9 to 12%, and the permeability is basically below 0.5 × 10⁻³ μm². The porosity of the glauberite mudstone varies widely (between 0% and 20%), and the permeability is concentrated below 0.5 × 10⁻³ μm², as shown in Figure 2. Mercury intrusion experiments show that the upper
carbon-rich laminated argillaceous dolomite facies (thickness of approximately 3–4 m) has a high average pore throat radius (>90 nm) and relatively good physical properties. The dolomite mudstone facies in the middle and lower parts of the reservoir have poor physical properties, the median pressure of the mercury injection is high, and the average pore throat radius is less than 40 nm.

### Table 1: Average pore throat radius of different lithofacies of Qian 34-10 cyclothem in the inter salts dolomite shale reservoir.

<table>
<thead>
<tr>
<th>Lithofacies</th>
<th>Average median pressure (MPa)</th>
<th>Average median pore throat radius (nm)</th>
<th>Average pore throat radius (nm)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Argillaceous dolomite</td>
<td>5.21</td>
<td>87.1</td>
<td>94</td>
</tr>
<tr>
<td>Dolomite mudstone</td>
<td>22.9</td>
<td>36.7</td>
<td>37.9</td>
</tr>
<tr>
<td>Glauberite mudstone</td>
<td>9.42</td>
<td>71.3</td>
<td>67.2</td>
</tr>
</tbody>
</table>

### Figure 1: Sedimentary facies histogram of Qian 34-10 cyclothem in the inter salts dolomite shale reservoir.

### 3. Microscopic Pore Structure and Salt Characteristics

Scanning electron microscopy observations were performed on the pores using a Zeiss focused ion beam scanning helium ion microscope. The Gatan 691 CS argon ion thinning instrument was used for argon ion polished sampling.
The main steps include sample pregrinding, high-energy argon ion thinning pretreatment, surface deposition of a conductive film, gold coating treatment, and electron microscope observation. The surface of the argon ion polished sample is smooth, and the backscattered electron imaging used is suitable for observing the microstructure and morphology of the nanopores.

3.1. Microscopic Pore Structure. Dolomites are widely distributed in argillaceous dolomite, often with rhombohedral crystals, and the grains are mostly less than 5 μm in size. It mainly develops dolomite and glauberite intergranular pores. The diameter of the dolomite intergranular pores is generally 0.5–10 μm, which are large intergranular pores supported by dolomite mineral lattice. The average pore diameter is nearly 2 μm, and the surface pore rate can reach 7.8%, as shown in Figure 3(a). The dolomite intergranular pores are one of the important oil storage spaces in the intersalt shale oil reservoir. Dissolution pores are generally distributed with a pore diameter of approximately 0.5–9 μm and a dissolution surface pore rate of around 5%–20%. They develop near the bedding plane and are also the main oil storage space. Most of the dissolution pores are effective pores, but some are filled with a large amount of salt particles, such as rock salt or glauberite, as shown in Figure 3(b). Second, a small number of intragranular pores are developed, including dolomite and quartz intragranular pores. The pore size is generally 5–200 nm with an average of around 1 μm, and the surface pore rate can reach up to 4.2%. The dolomitic mud shale also develops intergranular pores of clustered pyrite, bedding joints, and diagenetic shrinkage joints, as shown in Figures 3(c) and 3(d).

The dolomite mudstone mainly develops intergranular pores of clay minerals and dolomite. The pore diameter is generally 0.5–6 μm with an average of around 1 μm, and the surface pore rate can reach up to 6.1%. Rock salt particles are scattered on the pore surface. The rock salt crystals have a good original shape and are tetragonal. The length of the single particles is slightly shorter at around 200 nm. Clay minerals and dolomite intragranular pores are developed. The
pore size is generally 5–200 nm with an average of around 50 nm, and the surface pore rate can reach up to 3.1%. Compared with argillaceous dolomite, the dolomite mudstone pores are less developed, and the pore size is relatively small, as shown in Figure 3(e).

