

Research Article

Sulfate Sources of Thermal Sulfate Reduction (TSR) in the Permian Changxing and Triassic Feixianguan Formations, Northeastern Sichuan Basin, China

Pingping Li⁽¹⁾,^{1,2} Huayao Zou,^{1,2} Fang Hao,³ and Xinya Yu⁴

¹State Key Laboratory of Petroleum Resources and Prospecting, China University of Petroleum, Beijing 102249, China
 ²College of Geosciences, China University of Petroleum, Beijing 102249, China
 ³China University of Petroleum, East China, Qingdao, Shandong 266580, China
 ⁴China University of Geosciences, Wuhan, Hubei 430074, China

Correspondence should be addressed to Pingping Li; lpp@cup.edu.cn

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Thermal sulfate reduction (TSR) occurred throughout the Permian Changxing (P_2c) and Triassic Feixianguan (T_1f) dolostone reservoirs in the western and eastern parts of the Kaijiang-Liangping (K-L) trough in the northeastern part of the Sichuan Basin. To determine the sulfate sources of this TSR, fourteen solid bitumen samples and eight anhydrite samples were collected from the northeastern part of the Sichuan Basin. These samples were analyzed to determine their sulfur isotopes. In addition, untreated, HNO₃-treated, and CrCl₂-treated solid bitumen samples were analyzed to determine their sulfur isotopes in order to obtain reliable δ^{34} S data for the TSR solid bitumen. The results show that the HNO₃ method is more effective at removing pyrite from solid bitumen than the method using CrCl₂ thrice because the HNO₃-treated solid bitumen has lower sulfur contents and higher δ^{34} S of the T_1 f solid bitumen samples from the Puguang gas field (in the eastern part of the K-L trough, 12.0-24.0‰) is significantly lower than that of the samples from the Yuanba gas field (in the western part of the K-L trough, 24.1-34.2‰). The δ^{34} S of the T_1 f¹⁻² anhydrite is 18.1-26.6‰, which is lower than that of the T_1 f³⁻⁴ anhydrite samples (29.9-39.6‰). The TSR sulfates from the Puguang gas field were most likely from the coeval T_1 f¹⁻² evaporating seawater and were enriched during the reflux-seepage dolomitization process. The TSR sulfates from the Yuanba gas field were primarily caused by the evaporation of seawater during the T_1 f⁴. First, the evaporating seawater would flow vertically into the P₂c reservoirs in the adjacent area, and then, it would flow laterally into the P₂c reservoirs in the Yuanba gas field. Considering the fact that the sulfate sources of TSR and the δ^{34} S values of the TSR sulfates are different in the Puguang and Yuanba gas fields, the δ^{34} S of TSR solid bitumen cannot be simply used to show the extent of TSR.

1. Introduction

Thermal sulfate reduction (TSR) is the reaction of sulfate with a hydrocarbon (1) [1, 2]. As toxic and corrosive H_2S can be produced during TSR, the mechanism and products of TSR have been extensively studied [1–4]. The sulfur in the H_2S produced by TSR is derived from dissolved sulfate [1, 3, 5]. Different sulfate sources and whether the sulfates are sufficient or not can affect the H_2S concentration [3, 6] and sulfur isotopic composition of H_2S . Thus, determining the sulfate source is a key problem to predict the H_2S concentration and distribution.

$$\begin{array}{l} \text{Hydrocarbons} + \text{SO}_4{}^{2-} \rightarrow \text{Altered Hydrocarbons} \\ + \text{Solid Bitumen} + \text{CO}_2 + \text{H}_2\text{S} \end{array} \tag{1}$$

The sulfur in the sulfate is mainly transferred into sulfurrich solid bitumen, H_2S , and some solid sulfides (such as pyrite) during TSR [7, 8]. The sulfur isotope of sulfur-rich solid bitumen and H_2S is similar to the sulfur isotope of sulfate [9–11]. Thus, the sulfur isotopes of the sulfur-rich solid bitumen and the H_2S can be used to determine the sulfate source. The sulfur isotopes of the sulfur-rich solid bitumen can be obtained relatively easily compared to those of the toxic H_2S . As a result, we can use the sulfur isotopes of the sulfur-rich solid bitumen to determine the sulfate source.

