

## Research Article

# EPR, UV-Visible, and Near-Infrared Spectroscopic Characterization of Dolomite

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Dolomite mineral samples having white and light green colors of Indian origin have been characterized by EPR, optical, and NIR spectroscopy. The optical spectrum exhibits a number of electronic bands due to presence of Fe(III) ions in the mineral. From EPR studies, the parameters of  $g$  for Fe(III) and  $g$ ,  $A$ , and  $D$  for Mn(II) are evaluated and the data confirm that the ions are in distorted octahedron. Optical absorption studies reveal that Fe(III) is in distorted octahedron. The bands in NIR spectra are due to the overtones and combinations of water molecules. Thus EPR and optical absorption spectral studies have proven useful for the study of the solid state chemistry of dolomite.

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## 1. Introduction

Carbonate minerals in various forms such as limestone, dolomite, and calcite constitute the earth's largest CO<sub>2</sub> resource. The dolomite group of minerals have the general formula A<sup>++</sup>B<sup>++</sup>(CO<sub>3</sub>)<sub>2</sub>. When large amount of iron is present, the mineral ankerite forms and when excess manganese is present the mineral kuntnohorite forms instead of dolomite. All these minerals have the same internal structure but differ chemically from each other [1]. This group of minerals is classified [2] into several types as given below:

Ankerite: Ca(Fe,Mg,Mn)(CO<sub>3</sub>)<sub>2</sub>

Benstonite: (Ba, Sr)<sub>6</sub>(Ca, Mn, Mg)<sub>6</sub>(CO<sub>3</sub>)<sub>12</sub>

Dolomite: CaMg(CO<sub>3</sub>)<sub>2</sub>

Huntite: CaMg<sub>3</sub>(CO<sub>3</sub>)<sub>4</sub>

Kuntnohorite: Ca(Mn,Mg,Fe)(CO<sub>3</sub>)<sub>2</sub>

Minrecordite: (Ca,Zn)(CO<sub>3</sub>)<sub>2</sub>

Norsethite: BaMg(CO<sub>3</sub>)<sub>2</sub>.

In the above group of minerals, Mg(II) can be replaced by ferrous iron and manganese. In dolomites the Mn(II) ion is

a natural substitutional impurity at the sites of both Ca(II) and Mg(II) ions. Composition of low-Mn ankerite generally approached the formula CaMg(CO<sub>3</sub>)<sub>2</sub>-CaFe(CO<sub>3</sub>)<sub>2</sub> [3]. This group of minerals has the wide useful applications especially as fillers some of which are given below:

- (i) in cement paints, exterior paints, primers, putties, powder coatings, and industrial finishes;
- (ii) in PVC footwear, PVC pipes, cables, and others;
- (iii) in paper to give smoothness and gloss;
- (iv) in adhesives and carpet backing;
- (v) in toothpaste, cosmetic, and soap industry.

Dolomites are isostructural having space group R<sub>3</sub>. Dolomite structure reveals that there are two distinct cation sites denoted by A and B [4]. These both sites form nearly regular octahedra in which each corner of an octahedron is oxygen from a different CO<sub>3</sub> group. The site A is occupied by Ca and the B site is occupied by Mg in the ideally ordered case with layers of Ca octahedron alternating with layers of (Mg,Fe,Mn) octahedron. The octahedra are linked

TABLE 1: Chemical composition of dolomite.

Component	Percentage		
	1 (green)	2 (white)	3 (pale yellow)
CaCO <sub>3</sub>	53.6	56.6	
MgCO <sub>3</sub>	42.0	43.0	
CaO	31.7	31.7	31.27
MgO	20.5	20.5	21.12
Fe <sub>2</sub> O <sub>3</sub>	0.30	0.30	
SiO <sub>2</sub>	5.50	1.50	0.12
Al <sub>2</sub> O <sub>3</sub>	0.50	0.50	
FeO			0.22

1: The Good Earth Udaipur, Rajasthan, India; 2: Famous Minerals & Chemicals Private Limited; 3: Haley, Ontario, Canada, Mineral data publishing version I, 2001–2005.

by sharing corners simultaneously with octahedra of the opposite kind with CO<sub>3</sub> groups [4]. The two octahedra in dolomite are trigonally distorted by elongation along the three-fold axis. The larger A octahedron is always slightly more distorted than that in B [4, 5]. Detailed spectral studies have been reported on dolomite-ankerite series but not on dolomite of Indian origin [3, 6–8]. Several studies have been made on natural dolomites from different origins using EPR studies [9–11]. However, few EPR studies have been done which could identify paramagnetic radicals in dolomites around  $g = 2.0023$  [11]. Nevertheless, complete interpretation of the EPR spectra of dolomites is not available. Hence, optical absorption and EPR studies on some natural carbonate minerals have been made [12].