In addition, a small number of organic pores that developed inside the pyrolysis asphalt are found in the dolomite mudstone. The pore morphology is mostly irregular, bubble-like, and elliptical. The pore size is 5–300 nm, with an average of around 100 nm, and the surface pore rate can reach up to 15%, as shown in Figure 3(f). Secondary mineral particles with complete crystallization are usually developed in the edge or interior of the organic pores, and the pore volume usually accounts for 20%–50% of the total volume of organic particles with good pore connectivity.

### 3.2. Soluble Salt Minerals

The electron microscopy scanning results showed that the intersalt reservoir has microscopic salinity characteristics. Soluble salt minerals are widely distributed in the pores and cracks. The microscopic morphology of the reservoir salt particles consists of a single salt particle and a collection of salt particles. The intergranular pore size of the salt particles is less than 300 nm. The salt aggregates can completely fill the pores and also block the cracks, resulting in poor pore connectivity of the reservoir and hindering fluid flow in the reservoir.

Moreover, the intersalt cloud shale also has macroscopic salinity characteristics. As can be seen by the naked eyes, the core rock sample is mixed with unequal rock salt and glauberite particles, and some cracks are also filled with salt minerals. The salt characteristics of shale samples are quantitatively described by micro-CT scanning technique. The results show that in the Qian34-10 cyclothem, few salt particles are present in the argillaceous dolomite layer and shale in the middle dolomite mudstone layer, mainly distributed between 1.3% and 3.8%, as shown in Figures 4(a) and 4(b). At the same time, the cracks are partially filled with sodium chloride particles and glauberite minerals. The glauberite mudstone near the salt layer has high salt content, and the volume ratio of the sodium chloride and glauberite particles is around 15.7%–21.81%, as shown in Figure 4(c).

Fracturing practice shows that the salinity of the flowback fluid of the shale reservoir is generally high [18, 19]. With the extension of the backflow time, the salinity of the flowback fluid keeps increasing, reaching up to $10 \times 10^4$ mg L$^{-1}$, whereas the salinity of the slick water injected during fracturing is very low (approximately 1000 mg L$^{-1}$) [18, 19]. The salt ions in the flowback fluid are mainly derived from the intersalt clouds.
dissolution of the shale’s own minerals and the crystalline salts of the pore walls. When the fracturing fluid with low salinity enters the reservoir, the large difference in salinity between the fracturing fluid and the formation results in a remarkable chemical potential difference between them, which becomes the driving force for the low salinity liquid to enter the shale interior [20–22]. As for the intersalt dolomite shale reservoir, the interior has microscopic and macroscopic salinity properties, and the high-salinity difference will lead to a high chemical potential difference. After the fracturing fluid with low salinity infiltrates into the reservoir, it dissolves the salt minerals inside the reservoir continuously. Superficially, if the matrix and soluble salts in the fractures of the shale reservoir are dissolved, additional circulation channels can be obtained for reservoir oil and gas. This is also one of the main reasons for using low-viscosity slick water fracturing or water injection to increase shale oil production. However, further studies are needed to determine the specific effects of dissolution, migration, and recrystallization of soluble salt minerals on the physical properties of a reservoir.

4. Effects of Different Fluid Types

4.1. Sample Collection and Processing. The core samples are taken from the downhole cores of the Qian3–10 cyclothem at varying depths in the Qianjiang Depression, Jianghan Basin. As the intersalt shale is a typical stratified rock mass, the bedding surface is prone to open when exposed to water. However, factors such as the water solubility of salt rock can also contribute to this problem. Therefore, to avoid the opening of the bedding surface and dissolution of soluble minerals within the rock caused by hydraulic cutting, all samples in the experiment undergo anhydrous processing, and the samples are processed into a standard core column with a diameter of 2.54 cm. Then, the side of the core is wrapped and sealed with epoxy resin, and only the two end faces of the core are in contact with the fluid when immersed.

