

Research Article

Sensitivity Analysis of Factors Affecting Time-Dependent Slope Stability under Freeze-Thaw Cycles

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Received 23 January 2018; Revised 14 April 2018; Accepted 8 May 2018; Published 4 June 2018

Academic Editor: Marek Lefik

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With open-pit mines excavated, the slopes will be exposed to the natural environment for a long time. Affected by factors like temperature, seepage, mining, freeze-thaw, etc., slope structural integrity and strength will gradually decline as slope exposure time extends. Besides, the development of defect structure within the rocks is closely correlated with time. In this paper, freeze-thaw cycle tests were conducted on the saturated sandstones collected from a certain open-pit mine. According to the test results, the mass density and longitudinal wave velocity gradually increased with more times of freeze-thaw cycles while mechanical properties such as internal friction angle, cohesion, elastic modulus, and uniaxial compressive strength decreased instead. The constitutive model of saturated rock deterioration was established by taking the volume of phase transition of water in microcracks as a variable. Based on the tests results and theoretical analysis, the sensitivity of the factors affecting slope stability under freeze-thaw damage was studied by Control Variable Method (CVM) and Orthogonal Design Method (ODM). It was determined that the internal friction angle and cohesion had a highly significant effect on the test results, while the mass density had a significant effect. The conclusions may play a certain role in guiding slope construction and protection.

1. Introduction

Rock is typical of highly heterogeneous material and there exist various microdefects in different scales within it. And their formation, expansion, and perforation have a great influence on rock deformation and even failure. The development of internal defect structure in rock is closely correlated with time. During the whole formation and evolution of rock engineering, a variety of environmental factors may lead to the gradual deterioration of mechanical properties. For years, many scholars have devoted themselves to the study of time-dependent characteristics of rock strength gradually weakening over time. A. Dragon et al. (1979) studied the brittle and plastic failure of rock according to the fracture surface and established corresponding time-dependent continuum model[1]. M. A. Kaehanov (1982) studied the failure mechanism of rock materials according to the structure and

established corresponding damage constitutive model [2]. J. Kemeny (2003) derived a fracture mechanics model by utilizing subcritical crack growth, which resulted in a closed-form solution for joint cohesion as a function of time [3]. In [4], the creep properties of rock were simulated from the point of view that the mechanical properties of materials deteriorate with time. Having considered slope exposure time and service cycle, Q. X. Cai et al. (2008) proposed a theory of time-dependent slope, which has been successfully applied to steep slope mining practice [5]. This basic research shows that the time-dependent characteristics, namely, internal rock structure gradually weakening over time, play a crucial role in transforming microslope damage into macroinstability.

Open-pit coal mining is a complex and systematic engineering. Generally speaking, it takes years, even decades of service period, to complete the surface excavation, slope outcropping, and the inner dumping and reclamation. Exposed

to natural environment, slopes are easily affected by temperature, seepage, mining, freeze-thaw, and other factors. And as the exposure time extends, their structural integrity and strength gradually decline [6]. For open-pit mines located in seasonally frozen regions, the exposed soil and rock mass freeze in winter and thaw in spring. And such alternation of freezing and thawing brings about radical changes to their physical and mechanical properties. As a result, the long-term stability of slope engineering is directly compromised.

The slope stability of open-pit coal mines under normal temperature conditions has been studied by many scholars and lots of results have been obtained [7–10]. However, as to freeze-thaw related slope disaster and its prevention in seasonally frozen regions, quite limited research has been done so far. As a matter of fact, certain geotechnical engineering research has been conducted under the condition of freeze-thaw cycle and it is mainly for the sake of regional development and resources exploitation in permafrost regions of high elevation or high latitude or in polar regions. To study the slope stability of opencast coal mines under freeze-thaw cycle effect, mature theories and methods such as slope stability evaluation and numerical simulation were quoted herein to analyze the variation pattern of physical and mechanical characteristics of soil and rock mass. On this basis, slope parameters and support programs were further optimized. M. G. Ferrick et al. (2010) completed three groups of experiments, with six experiments in each group, to quantify the differences in terms of soil erosion and rill development following a single freeze-thaw cycle [11]. K. M. Skarzynska (1985) conducted a series of experiments on fine-grained soil under freeze-thaw cycle effect. The experiment results showed that the soil structure, density, and porosity change under such effect, which brings about water redistribution [12]. K. D. Eigenbrod (1995) and J. Graham (2011) found that the strength, compressibility, and pore water pressure of the fine-grained soil were affected in the case of soil melting [13, 14]. Besides, by simulating rock freeze-thaw process, L. Wen et al. (2015) studied the changes of mass, volume, uniaxial compressive strength, tensile strength, freeze-thaw coefficient, and weathering degree of saturated hard rock under freeze-thaw cycle effect. In addition, the factor of safety (Fos) before and after going through such cycles was calculated by using strength reduction method [15]. In combination with theoretical research and numerical analysis in [16], a three-field coupling of stress field, temperature field, and water field was carried out with FLAC3D to simulate the influence that freeze-thaw cycles exert on the slope stability of open-pit copper mines. As a matter of fact, the fracture procedure is a whole process of crack initiation, crack propagation, and final fracture. P. Valko et al. (1994) studied the mechanism of crack propagation resulting from frost heave in heterogeneous rocks [17] by using continuum damage mechanics.