Pyrite is a byproduct of TSR [1], and it usually resides within the solid bitumen. Thus, we can obtain accurate sulfur isotopes for the solid bitumen only after the pyrite has been removed from the solid bitumen. There are two main methods that can be used to remove pyrite from solid bitumen. One method is to use $CrCl_2$ to reduce the pyrite to H_2S (2) [12]. Cai et al. [13, 14] successfully removed pyrite from kerogen and solid bitumen by this method. The second method is to use dilute HNO₃ to oxidize the pyrite to H_2SO_4 (3) [15, 16]. The $CrCl_2$ method is more complicated than the HNO₃ method, but the HNO₃ method is time consuming. However, there is no research on the use of both methods simultaneously to remove pyrite from solid bitumen, nor has an experiment been conducted to compare the two methods and to determine which method is more effective.

$$FeS_2 + 2CrCl_2 + 4HCl \rightarrow FeCl_2 + 2CrCl_3 + 2H_2S$$
(2)

$$FeS_2 + 8HNO_3 \rightarrow Fe(NO_3)_3 + 2H_2O + 2H_2SO_4 + 5NO$$
(3)

The relatively high concentration of H_2S (5-20%) in the Permian Changxing (P_2c) and Triassic Feixianguan (T_1f) Formations in the northeastern part of the Sichuan Basin has been concluded to be from TSR [4, 6, 17-19], whereas sulfur-rich solid bitumen has been found in other gas reservoirs [6, 14, 20, 21]. The H₂S concentrations of the gas reservoirs in the eastern part of the Kaijiang-Liangping (K-L) trough (see detailed information in the background section) are higher than those of the gas reservoirs in the western part of the K-L trough [6]. Sulfate from the evaporative platform in T₁f has been suggested as the sulfate source of the TSR [4, 18]. However, whether the sulfate was derived from the early (the first and second members of the T_1f) or late T_1f (the third and fourth members of the T_1f) is unknown. In addition, there is no such evaporative platform in the P_2c in the western part of the K-L trough, so the sulfate source of the TSR in this area is also unknown.

The purpose of this paper is to determine which method is more effective at removing pyrite from solid bitumen by comparing the $CrCl_2$ and HNO_3 methods, to obtain reliable sulfur isotopic data for the solid bitumen, and to analyze the sulfate sources of TSR in the eastern and western parts of the K-L trough in the northeastern part of the Sichuan Basin based on the systematic correlation of the sulfur isotopes of the solid bitumen and anhydrite.

2. Geologic Setting

The Sichuan Basin is a rhombic basin in southwestern China (Figure 1). The general evolutionary history of the Sichuan Basin can be divided into two main stages [22]. Before the

Late Indosinian movement, the Sichuan Basin was primarily characterized by subsidence and uplift, and thick marine carbonates and shales were deposited at this stage. Since the Late Indosinian movement, large-scale lateral compression has been occurring, which has resulted in two obvious episodes of uplift, i.e., the Yanshan and Himalayan movements [22]. In addition, terrestrial fluvial-lacustrine deposits were formed.

The Permian strata primarily consist of marine carbonates and shales, and from bottom to top, it can be divided into the Liangshan (P1l), Qixia (P1q), Maokou (P1m), Longtan (P₂l)/Wujiaping (P₂w), and Changxing (P₂c) Formations [22]. The Early Triassic can be divided into the Fexianguan (T_1f) and Jialingjiang (T_1j) Formations, and the T_1f can be furtherly divided into four members (T_1f^1 , T_1f^2 , T_1f^3 , and T_1f^4) from bottom to top. During the P_2c , the northwest-southeast trending K-L trough developed in the northeastern part of the Sichuan Basin (Figure 2). An open platform and a relatively isolated open platform developed in the western and eastern parts of the K-L trough, respectively [23]. During the early T₁f, a restricted evaporative platform (characterized by anhydrite layers) developed in the eastern part of the K-L trough, whereas the western part of the K-L trough was still an open platform and contained no anhydrite layers. Until the end of the T₁f, the evaporative platform remained restricted and widespread anhydrite layers continued to develop in the northeastern part of the Sichuan Basin (Figure 3). The $T_1 j$ and Middle Triassic Leikoupo (T₂l) Formations developed on restricted evaporative platforms and consist of limestone and widespread anhydrite [22].