Transition metal ions such as Fe(III), Mn(II), and V(IV) display a rich coordination in minerals. For the identification and characterization of these metal ions in natural systems, electron paramagnetic resonance has been found to be the most powerful analytical technique. In India, Kadapa Basin is endowed with rich mineral resources. In this paper, we report the EPR and optical absorption spectral studies of dolomites obtained from Vempalli Mandal of Kadapa district, Andhra Pradesh, India.

## 2. Theory

Several EPR parameters (i.e.,  $g$ ,  $A$ ,  $D$ , and  $E$ ) are used while interpreting EPR data. The  $g$  parameter is a measure of the coupling between the unpaired electron's spin angular momentum ( $S$ ) and its orbital angular momentum ( $L$ ) [13].

The unpaired electron interacts (couples) with the nuclear spin ( $I$ ) to form a  $(2I + 1)$  line hyperfine structure centered on  $g$  and spaced with the distance quantified by the hyperfine coupling parameter  $A$ . The coupling between the nuclear and electron spins becomes stronger as the  $A$  parameter becomes larger. The combination of  $g$  and  $A$  parameters can be utilized to differentiate between electron environments of Fe<sup>3+</sup> and Mn<sup>2+</sup> ions. The EPR zero field splitting (ZFS) parameters  $D$  and  $E$  measure the deviation of the ion crystal field from ideal tetrahedral or

octahedral symmetries and they apply to ions with more than one unpaired electron (e.g., low field Fe<sup>3+</sup> and Mn<sup>2+</sup>). However, the broad nature of EPR spectra of Fe<sup>3+</sup> makes the determination of  $D$  and  $E$  difficult [14].

Mn(II), being a d<sup>5</sup> ion, has total spin  $S = 5/2$ . This state splits into three Kramers' doublets,  $\pm 5/2>$ ,  $\pm 3/2>$ , and  $\pm 1/2>$  separated by  $4D$  and  $2D$ , respectively, where  $D$  is the zero-field splitting parameter. The deviation from axial symmetry leads to a term known as  $E$  in the spin Hamiltonian. The value of  $E$  can be easily calculated from single crystal measurements. A nonzero value of  $E$  results in making the spectrum unsymmetrical about the central sextet.

## 3. Results and Discussion

### 3.1. Total Chemical Analyses

Dolomite minerals examined in this study are ankerite and dolomite. Elemental compositions of metal are listed in Table 1 [15–17].

### 3.2. EPR Spectral Analysis

EPR spectrum of the dolomite mineral samples recorded at room temperature is given in Figures 1(a) and 1(b). The overall EPR spectral features are similar in both dolomites. It is well known that the EPR studies of transition metal ion Mn(II) is easily observed at room temperature, even if present in minute levels, compared to other ions. Both spectra (Figures 1(a) and 1(b)) clearly indicate the presence of Mn(II) in the sample. An expanded version of central sextet is shown in Figures 2(a) and 2(b).

Figures 1(a) and 1(b) contain a strong sextet at the centre of the spectrum corresponding to the electron spin transition  $+1/2>$  to  $-1/2>$ . Generally, in most of the cases, the powder spectrum is characterized by a sextet, corresponding to this transition. The other four transitions, corresponding to  $\pm 5/2> \leftrightarrow \pm 3/2>$  and  $\pm 3/2> \leftrightarrow \pm 1/2>$  are not seen due to their high anisotropy in  $D$ . If  $E \neq 0$ , the EPR spectrum will not be symmetrical about the central sextet. The expanded version of the powder spectrum (Figures 2(a) and 2(b)) of both dolomite mineral samples indicate the presence of at least four types of Mn(II) impurities in the mineral. The sixth manganese hyperfine resonance clearly contains two lines (marked by a and b). This is also followed by two more sets of weak lines (marked by \* and \*\*) on either side of the intense Mn(II) signals as shown in Figures 2(a) and 2(b). This is the same in both samples. This is due to the presence of Mn(II) ions in a low symmetry environment [18]. The EPR line intensity of Mn(II) hyperfine lines observed in green dolomite is more intense, when compared to that in the white dolomite mineral. This indicates clearly that the concentration of Mn(II) ions in green dolomite sample is more than that in the white dolomite. Even when viewed with a naked eye, the sample is pale greenish pink in color indicating the substitution of manganese ion in larger concentration in this mineral.