4.2. Methods of Experiment. The shale sample is preplaced in a high-temperature and high-pressure reaction tank with a volume of 157 mL and evacuated. Then, fluid is injected into the reaction tank until the pressure becomes constant at 20 MPa, and the temperature in the incubator is kept constant at 80°C. The experimental immersion fluid includes distilled water, CO₂ (CO₂ is in a supercritical state at 80°C and 20 MPa), and a liquid mixture of CO₂ and distilled water (supercritical CO₂ saturated solution). When the immersion fluid medium is a supercritical CO₂ saturated solution, the immersion step differs from distilled water or pure CO₂. First, 100 mL of distilled water is injected into the reaction tank. The shale sample is placed in an aqueous phase, evacuated, and heated at a constant temperature of 80°C. Thereafter, the high-temperature and high-pressure reaction tank is charged with CO₂ until the fluid pressure in the tank reaches 20 MPa.

After the shale samples are immersed under high-temperature and high-pressure conditions, the porosity and permeability of different shale samples are measured by using helium porosimetry and ultralow permeability meter (pulse attenuation method). Moreover, helium is used as the test gas.

4.3. Solubility of CO₂. When supercritical CO₂ is mixed with water, hydrogen ions are formed through ionization, which makes the solution acidic, as shown in formula (1). Related literatures show that the higher the pressure, the higher the solubility of CO₂, the more hydrogen ions are ionized in the solution, and the lower the pH of the solution [8, 23, 24]. At the same time, the solubility of CO₂ in water at different temperatures under 20 MPa is measured, as shown in Figure 5. As seen in the figure, with the increase in temperature, the solubility of CO₂ in water decreases and tends to be stable when the temperature is greater than 100°C:

\[ \text{CO}_2 + \text{H}_2\text{O} \rightarrow \text{H}^+ + \text{HCO}_3^- \]  \hspace{1cm} (1)
For intersalt shale reservoirs, active H\(^+\) and HCO\(_3^-\) interact with carbonates, such as calcite and dolomite, to dissolve carbonate rocks:

\[
\text{CaCO}_3 + \text{H}^+ \rightarrow \text{Ca}^{2+} + \text{HCO}_3^- \tag{2}
\]

\[
\text{CaMg(CO}_3)_2 + 2\text{H}^+ \rightarrow \text{Ca}^{2+} + \text{Mg}^{2+} + 2\text{HCO}_3^- \tag{3}
\]

H\(^+\) and HCO\(_3^-\) can also interact with the feldspar minerals inside the reservoir. Aside from causing dissolution and sedimentation of silicate, they can also form kaolinite:

\[
2\text{KAlSi}_3\text{O}_8 + 2\text{H}^+ + 9\text{H}_2\text{O} \rightarrow 2\text{K}^+ + 4\text{H}_4\text{SiO}_4 + \text{Al}_2\text{SiO}_5(\text{OH})_4 \tag{4}
\]

\[
\text{NaAlSi}_3\text{O}_3 + \text{CO}_2 + 5.5\text{H}_2\text{O} \rightarrow \text{Na}^+ + \text{HCO}_3^- + 2\text{H}_4\text{SiO}_4 + 0.5\text{Al}_2\text{SiO}_5(\text{OH})_4 \tag{5}
\]

Therefore, the interaction of CO\(_2\)-water-reservoir rocks will lead to the dissolution of reservoir rocks and the formation of sediments, which will directly result in the change of pore structure and permeability of reservoirs.

4.4. Results and Analysis of the Experiment

4.4.1. Changes in the Microscopic Pore Structure. Under the conditions of 80°C and 20 MPa, the shale samples are immersed in distilled water, supercritical CO\(_2\), and supercritical CO\(_2\) saturated solution. Then, microscopic pore scanning electron microscopy is conducted to compare the microstructure changes in the intersalt cloud shale before and after the fluid action. The results are shown in Figure 6.