Many gas fields, such as the Puguang, Dukouhe, Tieshanpo, Yuanba, and Longgang, are located in the P_2c and T_1f reefs and shoals in the eastern and western parts of the K-L trough [24, 25]. The natural gas in these gas fields was primarily produced by oil-cracking and was primarily derived from the Upper Permian P_2l/P_2w source rocks [4, 18, 21].

3. Sampling and Analytical Methods

Fourteen solid bitumen-bearing core samples (5 samples from the T_1f in the Puguang gas field and 9 samples from the P_2c in the Yuanba gas field) were collected from the northeast part of the Sichuan Basin, and eight anhydrite samples (7 samples from the T_1f and 1 sample from the T_1j) were collected for this study. The solid bitumen was primarily present as fill in the pores, while the anhydrite primarily occurred as nodules and layers (Figure 4).

The anhydrite nodules and layers were crushed and powdered to less than 100 mesh using an agate mortar and pestle. The solid bitumen-bearing core samples were crushed in a rock crusher to less than 100 mesh, and then, this powder was treated with dichloromethane (DCM) to remove any soluble organic matter and elemental sulfur. The liquid and residue were filtered and flushed with DCM. The residue was dried and treated with hot 6 N HCl to remove carbonate minerals. This solution was heated for 2 hours at 70°C while using a magnetic stirrer. Next, the solid bitumen residue was filtered and rinsed with deionized water, and then, it



FIGURE 1: Main gas fields and sedimentary facies of the upper part of the Changxing Formation (P_2c) in the northeastern part of the Sichuan Basin (modified from [23, 24]). YB = Yuanba; LG = Longgang; HB = Heba; PG = Puguang; TSP = Tieshanpo; LJZ = Luojiazhai; DKH = Dukouhe; JZP = Jinzhuping; JST = Jiaoshutang; K-L = Kaijiang-Liangping; CK = Chengkou.

was dried at 70°C. The solid bitumen residue was mainly between 100 and 300 mg, so the pyrite content cannot be evaluated precisely using XRD (X-ray diffraction). However, pyrite grains can be easily found using the SEM (Scanning Electron Microscopy) and EDS (Energy Dispersive Spectrometer) analysis (Figure 5). As a result, it is necessary to remove pyrite to get accurate sulfur isotope of the solid bitumen. So the solid bitumen residue was treated with HNO₃ and CrCl₂ to remove the pyrite, respectively.

The HNO₃ method requires the solid bitumen residue to react with 1 N HNO_3 in a glass centrifuge tube at room temperature for two weeks. The centrifuge tubes were sonicated for 30 minutes every day. After reacting with 1 N HNO_3 for two weeks, deionized water was added to the tubes, they were centrifuged again, and the liquid was poured off. This

procedure was repeated twice more. Finally, the solid bitumen was dried at 70°C for 12 hours.

The $CrCl_2$ method used was similar to the method reported by Cai et al. [13, 14]. First, 12 N HCl and deionized water were added to solid $CrCl_3 \cdot 6H_2O$ to prepare a green $CrCl_3$ solution. Then, the green $CrCl_3$ solution was allowed to slowly flow past a flask column filled with granular zinc to convert it to a dark blue $CrCl_2$ solution, which was stored in a ground glass stoppered bottle to prevent oxidation. Next, 6 N HCl and the $CrCl_2$ solution were added to the solid bitumen residue in a glass tube, and then, the tube was covered with a plastic cap with a syringe needle, which allowed the generated H_2S to flow out. The tube was heated at 80°C for 3 hours, then deionized water was added, the tube was centrifuged twice, the liquid was poured off, and the solid bitumen



FIGURE 2: Key wells and sedimentary facies in the upper part of the second member of the Feixianguan Formation $(T_1 f^2)$ in the Sichuan Basin (modified from [24]).

residue was reground. Finally, this procedure was repeated twice, and the solid bitumen was dried.