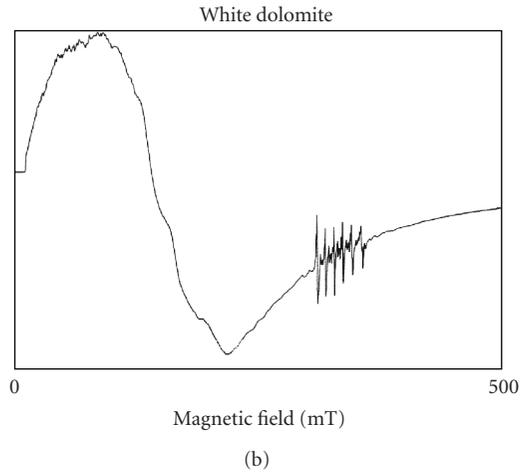
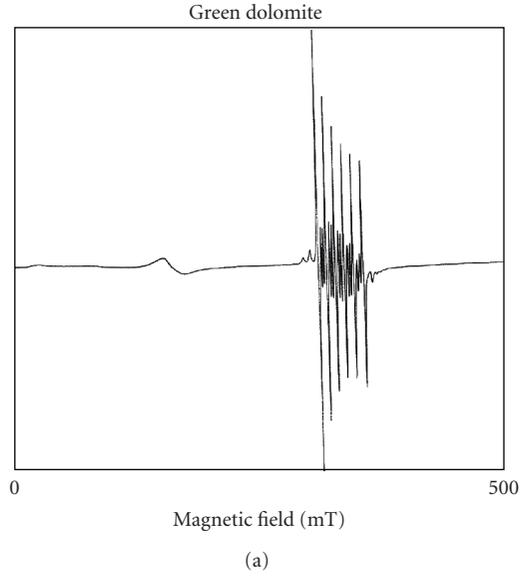


FIGURE 1: (a) EPR spectrum of green dolomite at RT ( $\nu = 9.39219$  GHz). (b) EPR spectrum of white dolomite at RT ( $\nu = 9.39208$  GHz).

The observed EPR spectrum can be explained by the spin-Hamiltonian of the form as shown in the following equation (see [19]):

$$H = \beta B g S + D \left\{ S_Z^2 - \frac{1}{3} S(S+I) \right\} + S A I. \quad (1)$$

Here the first term represents the electron-Zeeman interaction. The second term represents the zero field contribution, and the third term represents the nuclear-Zeeman interaction. The extra set of resonances in between the main sextet is due to the forbidden transitions. From the forbidden doublet lines, the Zero field splitting parameter,  $D$ , has been calculated using the formula (see [20, 21])

$$\Delta H = \left( \frac{2D^2}{H_m} \right) \left[ \frac{1 + 16(H_m - 8Am)^2}{9H_i H_m - 64Am} \right]. \quad (2)$$