In this paper, freeze-thaw cycle tests were conducted on sandstone collected from the slope of a certain open-pit mine. And they are designed to study the variation pattern, the damage and deterioration mechanism of physical and mechanical properties. Specifically, on the basis of ODM and FLAC3D, sensitivity analysis was conducted on the main

controlling factors of the landslide. Meanwhile, the factors were also compared to determine their order of importance and correlation so as to provide theoretical basis for the safe construction and landslide protection of freeze-thaw slopes.

2. Properties of Saturated Rock in the Condition of Freeze-Thaw Cycles

The drilling cores of the sandstone herein were taken from a slope of an open-pit coal mine in Xinjiang province (a seasonally frozen region in northwest China). The samples were processed into standard cylinders of 50 mm diameter and 100 mm height. Considering the local climate and the service period of the open-pit mine, a total of 24 rock samples were divided into four groups, with six samples in each group. The four groups of samples were set to go through 0, 10, 20, and 30 times of freeze-thaw cycles, respectively. At the beginning of the test, a group of samples was randomly selected and soaked for 48 hours at standard atmospheric pressure before measuring their mass density and geometry. The measurement was followed by triaxial compression tests at a constant displacement loading rate of 2×10^{-3} mm/s. The confining pressures of the two rock samples were 0 and the other four rock samples were provided with two confining pressure levels. The frictional stress circle was plotted according to the maximum principal stress under different confining pressures, and the corresponding internal friction angle and cohesion were obtained and recorded. At the same time, the remaining samples were placed in a closed system equipped with two sets of independent refrigeration systems, whose control precision and display precision were, respectively, 0.5°C and 0.1°C, as shown in Figures 1(a) and 1(b). A freeze-thaw cycle was set as 24 hours, of which 14 hours was at a low-temperature of -20°C so as to ensure the complete freezing of the samples. During the rest of 10 hours, however, they were put in pure water of +20°C to thaw out. Similarly, the properties tests and triaxial compression tests were performed on the other three groups of samples that had experienced 10, 20, and 30 times of freeze-thaw cycles, as mentioned above. And the testing results were recorded accordingly, as shown in Table 1.

(1) The tests showed that, with freeze-thaw cycle times increasing, certain sandstone properties began to change. For instance, the uniaxial compressive strength, Young's modulus, and cohesion obviously decreased as a result. In addition, since the rock samples were saturated when their mass was measured, it is reasonable to conclude that the rock mass increase was caused by water mass increase amongst the cracks. In other words, the space in the microcracks enlarged and the microcracks rate picked up accordingly. It indicated that freeze-thaw cycles resulted in the expansion and coalescence of microcracks in rock samples, which in turn led to internal structural destruction. At the same time, the proportion of effective rock per unit area decreased with gradual microcrack expansion; i.e., the effective stress-bearing area decreased as well, thus lowering the ultimate bearing capacity of the rock.



FIGURE 1: (a) Refrigeration systems; (b) temperature control system.

TABLE 1: Properties of sandstone with different times of freeze-thaw cycles.

Freeze-thaw cycles	Sample mass/(g)	Wave velocity/(m/s)	Compressive strength/(Mpa)	Young's modulus/(GPa)	Poisson's ratio	Cohesion/(MPa)	Angle of internal friction/(°C)
0 times	440.79	5140	138.68	8.3	0.214	28.6	25.6
10 times	440.86	5256	126.6	7.8	0.215	25.07	25.2
20 times	440.95	5441	118.87	7.2	0.214	24.55	24.6
30 times	441.01	5492	115.67	6.7	0.212	23.56	23.6

(2) According to the tests, there existed no visible natural spalling after every 10 times of freeze-thaw cycles. However, as can be seen from Table 1, the average mass density of saturated rock samples gradually enlarged with more times of freeze-thaw cycles. Obviously, affected by frost heave action, water freezing led to the microcrack growth. However, in the process of ice melting, the expanded microcrack space was filled with new water, which resulted in the increase of both water content and mass density of rock samples. In addition, the test data showed that the longitudinal wave velocity increased with more times of freeze-thaw cycles. As explained earlier, the closed cracks or unsaturated cracks in the rock sample have to go through frost heave. Considering the fact that wave propagates faster in water than in air, it is reasonable to deduce that it is because the air in closed cracks is replaced by water that not only the mass of the rock samples increased, but also the longitudinal wave velocity increased. They are positively related to each other, as shown in Figure 2.

(3) It should be noted that after 30 times of freeze-thaw cycles, the most remarkable changes in properties of sandstone were Young's modulus, cohesion, and uniaxial compressive strength, with a reduction of 19.3%, 17.6%, and 16.6%, respectively. The internal friction angle increased by 7.8% at maximum and the Poisson's ratio hardly changed.