The sulfur isotopes of the anhydrite powder, the untreated solid bitumen residue, and the HNO₃-treated and CrCl₂-treated solid bitumen were analyzed at the California Institute of Technology. Samples were combusted in a Costech elemental combustion system at 1000°C, and then, the isotopic ratios were determined using a Delta Plus XL mass spectrometer calibrated using International Atomic Energy Agency (IAEA) standards. The results are reported as δ^{34} S relative to the Vienna Canyon Diablo Troilite (VCDT) standard. Average precision of analyses based on replicate analyses of standards was ±0.2‰.

4. Results

4.1. Sulfur Content and Isotope Ratios of the Solid Bitumen Samples. The sulfur contents and sulfur isotopes (δ^{34} S) of the untreated solid bitumen, the HNO₃-treated solid bitumen, and the CrCl₂-treated solid bitumen are listed in Table 1. As the weight of the untreated solid bitumen is small, two samples (SB10 and SB12) were not treated with CrCl₂.

The sulfur contents of the untreated solid bitumen, the HNO_3 -treated solid bitumen, and the $CrCl_2$ -treated solid bitumen were different (Figure 6). Except for one sample

(SB16), the sulfur contents of the HNO₃-treated and $CrCl_2$ -treated solid bitumen samples were lower than those of the untreated solid bitumen samples. In addition, the sulfur contents of the HNO₃-treated and $CrCl_2$ -treated solid bitumen samples were similar, but the sulfur contents of three of the HNO₃-treated solid bitumen samples (SB2, SB5, and SB13) were lower than those of the $CrCl_2$ -treated solid bitumen samples (Figure 6).

Similarly, the δ^{34} S values of the untreated solid bitumen samples, the HNO₃-treated solid bitumen samples, and the CrCl₂-treated solid bitumen samples were different (Figure 7). Except for two samples (SB4 and SB11), the δ^{34} S values of the untreated solid bitumen samples were lower than those of the HNO3- and CrCl2-treated solid bitumen samples. In addition, the δ^{34} S values of the HNO₃- and CrCl₂-treated solid bitumen samples were similar. However, in the case of samples SB2, SB5, SB13, and SB17, the δ^{34} S values of the HNO3-treated solid bitumen samples were higher than those of the CrCl₂-treated solid bitumen samples (Figure 7), while the sulfur contents of the HNO₃-treated solid bitumen samples were lower than those of the CrCl₂treated solid bitumen samples (Figure 6). The δ^{34} S difference between these four HNO3-treated solid bitumen samples and corresponding untreated solid bitumen was 3.2‰, 4.4‰, 9.7‰, and 1.8‰, respectively, and the δ^{34} S difference

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FIGURE 3: The lithologic variations in the Changxing (P_2c) and Feixianguan (T_1f) Formations in the Sichuan Basin from southwest to northeast. The data for wells Z4, GC2, LG1, and JZ2 are from Du et al. [24]. The figures to the left of each lithologic column are the burial depths (in m). T_1f^{1-2} = the first and second members of T_1f ; T_1f^{3-4} = the third and fourth members of T_1f . The well locations are shown in Figure 2.

between these four $CrCl_2$ -treated solid bitumen samples and corresponding untreated solid bitumen were 1.2‰, 3.6‰, 7.3‰, and 1.6‰, respectively.

Overall, there is no obvious difference in the sulfur contents of the solid bitumen samples from the Puguang and Yuan gas fields, and the sulfur contents ranged from 4.0% to 15.0% (Figure 6), which are roughly similar to the sulfur content ranges reported for solid bitumen from the Puguang gas field (11.75-11.99%; [4]) and from the Yuanba gas field (6.64-17.99%; [6]). However, the δ^{34} S of solid bitumen from the Puguang gas field (generally lower than 20.0‰) is significantly lower than that of solid bitumen from the Yuanba gas field (generally higher than 24.0‰).