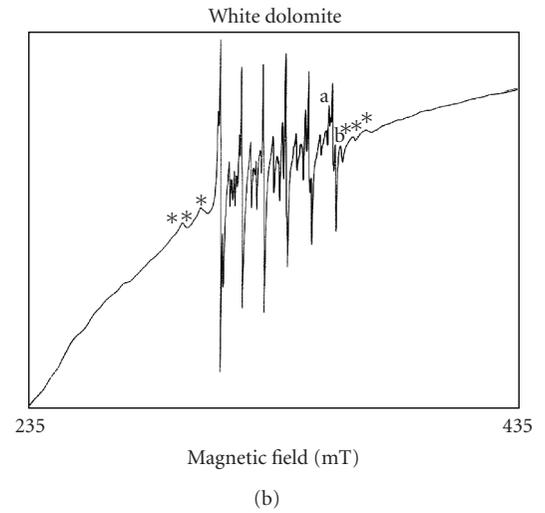
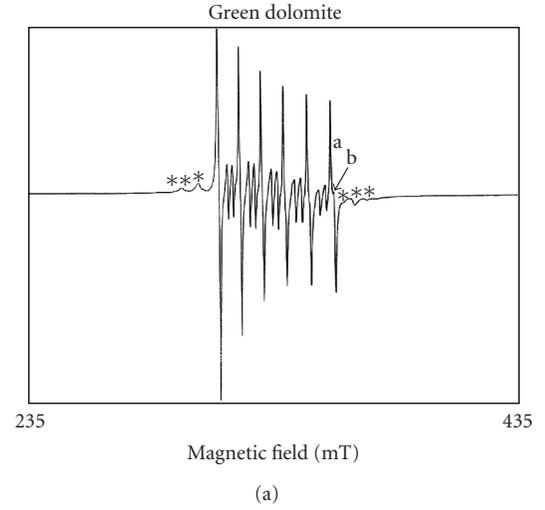


FIGURE 2: (a) Expanded form of EPR spectrum of green dolomite at RT ( $\nu = 9.39128$  GHz). (b) Expanded form of EPR spectrum of white dolomite at RT ( $\nu = 9.39183$  GHz).

Here

$$H_m = H_o - Am - \frac{[I(I+1) - m^2]A^2}{2H_o}. \quad (3)$$

From the powder spectrum of the mineral, the following parameters have been calculated:

White dolomite

$$g = 1.992 \pm 0.015, \quad A = 9.15 \pm 0.43 \text{ mT}, \quad (4)$$

$$D = 34.45 \pm 0.36 \text{ mT}.$$

Green dolomite

$$g = 1.993 \pm 0.017, \quad A = 9.35 \pm 0.43 \text{ mT}, \quad (5)$$

$$D = 34.25 \pm 0.35 \text{ mT}$$

The observed  $g$ ,  $A$ , and  $D$  values are comparable with other similar systems and agree well within the experimental

error [8, 22–25]. The large value of  $D$  indicates a considerable amount of distortion around the central metal ion. As EPR is highly sensitive to Mn(II) impurity, four such sites have been noticed in the EPR spectrum. A close look at the EPR spectrum indicates a nonzero value for  $E$ , which is very difficult to estimate from powder spectrum. A close look at the centre of the EPR spectrum indicates a broad line underneath the Mn(II) sextet. This can be attributed to the iron impurity in the sample. Generally, Fe(III) and Fe(II) impurities will give rise to a broad line around with a  $g$  value of 2. No extra information can be obtained from the powder spectrum recorded at room temperature and low temperature except that iron is present as an impurity. Even though the percentage of iron is more than that of manganese, its intensity is relatively smaller compared to the signals obtained from the manganese impurity. The EPR spectrum of white dolomite exhibits a broad resonance centered when  $g$  is 4.20; similarly a weak resonance is observed when  $g$  is 4.34 in the case of green dolomite. Those are attributed to the presence of isolated Fe(III) ions incorporated as an impurity in the mineral.

The hyperfine constant “ $A$ ” value provides a qualitative measure of the ionic nature of bonding with Mn(II) ion. The percentage of covalency of Mn-ligand bond in green dolomite has been calculated using “ $A$ ” (9.35 mT) value obtained from the EPR spectrum and Matumura’s plot [26]. It corresponds to an ionicity of 94%, where, as in the case of white dolomite, “ $A$ ” with a value of 9.15 mT gives an ionicity of 91%. Also the approximate value of hyperfine constant ( $A$ ) is calculated for both dolomites by using covalency ( $C$ ) using the following equation (see [27, 28]):

$$A_{\text{iso}} = (2.04C - 104.5)10^{-4} \text{ cm}^{-1}. \quad (6)$$

The value obtained is  $92 \times 10^{-4} \text{ cm}^{-1}$  for green dolomite, whereas for the white dolomite it is  $91 \times 10^{-4} \text{ cm}^{-1}$ . This calculated value agrees well with the observed hyperfine constant for both samples indicating ionic character for Mn–O bond in the mineral under study. The number of ligands around Mn(II) ion is estimated using the covalency [27] equation for  $C$