3. Mechanism of Freeze-Thaw Damage for Saturated Rock

The freeze-thaw cycle is actually the phase transition process of water amongst cracks. To be specific, when crack water is

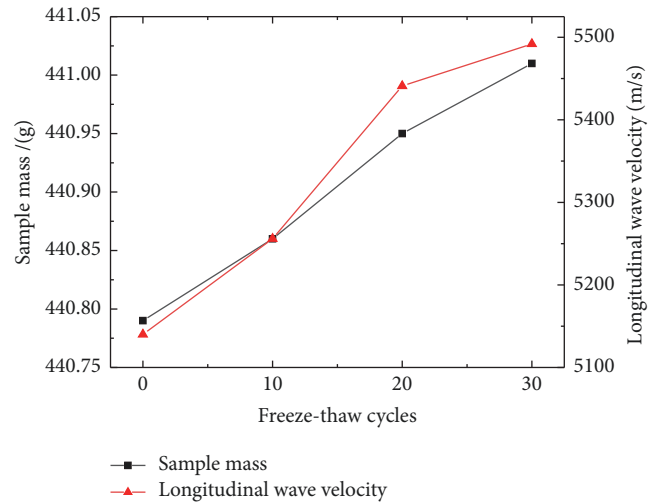


FIGURE 2: Mass and longitudinal wave velocity under the condition of freeze-thaw cycles.

frozen into ice, the volume expansion leads to the extension of microcracks. And once it melts into water again, it would in turn fill up the newly expanded cracks. In this sense, long-term freeze-thaw cycles are sure to bring about continuous crack expansion, porosity increase, rock damage accumulation, and, finally, its mechanical strength decrease. With such influence taken into account, a constitutive model of rock deterioration is established herein by taking the phase transition volume of the water in microcracks as a variable.

When the temperature drops below freezing point, the water in cracks will be frozen into ice. The microcracks in the rock mass were assumed to be uniformly distributed and the constraints of crack walls on volume expansion were ignored. In this case, the frozen volume of water in microcracks can be expressed as follows:

$$\Delta V_i^F = \beta V_w \mu^T \quad (1)$$

where ΔV_i^F is the frozen volume of water in microcracks at standard atmospheric pressure, β is the frost heaving coefficient of water at standard atmospheric pressure, V_w is the volume of water in saturated microcracks, and μ^T is the freezing rate of water at T degrees centigrade.

When the water in saturated microcracks is frozen into ice, the crack wall will exert inward compressive stress on the ice while being subjected to outward expansion force. At the same time, with temperature lowering, the rock contracts and the resulting frozen contraction force would impose on the ice as well. Therefore, in addition to the free expansion, the ice volume would experience certain compressive deformation under the action of composite forces mentioned above. It is assumed that rock and ice are isotropic elastic media and the following equation can be obtained:

$$\Delta V_i^P = (\varepsilon_x + \varepsilon_y + \varepsilon_z) V_w = \frac{1 - 2\nu_i}{E_i} (\sigma_x + \sigma_y + \sigma_z) V_w \quad (2)$$

where ΔV_i^P is the volume compression of ice due to crack wall load, $\varepsilon_x, \varepsilon_y, \varepsilon_z$ are the strain of ice in the directions of x -, y - and z -axis, ν_i is Poisson's ratio of ice, E_i is Young's modulus of ice, and $\sigma_x, \sigma_y, \sigma_z$ are the stress of ice in the directions of x -, y -, and z -axis. Since the expansion force of ice is equal to the frozen contraction force of rock mass in every direction, (2) can be hence expressed as

$$\Delta V_i^P = \frac{3(1 - 2\nu_i)}{E_i} \sigma_{ri} V_w \quad (3)$$

where σ_{ri} is the composite force exerting on the ice amongst the microcracks in arbitrary directions, $\sigma_{ri} = \sigma_x = \sigma_y = \sigma_z$.

So, when the water in the cracks is frozen into ice, the actual volume expansion ΔV_i is

$$\Delta V_i = \Delta V_i^F - \Delta V_i^P = \left[\beta \mu^T - \frac{3(1 - 2\nu_i)}{E_i} \sigma_{ri} \right] V_w \quad (4)$$

When the ice expansion stress exceeds the rock fracture strength R_t , the volume of microcracks would expand. At the moment of rock cracking, the composite force σ_{ri} is equal to the fracture strength R_t in terms of their numerical value. So the expanded volume can be expressed as

$$V_i = V_w + \Delta V_i = \left\{ 1 + \left[\beta \mu^T - \frac{3(1 - 2\nu_i)}{E_i} R_t \right] \right\} V_w \quad (5)$$

where V_i is the microcrack volume after one freeze-thaw cycle.