For argillaceous dolomite samples, some of the pores are filled with many salt crystals without treatment, as shown in Figure 3(a). The filled salt minerals occupy most of the fluid passages, hindering fluid migration. After immersing in supercritical CO\(_2\), the surface pore rate reaches 9.1%, as shown in Figure 6(a). This result may be due to the expansion...
of CO₂. On the one hand, the pore throat radius increases and the effective stress decreases due to the volume expansion of CO₂. On the other hand, the CO₂ expansion affects and disperses the collection of salt particles in the pore channels into salt particles, thereby dredging pores and reducing the flow resistance of the reservoir fluid. However, under the reservoir conditions, the salt minerals still remain in the reservoir and occupy the flow channels of the reservoir fluid.

When immersed in distilled water, the salt crystals in the pores dissolve, and the pore channels expand, significantly increasing the surface pore rate (around 40%) and greatly improving pore connectivity, as shown in Figure 6(b). However, distilled water weakens the cohesive force between the rock mineral particles, which loosens the rocks and even produces microcracks. In addition, the clay mineral undergoes hydration expansion under the action of water, as shown in Figure 6(c). However, due to the low clay mineral content of the inter_salt shale reservoir, the hydration of clay minerals has minimal effect on the shale skeleton or pore structure.

When the shale sample is immersed in a supercritical CO₂ saturated solution, the saturated solution has the greatest impact on the micropore structure of shale. First, distilled water can dissolve salt, and the pore channels are enlarged when salt crystals dissolve inside the pore channels. In addition, when supercritical CO₂ dissolves in distilled water, it will produce a weak acid that will dissolve the cement and mineral components of the rock particles, thereby greatly increasing the pore radius and even producing the dominant channel of the acid-rock reaction. The surface pore rate can reach up to 23.6%, as shown in Figure 6(d). Second, the weak acid solution will dissolve the skeleton particles, resulting in dissolved pores with a diameter of 50–100 nm inside the

![Graphs showing the change of porosity after immersion in fluid at different times.](image)
particles, as shown in Figure 6(e). The saturated solution destroys the stability of the clay minerals and causes the dissolution of the cement of the skeleton particle, leading to the increase of intergranular pores and considerable migration of clay particles, as shown in Figure 6(f).

4.4.2. Change of Porosity. Figure 7 shows the results of porosity change of the shale samples after immersing in supercritical CO₂, distilled water, and supercritical CO₂ saturated solution at 20 MPa and 80°C.

The results of the experiment show that the porosity increased at varying degrees after immersing the samples in different fluids. The supercritical CO₂ saturated solution obtained the largest increase in porosity. The salt solution effect of distilled water combined with the dissolution effect of the weak acid generated by the dissolution of supercritical CO₂ in water will greatly expand the pore and increase the pore volume. The porosity of argillaceous dolomite and glauberite mudstone with high total content of salt minerals and carbonates increased by 210.7% and 296.7%, respectively, after immersion for 168 h. The pore-enlarging effect of distilled water was mainly derived from the dissolution of salt minerals. The effect was most obvious in the glauberite mudstone with high salt content. The porosity of argillaceous dolomite and dolomite mudstone with lower salt mineral content increased by 43% and 31.9%, respectively, after immersion in distilled water for 168 h. By contrast, the pore-enlarging effect of supercritical CO₂ was relatively small, and the porosity of different samples increased by approximately 6.0% to 13.2% after 168 h immersion.

4.4.3. Change of Permeability. After immersion in different fluids, an ultralow permeability meter is used to measure the change in permeability of the shale samples over time. The results are shown in Figure 8.
As can be seen from the figure, the effect of supercritical CO₂ on permeability improvement is not obvious. For the argillaceous dolomite samples with relatively good porosity and permeability, the permeability increases by 13.7% through flushing the salt crystal minerals and dredging the pore channels. For the glauberite mudstone samples, the permeability only changes by 4.2% due to its small internal pores that are mostly blocked by salt minerals. The permeability of distilled water is greatly improved by dissolving the salt minerals in the pores. After immersion for 168 h, the permeabilities of the argillaceous dolomite and dolomite samples increase by 97.6% and 46.2%, respectively. By contrast, the permeability of glauberite mudstone samples with high salt minerals increases by 14.7% after immersion for 168 h. Therefore, the better the initial porosity and permeability conditions, the greater the salt content and the increase in permeability. When the supercritical CO₂ is dissolved in distilled water, the weak acid dissolves and expands the pores, which then improves interporosity connectivity, enlarges the range of water-salt-rock action, and strengthens the salt solution of distilled water. The permeability of the glauberite mudstone minerals increases by 21.1 times after immersion in supercritical CO₂ saturated solution for 168 h.