$$C = \frac{1}{n} [1 - 0.16(X_p - X_q) - 0.035(X_p - X_q)^2]. \quad (7)$$

Here  $X_p$  and  $X_q$  represent electronegativities of metal and ligand. Assuming  $X_p = X_{\text{Mn}} = 1.4$  and  $X_q = X_{\text{O}} = 3.5$ , the number of ligands ( $n$ ) obtained are 20. This suggests that Mn(II) may be surrounded by four  $\text{CO}_3^{2-}$  and two oxygens. Further the  $g$  value for the hyperfine splitting is indicative of the nature of bonding. If the  $g$  value shows negative shift with respect to free electron  $g$  value of 2.0023, the bonding is ionic and conversely, if the shift is positive, the bonding is more covalent in nature [29]. In the present work, from the observed negative values of 0.0103 for white dolomite and that of  $-0.0093$  for green dolomite, it is apparent that the Mn(II) is in an ionic environment. EPR spectrum of white dolomite (Figure 2(b)) reveals that in between first and second manganese hyperfine resonance lines a very weak sharp peak at (319.2 mT) having a  $g$  value 2.1019 is observed.

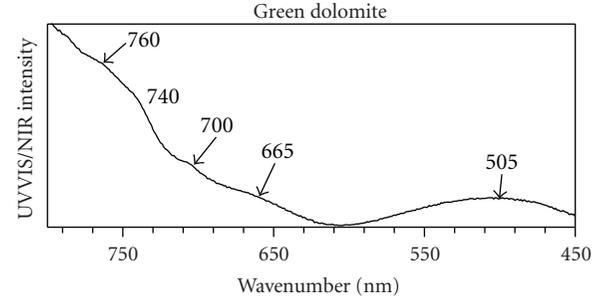


FIGURE 3: Optical absorption spectrum of green dolomite at room temperature.

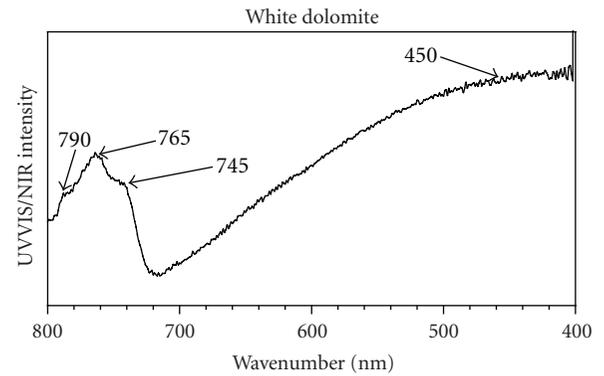


FIGURE 4: Optical absorption spectrum of white dolomite at room temperature.

This can be ascribed to a radical such as  $\text{CO}_3^{2-}$  or  $\text{NO}_3^{2-}$ . The absence of any super hyperfine structure in the radical suggests that the radical may be  $\text{CO}_3^{2-}$  [11].

### 3.3. Optical Absorption Spectral Analysis

Ferric iron, Fe(III), has the electronic configuration  $\text{Ar } (3d^5)$  with a half filled d-shell having one unpaired electron in each of the orbitals. Hence ground state configuration is  $t_{2g}^3 e_g^2$ . The free ion levels of Fe(III), in the order of increase in energy, are  ${}^6S, {}^4G, {}^4P, {}^4D, {}^4F$ . The energy levels for Fe(III) in an octahedral environment are  ${}^6A_{1g}({}^6S), {}^4T_{1g}({}^4G), {}^4T_{2g}({}^4G), {}^4A_{1g}({}^4G), {}^4E_g({}^4G), {}^4T_{2g}({}^4D),$  and  ${}^4E_g({}^4D)$ . The  ${}^4A_{1g}({}^4G)-{}^4E_g({}^4G)$  and  ${}^4E_g({}^4D)$  levels have relatively less influence compared to other levels by crystal field. It means that the relative sharp levels can be expected in the absorption spectrum which is the criterion for assignment of levels for Fe(III) ion. Since all the excited states of Fe(III) ion will be either quartets or doublets, the optical absorption spectra of Fe(III) ions will have only spin forbidden d-d transitions. These occur from the ground state  ${}^6A_{1g}(S)$  to the excited states  ${}^4T_{1g}(G), {}^4T_{2g}(G), {}^4A_{1g}(G), {}^4E_g(G), {}^2T_{1g}(D), {}^4E_g(D),$  and  ${}^4T_{1g}(P)$  of regular octahedron [30].