4.2. Sulfur Isotope Ratios of the Anhydrite Samples. The δ^{34} S of the eight anhydrite samples analyzed and the δ^{34} S reported for other anhydrites are listed in Table 2.

For the eight anhydrites analyzed in this study, the δ^{34} S of the T₁f¹ anhydrite (26.2–26.6‰) was lower than that of the T₁f³-T₁f⁴ anhydrites (29.9-34.4‰) and the T₁j anhydrite (31.7‰). The δ^{34} S of the T₁f¹ anhydrite is similar to the

 δ^{34} S range reported for the $T_1f^1-T_1f^3$ anhydrites (18.12-30.7‰; [19, 26]), and the δ^{34} S of the T_1f^4 anhydrite is close to that reported for the T_1f^4 anhydrite (35.88-39.64‰, [27]) and the T_1j anhydrite (24.7-32.5‰, [28]; 29.72‰, [27]). This suggests that the δ^{34} S values of the anhydrite samples are reliable. The δ^{34} S of the $T_1f^1-T_1f^3$ anhydrites (18.12-30.7‰) reported by Wang et al. [26] and Zhu et al. [19] is mainly for the $T_1f^1-T_1f^2$ because the anhydrite was mainly located within the $T_1f^1-T_1f^2$ in the Dukouhe, Tieshanpo, and Jinzhuping gas fields (Figure 3).

Overall, the δ^{34} S of the T₁f¹-T₁f² anhydrites (18.12-26.6‰) was lower than that of the T₁f³-T₁f⁴ anhydrites (29.9-39.6‰), which is consistent with the conclusion that the δ^{34} S of seawater sulfate increased during the Early Triassic [29].

5. Discussion

5.1. Which Method Is More Effective at Removing Pyrite from Solid Bitumen? Pyrite can be a byproduct of TSR [1, 8]. Pyrite grains were observed in the T₁f reservoir cores and have been





(d)





FIGURE 4: Photographs showing anhydrite (Anh) in an outcrop sample (a) and in cores (b-h), and black solid bitumen in thin section (i). (a) Outcrop JST, T_1f^4 , anhydrite nodules in dolostone. (b) Well HB1, 4847.6 m, T_1f^4 , anhydrite layers in mudstone. (c) Well HB1, 4501 m, T_1j^2 (second member of the Jialingjiang Formation), anhydrite nodules in dolostone. (d) Well MB6, 4015.9 m, T_1f^1 , anhydrite nodules in limestone. (e) Well XHC1, 5468.2 m, T_1f^1 , anhydrite nodules in dolostone. (f) Well DW2, 4750.7 m, T_1f^3 , anhydrite nodules in dolostone. (g) Well MB1, 4302.3 m, T_1f^4 , anhydrite layers in mudstone. (h) Well MB3, 3873 m, T_1f^4 , large anhydrite nodules in dolostone. (i) Well YB271, 6327.4 m, P_2c , black solid bitumen in the pores.



FIGURE 5: SEM (Scanning Electron Microscopy) and EDS (Energy Dispersive Spectrometer) photographs showing pyrites in the untreated SB2 (a, c) and SB13 (b, d) solid bitumen samples.

suggested to be formed by TSR because their δ^{34} S is 18.7-20.16‰ [19], which differs from that of the pyrite from the source rock layers. Except for one sample (SB16), the sulfur contents of the HNO₃-treated and CrCl₂-treated solid bitumen samples are lower than those of the untreated solid bitumen samples (Figure 6), which support the conclusion that some of the sulfur was removed by the HNO₃ and CrCl₂ methods. The untreated solid bitumen samples were extracted by DCM so that the organic matter and elemental sulfur were removed from the solid bitumen. As a result, the decrease in the sulfur contents of the HNO₃-treated and CrCl₂-treated solid bitumen samples was most likely caused by the removal of some of the pyrite. In addition, the special smell of released H₂S was smelled when the solid bitumen was treated with CrCl₂, and small pyrite grains were observed in the untreated solid bitumen powder with SEM (Figure 5). XRD had been used to determine the pyrite content in kerogen and solid bitumen [13, 14], and an improved method to determine the pyrite content by measuring the dissolved iron at pH < 2 using an atomic absorption spectrometer was developed by Cai et al. [30]. However, XRD and the dissolved iron analysis were not performed in this study because the obtained untreated solid bitumen samples weighed less than 100-300 mg each, so the pyrite content could not be

determined. However, it is reasonable that the decrease in the sulfur contents of the HNO_3 - and $CrCl_2$ -treated solid bitumen samples was caused by removing some of the pyrite.