5. Effects of Recrystallization of Salt Minerals

During the mining process, due to changes of formation temperature and pressure in and near the wellbore, the solubility of the salt minerals changes, causing the precipitation of salt minerals and salt crystallization. Then, salt crystallization will lead to damage near the wellbore reservoir. The static evaporation salting-out method is used to simulate the salt crystallization phenomenon in the formation so as to evaluate the damage degree of salt crystallization on the porosity and permeability of the intersalt shale oil reservoir. The specific steps are as follows: (1) under the conditions of 80°C and 20 MPa, the shale sample is placed in a saturated sodium chloride solution for 24 h using a high-temperature and high-pressure reactor; (2) the shale immersed in the saturated sodium chloride solution is heated in an oven at a constant temperature 60°C and evaporated to constant weight; and (3) the morphology and microscopic

![Figure 9: Pore microstructure changes after static evaporation salting out.](image-url)
distribution of salt crystals are observed under scanning electron microscopy to determine the effect on the porosity and permeability of the inter shale.

5.1. Morphology and Microscopic Distribution of Salt Crystal. Scanning electron microscopy observation of salt shale after salt crystallization shows three main forms of salt crystals, including clumps of superimposed growth morphology of intergranular pore/seam filling type, surface-attached layered or flocculent growth morphology, and single-particle dispersion morphology, as shown in Figure 9.

The filling-type agglomerated crystalline salt usually aggregates in the dominant channel with good evaporation environment. These channels generally have large pore size and good connectivity because the fluid saturation in the large pores is usually high and the large pore size also provides a good evaporation environment for the liquid [25–27]. The salt crystals are stably stacked in the large pores and gradually fill the entire pores. However, these filled salt crystals are densely packed and have no internal permeability, as shown in Figure 9(b). Therefore, the filling of the cluster salt crystals may partially cause the dominant pores to fail, as shown in Figure 9(c). At the corners of pores or on the surfaces of rocks, scattered regular single-crystal crystalline salts with a length of tens to hundreds of nanometers are present, as shown in Figures 9(a) and 9(d). Given the small particle size, such salt crystals tend to migrate and block small pores.

When the flow and evaporation environments of the tunnel are relatively poor, irregular flocculated or layered crystals attached to the surface of rock particles are generated due to unstable salt crystal deposition (Figures 9(d)–9(f)). On the one hand, the layered adhesion and flocculation of the salt crystals will make the pore flow channel narrower. On the other hand, the flocculent salt crystals are easy to disperse, accumulate, and block the pore due to the lack of framework support.

5.2. Influence on Porosity and Permeability. The microscopic scanning results show that salt crystallization will occupy a certain reservoir pore space and block the shale pore channels. By measuring the changes of shale porosity and permeability before and after static evaporation salting-out experiments, the salt crystals will lead to the decrease of porosity and permeability of shale. The average porosity and permeability decreased by 17.8% and 37.5%, respectively, as shown in Figure 10. The smaller the porosity and permeability of the shale, the greater the damage of salt crystals to the porosity and permeability. This is due to the small pore throat of rocks with low permeability, most of which contain small pores or micropores. These micropores are easily blocked by fine salt crystals, so the permeability is drastically lowered. For rocks with high permeability, most of the pore channels are relatively large and are not easily blocked by salt crystals. In addition, the dispersed crystalline salt particles are more likely to migrate in cores with better permeability [26–28]. Therefore, the permeability of the shale sample with high permeability will slightly decrease after the salt crystallizes.