Optical absorption spectrum of green dolomite recorded in the mull form at RT in the range of 200–1200 nm is shown in Figure 3 and that of white dolomite is shown in Figure 4. The green dolomite spectrum shows energies at  $13160 \text{ cm}^{-1}$

TABLE 2: Band headed data with assignments for Fe(III) in natural dolomite mineral.

Dolomite green			Dolomite white			Transition from ${}^6A_{1g}$
$Dq = 900, B = 605, C = 2600 \text{ cm}^{-1}$			$Dq = 900, B = 600, C = 2600 \text{ cm}^{-1}$			
$\alpha = 90 \text{ cm}^{-1}$			$\alpha = 90 \text{ cm}^{-1}$			
Wave length (nm)	Wavenumber ( $\text{cm}^{-1}$ )		Wave length (nm)	Wavenumber ( $\text{cm}^{-1}$ )		
	Observed	Calculated		Observed	Calculated	
			790	12660		
760	13160	13161	765	13070	13087	
740	13515		745	13425		${}^4T_{1g}(G)$
700	14285		—	—		
665	14995	16951	—	—	16852	${}^4T_{2g}(G)$
505	19800	20813	450	22220	20739	${}^4E_g(G)$
						${}^4A_{1g}(G)$

(760 nm),  $13515 \text{ cm}^{-1}$  (740 nm),  $14285 \text{ cm}^{-1}$  (700 nm),  $14995 \text{ cm}^{-1}$  (665 nm), and  $19800 \text{ cm}^{-1}$  (505 nm)  $\text{cm}^{-1}$  in the UV-Vis region whereas white dolomite spectrum shows energies at  $12660 \text{ cm}^{-1}$  (790 nm),  $13070 \text{ cm}^{-1}$  (765 nm),  $13425 \text{ cm}^{-1}$  (745 nm),  $22220 \text{ cm}^{-1}$  (450 nm)  $\text{cm}^{-1}$  in the UV-Vis region. In green dolomite, the band observed at  $13515 \text{ cm}^{-1}$  with split component on left side at  $13160 \text{ cm}^{-1}$  is assigned to the transition  ${}^6A_{1g}(S) \rightarrow {}^4T_{1g}(G)$  whereas the band at  $14995 \text{ cm}^{-1}$  is assigned to  ${}^4T_{2g}(G)$  transition. The third band visible around  $19800 \text{ cm}^{-1}$  is assigned to  ${}^4A_{1g}(G)$ ,  ${}^4E_g(G)$  transition. These bands are characteristic of Fe(III) ion in octahedral symmetry in the mineral. This is also further supported by EPR studies. Using Tree's polarization term  $\alpha = 90 \text{ cm}^{-1}$  [26], the energy matrices of the  $d^5$  configuration are solved for various  $B$ ,  $C$ , and  $Dq$  values. The evaluated parameters are presented in Table 2. A comparison is also made between the calculated and observed energies of the bands and these are presented in Table 2. The sharp band observed in green dolomite at  $14285 \text{ cm}^{-1}$  may be due to Mn(II). The very sharp band observed at  $13070 \text{ cm}^{-1}$  in white dolomite with split component on either side (at  $12660$  and  $13425 \text{ cm}^{-1}$  with average of  $13043 \text{ cm}^{-1}$ ) is assigned to the transition  ${}^6A_{1g}(S) \rightarrow {}^4T_{1g}(G)$ . Similar assignment is also made in green dolomite. The intensity of the band in white dolomite is large when compared to green dolomite. Further in white dolomite the band visible around  $16000 \text{ cm}^{-1}$  due to the transition of  ${}^6A_{1g}(S) \rightarrow {}^4T_{2g}(G)$  is not seen. This may be due to low concentration of iron in the mineral. The third band at  $22220 \text{ cm}^{-1}$  is assigned to  ${}^4A_{1g}(G)$ ,  ${}^4E_g(G)$  transition. These bands are characteristic of Fe(III) ion in octahedral symmetry in the mineral. This is also further supported by EPR studies. Using Tree's polarization term  $\alpha = 90 \text{ cm}^{-1}$  [31], the energy matrices of the  $d^5$  configuration are solved for various  $B$ ,  $C$ , and  $Dq$  values. A comparison is also made between the calculated and observed energies of the bands and these are presented in Table 2. These results agree well with the reported values of different dolomite minerals of Indian origin [12].