Approximately 88.9-99.7% of pyrite was removed when the kerogen was treated twice with CrCl₂ [12]. Cai et al. [13, 14, 30] used a similar method and successfully removed the pyrite from kerogen and solid bitumen samples, and sulfur isotopes were analyzed only when pyrite sulfur/total sulfur values were less than 0.08 [30]. However, no studies have been conducted to compare the HNO₃ and CrCl₂ methods. In this study, the sulfur contents and sulfur isotopes of most of the solid bitumen samples treated using the HNO₃ and CrCl₂ methods were quite similar (Figures 6 and 7), which demonstrate that pyrite was removed by both methods. However, some of the HNO3-treated solid bitumen samples (SB2, SB5, SB13, and SB17) had significantly lower sulfur contents and significantly higher δ^{34} S than the CrCl₂-treated solid bitumen samples (Figure 6), which suggests that the HNO₃ method is more effective at removing pyrite from solid bitumen than the method using CrCl₂ thrice. In particular, the δ^{34} S of the CrCl₂-treated sample SB13 (16.8‰) is still less than 20.0‰, but it is 2.4‰ higher than that of untreated sample SB13. The δ^{34} S of HNO₃treated sample SB13 is 24.1‰, which is similar to the δ^{34} S

Gas field	Well	Sample no.	Strata	Depth (m)	Untreated SB S% δ^{34} S%		$CrCl_2$ -treated SB S% $\delta^{34}S\%$		HNO ₃ -treated SB S% δ^{34} S%	
Puguang	MB4	SB1	$T_1 f^2$	3814.5	12.41	17.6	10.47	17.5	11.04	17.9
Puguang	MB4	SB2	T_1f^2	3836.3	6.18	15.4	5.76	16.6	4.63	17.8
Puguang	PG12	SB3	T_1f^2	6007.5	8.17	16.9	7.52	18.1	7.58	17.9
Puguang	PG2	SB4	T_1f^1	5059.0	11.53	20.1	10.57	19.0	10.54	18.5
Puguang	DW102	SB5	T_1f^1	4819.7	5.68	13.3	5.04	14.1	3.51	17.7
Yuanba	YB29	SB7	P_2c	6642.4	12.4	27.7	11.6	27.9	12.1	28.5
Yuanba	YB205	SB8	P_2c	6461.4	9.37	34.2	7.92	34.3	7.89	34.2
Yuanba	YB27	SB10	P_2c	6295.3	11.53	25.9	/	/	11.08	26.5
Yuanba	YB27	SB11	P ₂ c	6301.1	10.17	31.1	8.87	29.8	8.89	29.4
Yuanba	YB273	SB12	P ₂ c	6826.2	10.34	25.1	/	/	10.29	25.6
Yuanba	YB271	SB13	P_2c	6320.5	9.87	14.4	7.55	16.8	4.52	24.1
Yuanba	YB28	SB15	P_2c	6807.3	11.75	30.5	11.6	29.8	11.1	30.3
Yuanba	YB224	SB16	P_2c	6627.9	10.34	25.2	13.4	25.4	12.77	25.9
Yuanba	YB224	SB17	P_2c	6641.2	14.57	24.7	14.28	24.9	13.66	26.5

TABLE 1: Sulfur content and isotope composition of solid bitumen from the Yuanba and Puguang gas fields in the Sichuan Basin.

Note: $P_2c = Changxing$ Formation (Fm); $T_1t^1 =$ the first member of Feixianguan Fm; $T_1t^2 =$ the second member of Feixianguan Fm.