K , the microcrack propagation factor, is defined as

$$K = \frac{V_w + \Delta V_i}{V_w} = 1 + \frac{\Delta V_i}{V_w} \quad (6)$$

$$= 1 + \left[\beta \mu^T - \frac{3(1 - 2\nu_i)}{E_i} R_t \right].$$

Then after N times of freeze-thaw cycles, K_n could be expressed as

$$K_n = \left\{ 1 + \left[\beta \mu^T - \frac{3(1 - 2\nu_i)}{E_i} R_t \right] \right\}^n \quad (7)$$

For saturated rock mass, as soon as the microcracks within extend, there would be water filling them up in time. Therefore, the volume of microcracks after N times of freeze-thaw cycles could be expressed as

$$V_{in} = \left\{ 1 + \left[\beta \mu^T - \frac{3(1 - 2\nu_i)}{E_i} R_t \right] \right\}^n V_w = K_n V_w \quad (8)$$

where V_{in} is the volume of microcracks within rock mass after N times of freeze-thaw cycles.

Meanwhile, the microcrack rate of rock mass can be expressed as

$$e_n = \frac{V_{in}}{V_0 + \Delta V_{in}} = \frac{K_n V_w}{V_w/e_0 + K_n V_w - V_w} \quad (9)$$

where V_0 is the initial volume of rock, e_n is the microcrack rate of rock mass after N times of freeze-thaw cycles, and e_0 is the initial microcrack rate. Based on (9), the following can be obtained:

$$e_n = \frac{K_n e_0}{(K_n - 1)e_0 + 1} \quad (10)$$

$$K_n = \frac{e_n(1 - e_0)}{e_0(1 - e_n)} \quad (11)$$

Once the rock mass goes through many times of frost heaving, its microcrack rate changes and such change would directly affect its mechanical strength. A microunit of rock mass was taken as a sample herein and its arbitrary section is composed of microcolumn and microcracks. It showed that after several times of free-thaw cycles, microcracks per unit area extended while the number of microcolumns reduced, which resulted in less bearing area of effective stress. The ultimate bearing capacity of per unit volume before and after freeze-thaw cycles can be expressed as

$$F_0 = R_c \Delta S_0 \quad (12)$$

$$F_n = R_c \Delta S_n \quad (13)$$

where F_0, F_n are the ultimate bearing capacities of the initial state and after N times of freeze-thaw cycles, respectively. $\Delta S_0, \Delta S_n$ are the bearing areas of effective stress in micro-rock-unit and R_c is the compressive strength of microrock column per

unit area in microunit. Based on these two equations, we can obtain

$$\frac{F_0}{F_n} = \frac{\Delta S_0}{\Delta S_n}. \quad (14)$$

It is assumed that the microrock column is rigid and there exists no extrusion deformation in the condition of frost heave. The microcracks expand and contract with the phase change of crack water. The section area of the micro-rock-unit after N times of freeze-thaw cycles is expressed as follows:

$$S_n = \Delta S_0 + \frac{\Delta S_0}{1 - e_0} e_0 K_n \quad (15)$$

where S_n is the section area of the micro-rock-unit after N times of freeze-thaw cycles.

According to the assumption mentioned above, there is no compression deformation for microrock column when the cracks extend. So the section area of the extended unit will be larger than that of the initial state. By calculation, the effective area of the microrock column per unit area after expansion can be obtained as follows:

$$\Delta S_n = \Delta S_0 \frac{S_0}{S_n} \quad (16)$$

where S_0 is the section area of the initial unit.

Therefore, the ratio of the effective stress-bearing areas in these two cases is

$$\begin{aligned} \frac{\Delta S_0}{\Delta S_n} &= \frac{S_n}{S_0} = \frac{((1 + e_0 (K_n - 1)) / (1 - e_0)) \Delta S_0}{\Delta S_0 / (1 - e_0)} \\ &= 1 + e_0 (K_n - 1). \end{aligned} \quad (17)$$

In combination with (7), (14), and (17), the relationship of ultimate bearing capacity in micro-rock-unit between the initial state and after N times of freeze-thaw cycles can be obtained as follows:

$$F_n = \frac{F_0}{1 + e_0 \{1 + [\beta u^T - (3(1 - 2\nu_i) / E_i) R_i]\}^n}. \quad (18)$$

It shows that the ultimate bearing capacity of rock is closely related to the times of freeze-thaw cycles. The rock damage resulting from frost heave is by nature a process of gradual accumulation and deterioration.

4. Sensitivity Analysis of Slope Stability Factors

4.1. Establishment of Slope Model. Rock is typical of highly heterogeneous material, and its internal microcrack extension is closely correlated with time and the surroundings. When it comes to the rock mass in open-pit slope, its deterioration is a nonlinear cumulative process that is closely related to time. Hence, the slope stability has typical time-dependent feature. During the entire service period ranging from the open-pit excavation, slope exposure, to inner dumping and reclamation, all the external conditions,