### 3.4. Near Infrared Spectroscopy

Water has three fundamental modes. They are symmetric OH stretch ( $\nu_1$ ), H–O–H bending mode ( $\nu_2$ ), and asymmetric OH stretch ( $\nu_3$ ). In solid phase, these modes appear at  $3220$  ( $\nu_3$ ),  $1620$  ( $\nu_2$ ), and  $3400$  ( $\nu_1$ )  $\text{cm}^{-1}$  [32]. The shifting of  $\nu_1$  and  $\nu_3$  towards lower frequency and  $\nu_2$  towards higher frequency is due to hydrogen bonding [33].

#### 3.4.1. The 11500 to 5500 $\text{cm}^{-1}$ Spectral Region

Figure 5 shows near infrared spectrum of dolomite mineral from  $11500$ – $5500 \text{ cm}^{-1}$  region. The near-IR spectral regions may be conveniently divided as follows: the high wave number region  $>7500 \text{ cm}^{-1}$ : in this region, electronic bands due to the characteristic of  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$  ions are observed. The two bands observed in green dolomite at  $9286$  and  $8581 \text{ cm}^{-1}$  are broad and are of very low intensity which are assigned to the two components of the transition  ${}^2T_{2g} \rightarrow {}^5E_g$ . This may be due to trace of Fe(II) present in the sample. The average of these bands  $8934 \text{ cm}^{-1}$  is taken as  $10Dq$ . But in white dolomite no Fe(II) bands are noticed.

The high wave number region between  $7200 \text{ cm}^{-1}$  and  $6300 \text{ cm}^{-1}$  is attributed to the first overtone of the fundamental hydroxyl stretching mode. The very sharp group of bands at  $7127$  and split components on either side at  $7237$  and  $7036 \text{ cm}^{-1}$  with high intensity are due to  $2\nu_3$ , the asymmetric OH stretch in the green dolomite. Exactly the similar very sharp group of bands observed at  $7154 \text{ cm}^{-1}$  with components at  $7232$ ,  $7173$ , and  $6920 \text{ cm}^{-1}$  with maximum intensity are assigned to the same  $2\nu_3$  asymmetric OH stretch in the white dolomite.

#### 3.4.2. The 5500 to 4000 $\text{cm}^{-1}$ Spectral Region

The spectral region from  $5500 \text{ cm}^{-1}$  to  $4000 \text{ cm}^{-1}$  is made up of two parts: the  $4000$ – $4500 \text{ cm}^{-1}$  region in which OH combination bands are found and  $4500$ – $5500 \text{ cm}^{-1}$  region