FIGURE 6: Variation in the sulfur contents of the untreated, $CrCl_2$ -treated, and HNO_3 -treated solid bitumen (SB) samples.

of the other HNO₃-treated solid bitumen samples from the Yuanba gas field (greater than 24.0‰) (Figure 7). Therefore, the δ^{34} S of HNO₃-treated sample SB13 should be closer to that of solid bitumen without pyrite. In addition, the difference in sulfur content (Δ S) and the difference in the sulfur isotopes ($\Delta\delta^{34}$ S) of the HNO₃- and CrCl₂-treated solid bitumen samples are negatively correlated ($R^2 = 0.83$; Figure 8), indicating that the δ^{34} S of the HNO₃-treated solid bitumen samples increases more with decreasing sulfur content than the δ^{34} S of the CrCl₂-treated samples does. This also supports that the HNO₃ method is more effective at removing pyrite from solid bitumen than the method using CrCl₂ thrice.

Ås a result, the δ^{34} S of the solid bitumen discussed in the next section is the δ^{34} S of the HNO₃-treated solid bitumen

36 32 28 0 δ³⁴S (%0) 24 20 16 12 SB15 SB16 SB10 SB12 SB13 SB17 SB8 SB1 SB2 SB3 SB4 SB5 SB7 SB11 Untreated SB CrCl₂-treated SB O HNO3-treated SB

FIGURE 7: Variations in the sulfur isotopes of the untreated, CrCl₂-treated, and HNO₃-treated solid bitumen (SB) samples.

samples. Though we cannot be absolutely sure that all of the pyrite was completely removed from the solid bitumen by the HNO₃ method, we believe that the δ^{34} S of the HNO₃-treated solid bitumen samples is quite similar to the true δ^{34} S of solid bitumen without pyrite. Thus, the δ^{34} S of solid bitumen samples can be used to investigate the origin of the sulfur.

5.2. Sulfate Sources of TSR. The solid bitumens from the T_1f and P_2c reservoirs in the northeastern part of the Sichuan Basin are insoluble pyrobitumens, which were formed in the advanced stages of thermal maturity [4, 21]. Insoluble solid bitumen can form due to the thermal chemical alteration (TCA) or thermochemical sulfate reduction (TSR) of migrated petroleum [31]. The S/C ratio of TSR bitumen is

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TABLE 2: Sulfur isotope compositions of anhydrite from the northeastern Sichuan Basin.