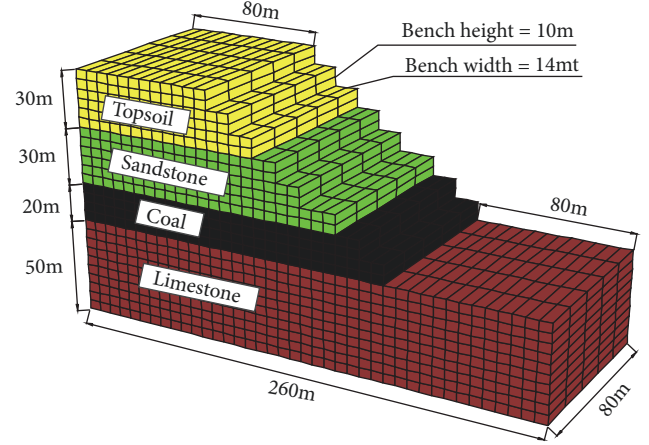


FIGURE 3: The details of modeled geometry.

such as blasting vibration, mining disturbance, weathering, groundwater infiltration, and frost heave, may cause the destruction and accumulation of fine structures inside the rock mass. As time goes on, the mechanical strength of rock mass attenuates and gradually converges to a lower limit value, which results in a gradual reduction of slope stability. Therefore, for the sake of safe construction and slope protection, it is of great significance to conduct sensitivity analysis on factors affecting slope stability.

The parameters of the computational model in this paper are 80m, 10m, and 14m for whole height, single bench height, and bench width, respectively. The artificial boundaries are set in places that are far enough from the area of the slope. Specifically, it is 80m at both slope top and bottom boundary and 50m at the lower boundary. The details of modeled geometry are shown in Figure 3. Boundary conditions of the model are set that nodes on the boundaries parallel to the X-Y plane, at $X=0$, $X=260m$, $Y=0$, and $Y=80m$, are constrained in the horizontal direction along X-axis and Y-axis. The nodes are free to move in the horizontal direction along Z-axis while bottom base nodes, at $Z=0$, are fully constrained. A Mohr-Coulomb material model is initially defined for all zones and the initial stress state is applied with gravity acting in the positive z direction. The model has 5200 zones and 6642 grid-points. The parameters and dimensionless values for each rock mass are shown in Table 2.

Parameters in Table 2 were substituted into the model for numerical simulation by using Flac3D. The convergence criterion for calculation is that the unbalanced force ratio falls below the default limiting value of 1×10^{-5} . In this model, the calculation stops at step 7704 and converges to an equilibrium state. The results of the simulation are shown in Figures 4–6.

As shown in Figure 4(a), there exist plastic zones within the slope that are parallel to the slope surface, while the basement and back of the slope are generally nondestructive and stable. Figure 4(b) demonstrates the shear-strain increment of the model. It is easy to observe that stress concentration appears at the bottom of the slope, which is close to the boundary of the plastic region. It is also in this area that the shear-strain increment reaches its maximum value. What is

TABLE 2: Properties of rock mass.

Stratum	Young's modulus (GPa)	Poisson's ratio	Mass density/ (kg/m ³)	Cohesion/ (MPa)	Angle of internal friction/(°)
Topsoil	1.80×10^{-2}	0.3	1786	13.4	21.3
Sandstone	8.30	0.214	2245	28.6	25.6
Coal	0.65	0.36	1540	13.0	20
Limestone	6.11	0.224	2130	27.0	22.5

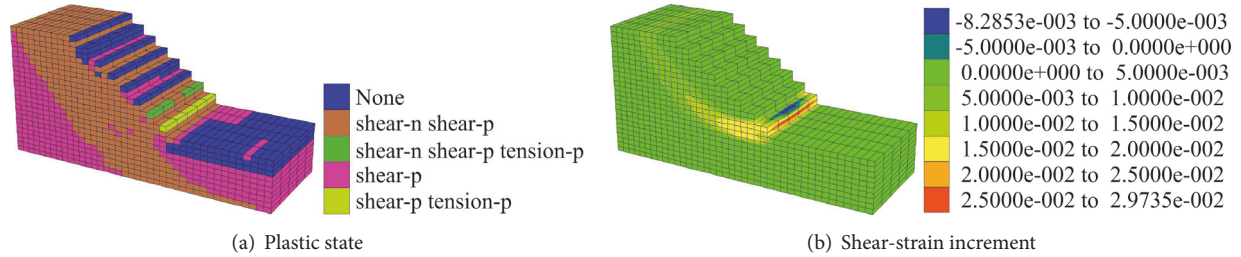


FIGURE 4: Plastic state and shear-strain increment of the slope model.

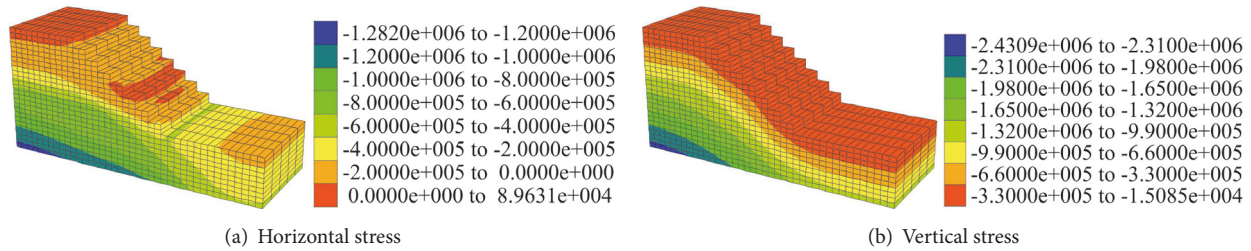


FIGURE 5: Horizontal and vertical stress of slope.