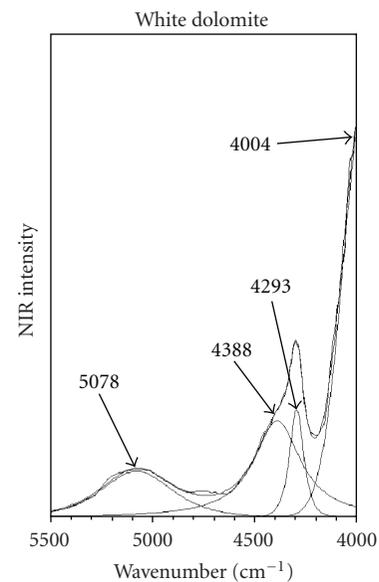
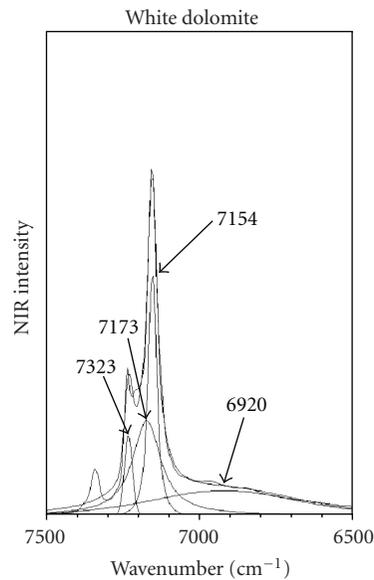
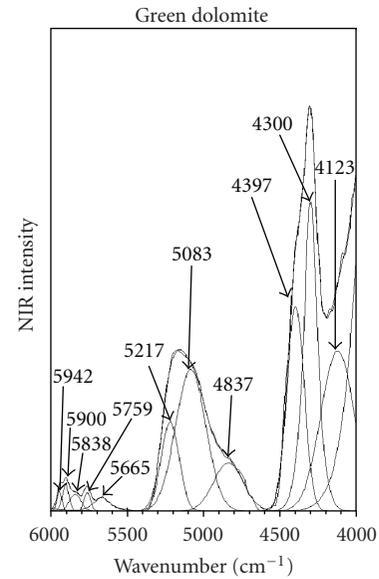
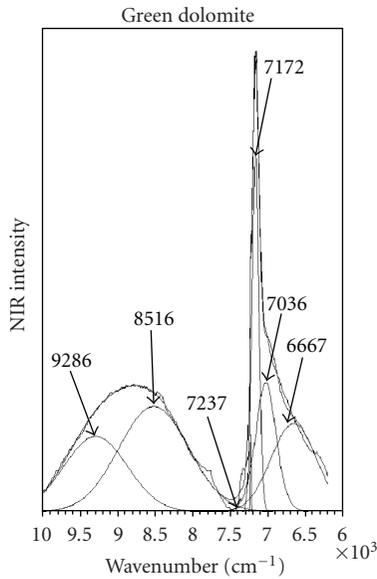


FIGURE 5: NIR spectrum of dolomite at room temperature from 11000–5500  $\text{cm}^{-1}$ .

FIGURE 6: NIR spectrum of dolomite at room temperature from 5500–4000  $\text{cm}^{-1}$ .

in which water combination bands are observed. Figure 6 shows a profile of bands at  $4300 \text{ cm}^{-1}$  with high intensity and with components on either side at  $4397$  and  $4123 \text{ cm}^{-1}$ . These bands in green dolomite are assigned to the water OH combination. In white dolomite, similar bands are observed with less intensity at  $4293 \text{ cm}^{-1}$  with a shoulder at  $4388 \text{ cm}^{-1}$  which may be attributed to the water OH combination. The intense band observed at  $5083 \text{ cm}^{-1}$  in green dolomite is attributed to the water OH overtone. The same band with very low intensity observed at  $5078 \text{ cm}^{-1}$  in white dolomite is assigned to the water OH overtone. In general, the NIR bands observed for white dolomite are less intense when compared to green dolomite which indicates that water is more in green dolomite than that present in white dolomite.

#### 4. Conclusion

EPR results indicate that Fe(III) and Mn(II) are present in the mineral in distorted octahedral environment. The  $g$  value of 4.34 in green dolomite and that of 4.20 in white dolomite are ascribed to Fe(III). This supports that the sample contains isolated Fe(III) ions in the lattice of the dolomite mineral. In both dolomites, the values of  $g$ ,  $A$ , and  $D$  in the spectra are  $g = 1.992 \pm 0.015$ ,  $A = 9.15 \pm 0.43 \text{ mT}$ , and  $D = 34.45 \pm 0.36 \text{ mT}$  (white dolomite) and  $g = 1.993 \pm 0.017$ ,  $A = 9.35 \pm 0.43 \text{ mT}$ , and  $D = 34.25 \pm 0.35 \text{ mT}$ , respectively, and they are due to Mn(II) indicating distorted octahedral environment in the mineral. These results suggest that Mn(II) is replaced either by Ca or Mg site. Also white dolomite shows a free

radical at  $g = 2.1019$  which is ascribed to  $\text{CO}_3^{2-}$  which may be a geological indicator. Thus EPR is a very sensitive technique for the magnetic analysis of paramagnetic components in diamagnetic (carbonate) minerals.

The optical absorption studies indicate that Fe(III) ion is present in major quantity in both dolomites whereas Fe(II) is in traces in the green dolomite mineral. These ions are also present in distorted octahedral environment. Green color of the mineral is due to Fe(II) which is evident from the optical absorption spectrum.

Near-infrared spectrum is due to overtones and combinations of water molecules. Also green dolomite sample contains more lattice water than the white dolomite mineral.

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