6. Discussion

With low water saturation and high salt content in pores, the formation water in the inter salt dolomite shale reservoir mainly exists in the form of salt crystal water. The salt crystals will block the pore channels, reduce the connectivity between pores, and increase flow resistance. Water-based fracturing fluids can greatly increase the porosity and permeability of reservoirs by dissolving the salt minerals in the matrix. The salt minerals are dissolved after contact with water, especially in the glauberite mudstone layer near the salt layer. Then, the rock porosity and permeability will be greatly improved. When the salt minerals are dissolved, they will migrate to the fracture and bottom of the well in the form of high-salinity water. Under the influence of temperature and pressure, high-salinity water recrystallizes and gradually blocks the matrix, cracks, and wellbore. At this point, the water extrusion process can be adopted to restore the production capacity of the reservoir by dissolving the salt minerals. However, the water injected in a later stage continuously acts on the glauberite mudstone layer and strengthens its permeability, thereby providing unfavorable conditions for the dissolution of the salt layer at the top of the reservoir and the migration of many salt minerals.
Supercritical CO₂ fracturing can avoid the problems caused by the dissolution, migration, and recrystallization of salt minerals due to water-based fracturing fluids. At the same time, supercritical CO₂ fracturing increases the complexity of hydraulic fractures and improves the fluidity of the mixed phase. However, the results of microscopic scanning observation and pore infiltration parameter test show that the salt content and pore connectivity of intersalt shale reservoirs are poor. Although supercritical CO₂ can disperse the salt particles, dredge the channels, and enlarge the pores due to expansion, its overall ability to change the structure and permeability of the pores is still weak. Therefore, it cannot meet the requirements of economic development for transformation effect by only using the supercritical CO₂ for the transformation of intersalt shale reservoirs.

At present, the fracturing technology combining CO₂ with water-based fracturing fluid has become an emerging technology for unconventional oil and gas development. For the intersalt shale oil reservoir, the supercritical CO₂ composite fracturing technology can reduce the amount of water-based fracturing fluid and the considerable dissolution and migration of salt minerals. At the same time, after the CO₂-water-shale interaction, the pore circulation channel is enlarged, and porosity and permeability are improved. To a certain extent, the damage of salt crystals to the physical properties of the reservoir near the wellbore is reduced. Therefore, the CO₂ composite fracturing technology combines the salt dissolution effect of water-based fracturing fluid and the unique effect in production increase of CO₂, which may cause effective long-term development of intersalt shale oil reservoirs. However, further research needs to be conducted on how to optimize the amount and proportion of CO₂ and water-based fracturing fluid based on the physical damage law of the reservoir under the interaction of fluid-rock-salt in the intersalt dolomite shale reservoir.

7. Conclusion

(1) The intersalt dolomite shale reservoir has microscopic and macroscopic saliferous characteristics, which lead to poor pore connectivity of the reservoir and hinder the flow of reservoir fluid. Distilled water can enlarge the flow passage by dissolving soluble salt minerals inside the reservoir. However, the dissolution and migration of many salt crystals will cause problems for later production.

(2) During the mining process, salt minerals are precipitated, and salt crystals are formed due to changes in formation temperature and pressure in the wellbore and near the well. Salt crystals will occupy a certain pore space and block the pore channel. The smaller the porosity of the shale and the lower the permeability, the greater the damage of salt crystals to the pores.

(3) Supercritical CO₂ disperses the salt particles, dredges the channels, and slightly enlarges the pores by expansion, but the overall ability to change the pore structure and connectivity is still weak. By using supercritical CO₂ to transform the intersalt shale reservoir, the problems caused by salt dissolution and recrystallization can be avoided. However, supercritical CO₂ cannot meet the requirement of economic development for the transformation effect.

(4) When supercritical CO₂ is dissolved in distilled water, it will produce a weak acid, which erodes the cement and the mineral components of the rock particles, greatly increasing the radius of the channel. At the same time, by reducing the amount of water-based fracturing fluid and increasing the porosity and permeability of the matrix, the cycle of salt plugging caused by the massive migration of salt minerals will be reduced, which is conducive to extending the effective period of increased production.

Data Availability

The data used to support this study are included within the article.

Conflicts of Interest

The authors declare that they have no conflicts of interest.

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