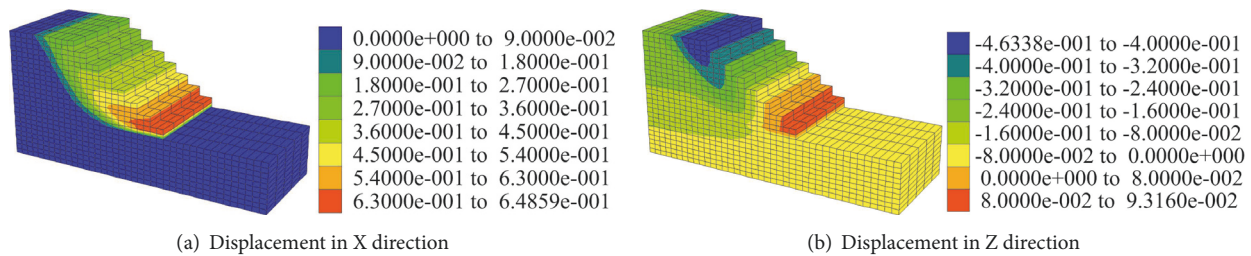


FIGURE 6: Displacement of slope in X and Z direction.

more, the stress concentration area and the interface of shear failure interpenetrate, thus bringing about potential landslide damage.

According to Figure 5(a), on one hand, the horizontal stress is small on the slope surface and gradually increases inward. It reaches its maximum at the bottom of the slope surface, which is not conducive to the slope stability. On the other hand, the vertical stress shown in Figure 5(b) enhances as the depth of the formation increases, indicating that, in vertical direction, the slope body is mainly affected by its own

gravity. And such results conform to the variation law of the ground stress.

Due to self-gravity, the slope body gradually settles from top to bottom and shear failure occurs at low-intensity places, causing the entire slope to gradually deform toward the empty surface of the slope. From Figures 6(a) and 6(b), the maximum displacement in the vertical direction at the top of the slope is 46.3cm, while the maximum horizontal displacement at the bottom of the slope is 64.9cm and the vertical displacement is 9.3cm, indicating that the ground heave

TABLE 3: The value of Fos for slope stability based on CVM.

Level	Fos Variables			
	Mass density	Young's modulus	Cohesion	Angle of internal friction
Level 1	1.6602	1.7070	1.6602	1.6602
Level 2	1.6289	1.6992	1.6211	1.6133
Level 3	1.5836	1.6992	1.5742	1.5664
Level 4	1.5566	1.7070	1.5332	1.5215

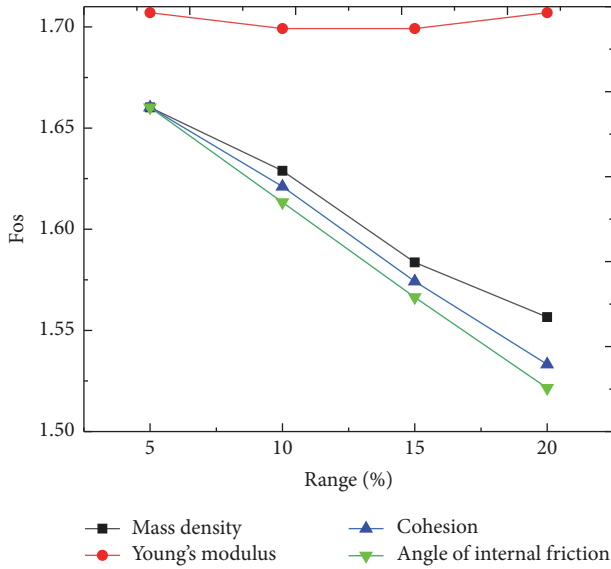


FIGURE 7: Trend of Fos.

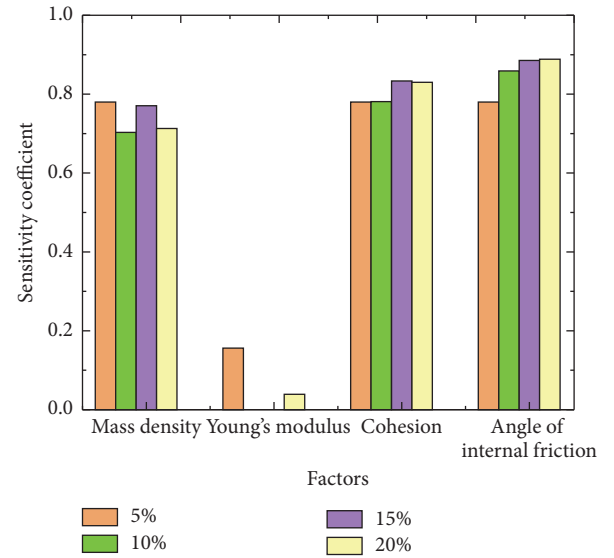


FIGURE 8: Histogram of sensitivity coefficient.

occurs at the bottom of the slope due to the compression of the upper landslide body, but no large-scale landslide has yet occurred. The corresponding Fos is 1.6992.

4.2. Control Variable Method (CVM) for Single Factor Sensitivity Analysis of Slope Stability. Controlling for a variable is often used in statistics for sensitivity analysis of multivariate problems; that is, when experiments are conducted in an attempt to evaluate the effect of one independent variable on another or more dependent variables, to ensure the measured effect is not influenced by external factors, other variables must be held constant. In this paper, the CVM is used to analyze the effect of a single factor on the slope stability. According to the degree of rock degradation measured in the tests, mass density, Young's modulus, cohesion, and internal friction angle are taken as independent variables and four levels are designed, respectively, in which the mass density increased by 5% per level, while Young's modulus, cohesion, and angle of internal friction reduced by 5% per level. The corresponding values of Fos are shown in Table 3.

In order to accurately and quantitatively analyze the sensitivity of each factor, the sensitivity coefficients are obtained as shown in Figure 7.

$$S_i = \frac{|\Delta f_i / f_i|}{|\Delta x_i / x_i|} \quad (19)$$

where S_i is the sensitivity coefficient of each factor, $|\Delta f_i / f_i|$ is the change rate of safety factor, and $|\Delta x_i / x_i|$ is the change rate of the evaluation index.

It can be seen that the sensitivity of all factors is relatively stable within the variation range of the factors studied. When it comes to the effect on slope stability, they rank like this: internal friction angle > cohesion > mass density > Young's modulus, as shown in Figure 8.

4.3. Orthogonal Design Method (ODM) for Multifactor Sensitivity Analysis of Slope Stability. In reality, there exist interactions amongst the factors affecting slope stability. Hence, when external conditions change, these influencing factors would change altogether. As can be seen from Figure 8, both the internal friction angle and cohesion have a great influence on the slope stability. Instead of probing into these two factors and their possible interactions, the CVM only takes the change of one single factor into consideration. If full factorials are carried out, $4^4=256$ tests should be required according to the factors and levels of this experiment, whose scale will be too large to implement. Therefore, the orthogonal design is used for test design.

To explore the influencing factors of slope stability, the orthogonal tests ($L16_4^5$) were designed and carried out. There were five factors in the tests, each at four levels, thus generating sixteen different combinations. Factors A-D,

TABLE 4: Orthogonal design and calculation results.

Test No.	A	B	C	D	C * D	Fos
1	1(105%)	1(95%)	1(95%)	1(95%)	1	1.5742
2	1	2(90%)	2(90%)	2(90%)	2	1.4863
3	1	3(85%)	3(85%)	3(85%)	3	1.4043
4	1	4(80%)	4(80%)	4(80%)	4	1.3223
5	2(110%)	1	2	3	4	1.4121
6	2	2	1	4	3	1.4082
7	2	3	4	1	2	1.4199
8	2	4	3	2	1	1.416
9	3(115%)	1	3	4	2	1.2988
10	3	2	4	3	1	1.3066
11	3	3	1	2	4	1.4629
12	3	4	2	1	3	1.4707
13	4(120%)	1	4	2	3	1.3223
14	4	2	3	1	4	1.4043
15	4	3	2	4	1	1.3105
16	4	4	1	3	2	1.3887

TABLE 5: Value of range analysis.

Parameter	A	B	C	D	C * D
K_{1j}	1.447	1.402	1.474	1.467	1.402
K_{2j}	1.414	1.401	1.42	1.422	1.414
K_{3j}	1.385	1.399	1.381	1.394	1.401
K_{4j}	1.372	1.415	1.343	1.335	1.4
R_j	0.075	0.016	0.131	0.132	0.014
Affect order	D>C>A>B>C * D				

respectively, represented the mass density, Poisson's modulus, cohesion, and internal friction angle while factor C * D stood for the interaction of factors C and D. The details and calculation results are shown in Table 4.

Set K_{ij} as the standard deviation for factor j at level i , and the difference between the maximum and minimum average standard deviations for the factor can be expressed as

$$R_j = \max \{K_{1j}, K_{2j}, \dots, K_{ij}\} - \min \{K_{1j}, K_{2j}, \dots, K_{ij}\}. \quad (20)$$

The calculation results are shown in Table 5. Larger range value indicates that the influencing factor makes a greater impact on system performance [18]. The influence of the five factors on slope stability was ranked as follows: D>C>A>B>C*D. In other words, the internal friction angle and cohesion exerted the greatest influence, followed by mass density. And it could also be seen that the delta R_j of B was obviously smaller than that of D and C, which indicated that the change of Poisson's modulus had little effect on the slope stability. Besides, the impact resulting from the interaction between internal friction angle and cohesion was much less than that of a single factor.

Furthermore, the F distribution of the factors was tested by variance analysis. The sums of squared deviations of each factor can be expressed as

$$S_j = \frac{r}{n} \sum_{i=1}^r K_{ij}^2 - \frac{1}{n} \left(\sum_{i=1}^n y_i \right)^2 \quad (21)$$

where S_j is the sum of squared deviations of each factor, y_i is the calculation result of test i , and $i=1,2,3,\dots,n$, r is the level number of one single factor. The degree of freedom of S_j can be calculated as $f_i=r-1$. Since all the columns in the orthogonal table have factors, that is, there are no empty columns herein, the minimum value in the S_j is taken as the sum of squared deviations for errors. The statistics can then be defined as follows:

$$F = \frac{S_j/f_j}{S_e/f_e} = \overline{S_j}/\overline{S_e}. \quad (22)$$

By the F distribution table, $F_{1-0.01}(3,3) = 29.5$, $F_{1-0.05}(3,3) = 9.28$, and $F_{1-0.1}(3,3) = 5.39$ can be obtained. Given that $F_C = \overline{S_C}/\overline{S_e} = (0.038/3)/(0.001/3) = 38 > 29.5 = F_{1-0.01}(3,3)$, $F_D = \overline{S_D}/\overline{S_e} = (0.037/3)/(0.001/3) = 37 > 29.5 = F_{1-0.01}(3,3)$, it is considered that the internal friction angle and cohesion have an extremely significant influence on the test results. $F_A = \overline{S_A}/\overline{S_e} = (0.013/3)/(0.001/3) = 13 > 9.28 = F_{1-0.05}(3,3)$; it shows that mass density has a significant influence. Finally, $F_{C*D} = \overline{S_{C*D}}/\overline{S_e} = (0.001/3)/(0.001/3) = 1 < 5.39 = F_{1-0.1}(3,3)$; it indicates that the effect that the interaction between internal friction angle and cohesion exerts on the test results is not significant. This is in accordance with the previous conclusions obtained by value analysis. The results of the tests are observed in Table 6.

TABLE 6: Variance analysis of various factors.

Source of variance	S_j	f_j	$F = \frac{S_j/f_j}{S_e/f_e}$	Significance
A	0.013	3	13.000	**
B ^A	0.001	3	1.000	
C	0.038	3	38.000	* * *
D	0.037	3	37.000	* * *
C * D	0.001	3	1.000	
Error	0.001	3		

* * * means extremely significant influence ($\alpha < 0.01$); ** means significant influence ($0.01 < \alpha < 0.05$); * means general influence ($0.05 < \alpha < 0.1$); other values are not significant ($0.1 < \alpha$)

5. Conclusions

(1) The freeze-thaw cycle tests of saturated sandstone showed that the properties of rock mass were highly sensitive to increasing freeze-thaw cycles. Of all the indexes, physical indexes such as mass density, microcrack rate, and longitudinal wave velocity increased, while mechanical indexes like internal friction angle, cohesion, Poisson's modulus, and compressive strength showed a decreasing trend. It is reasonable to conclude that freeze-thaw cycles caused the expansion and coalescence of microcracks within rock samples and the destruction of the internal rock structure and finally reduced the ultimate bearing capacity of rock mass.

(2) Considering the microcrack propagation factor, the constitutive model of saturated rock deterioration was presented by taking the volume of phase transition of water in microcracks as a variable. It was verified that the ultimate bearing capacity of rock mass gradually decreased under the freeze-thaw cycles, and the damage of frost heave to rock was a process of gradual accumulation and deterioration.

(3) Slope stability deteriorates due to freeze-thaw cycles. To study the sensitivity to such deterioration of influencing factors of slope stability, a slope model was established by Flac3D. Through the analysis of CVM and ODM, it was obtained that the sensitivity of all factors was relatively stable within the variation range of factors studied. Through variance analysis, $F_C = 38 > 29.5$ and $F_D = 37 > 29.5$ when the test level $\alpha = 0.01$ is given, which shows that the internal friction angle and cohesion are the most important determinant affecting the slope stability. Meanwhile, $F_B = F_{C*D} = 1 < 5.39$ when $\alpha = 0.1$; it indicates that Poisson's modulus as well as the interaction between internal friction angle and cohesion are not significant. When it comes to the effect on slope stability, internal friction angle ranks the first, followed by cohesion, mass density, and Poisson's modulus. Therefore, it is advisable to pay more attention to the change of significant factors in actual slope construction and protection.

Data Availability

The numerical simulation data files used to support the findings of this study are available from the corresponding author upon request.

Conflicts of Interest

The authors declare that there are no conflicts of interest regarding the publication of this paper.

Acknowledgments

This work was supported by the Open Projects of Research Center of Coal Resources Safe Mining and Clean Utilization (Liaoning) (LNTUI7KF08), National Key Research and Development Program of China (2016YFC0501100), the National Natural Science Foundation of China (51774271, 51604264, and 51464043), and the Foundation of Xinjiang Educational Committee (XJEDU2016S088). The authors are grateful to all the generous support.

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