Area/gas field	Well	Strata	Depth (m)	Sample description	δ^{34} S/‰
XHC	XHC1	$T_1 f^1$	5468.2	Anhydrite layer	26.2
PG	MB6	T_1f^1	4015.9	Anhydrite nodule	26.6
PG	DW2	$T_1 f^3$	4750.7	Anhydrite nodule	29.9
HB	HB1	$T_1 f^4$	4847.6	Anhydrite layer	34.3
PG	MB1	T_1f^4	4302.3	Anhydrite nodule	34.2
PG	MB3	T_1f^4	3873.0	Anhydrite nodule	34.4
JST	JST	T_1f^4	1	Anhydrite nodule	34.0
HB	HB1	$T_1 j^2$	4501.0	Anhydrite layer	31.7
DKH*	Du3	$T_1 f^{1-3}$	1	Anhydrite layer	26.5
DKH*	Du3	$T_1 f^{1-3}$	/	Anhydrite nodule	30.0
DKH*	D4	$T_1 f^{1-3}$	1	Anhydrite nodule	30.7
TSP*	Po1	$T_1 f^{1-3}$	1	Anhydrite nodule	30.1
TSP*	Po1	$T_1 f^{1-3}$	1	Anhydrite nodule	30.6
$HB^{\#}$	HB102	$T_1 f^3$	5176.8	Anhydrite vein	34.3
PG [#]	MB6	T_1f^1	4015.7	Anhydrite nodule	26.48
$HB^{\#}$	HB1	T_1f^4	4847.3	Anhydrite layer	35.88
$PG^{\#}$	MB2	T_1f^4	4144.0	Anhydrite layer	39.33
JZP [†]	Z2	T_1f	3416.8	Anhydrite	19.71
LJZ^{\dagger}	Lj2	$T_1 f^{1-3}$	2897.2	Anhydrite	22.59
DKH^\dagger	Du3	$T_1 f^{1-3}$	4290.0	Anhydrite	18.12
$\rm DKH^\dagger$	Du5	$T_1 f^{1-3}$	4740.3	Anhydrite	24.34
DKH^\dagger	Du5	$T_1 f^{1-3}$	4742.3	Anhydrite	22.8
$\rm DKH^\dagger$	Du5	$T_1 f^{1-3}$	4753.4	Anhydrite	22.83
TSP^\dagger	Po1	$T_1 f^{1-3}$	3302.7	Anhydrite	19.46
TSP^\dagger	Po3	$T_1 f^{1-3}$	3464.7	Anhydrite	18.92
QLX^{\dagger}	QL52	$T_1 f^{1-3}$	3536.0	Anhydrite	24.64
QLX^{\dagger}	QL52	$T_1 f^{1-3}$	3490.4	Anhydrite	23.57
JZP^{\dagger}	JZ1	$T_1 f^{1-3}$	2825.8	Anhydrite	19.35
JZP [†]	JZ1	$T_1 f^{1-3}$	2877.2	Anhydrite	22.07
JZP [†]	JZ1	$T_1 f^{1-3}$	4766.0	Anhydrite	22.13
JZP [†]	ZJ1	$T_1 f^{1-3}$	3941.9	Anhydrite	23.74
JZP [†]	Z1	$T_1 f^{1-3}$	3350.5	Anhydrite	18.09
JZP [†]	Z1	$T_1 f^{1-3}$	5648.9	Anhydrite	25.4
$PG^{\dagger\dagger}$	MB3	$T_1 f^4$	3872.0-3880.2	Anhydrite vein	39.64
$PG^{\dagger\dagger}$	MB3	$T_1 f^4$	3872.0-3880.2	Anhydrite vein	38.56
PG ^{††}	CY83	$T_1 j^4$	3552.0-3554.0	Anhydrite layer	29.72

*from Wang et al. [26], [#]from Zhang [39], [†]from Zhu et al. [19], ^{††}from Zhu et al. [27].

generally greater than 0.03 [31], while its δ^{34} S is close to the δ^{34} S of sulfate [11, 32, 33]. The S/C ratio of TCA bitumen is generally less than 0.03 [31], while its δ^{34} S is similar to the δ^{34} S of the source kerogen [20]. The S/C ratios of the T₁f and P₂c solid bitumen from the Puguang and Yuanba gas fields are greater than 0.03 [4, 6]. In addition, the δ^{34} S of this solid bitumen is 12.0-35.0‰ (Figure 9), which is significantly higher than that of the kerogen from the Permian source rocks from well HB1 (-26.7‰; [30]), but it is close to the δ^{34} S of the T₁f anhydrite (18.0-35.0‰; Figure 9). The S/C ratios and the δ^{34} S values of the solid bitumen both support its TSR origin.

The mechanism by which the sulfur in the sulfate is transferred into the solid bitumen is not yet clearly understood. However, when the TSR proceeds gradually, the saturate/aromatic ratio of the TSR-altered oil decreases, while the sulfur and oxygen contents of the solid bitumen increase [32, 34], which suggests that the sulfur was transferred from the sulfate into the hydrocarbon. In addition, the TSR-altered oils are rich in sulfur compounds, such as thiophenes, benzothiophens, and dibenzothiophenes [32, 33]. The isotopic fractionation of sulfur that occurs during the TSR process is minimal, so the δ^{34} S of the TSR-altered oil and solid bitumen is close to that of the TSR-involved sulfate [9–11, 28, 32, 33].



FIGURE 8: Variation in differences in the sulfur contents (Δ S) and the differences in the sulfur isotopes ($\Delta \delta^{34}$ S) of the HNO₃-treated and CrCl₂-treated solid bitumen samples.

TSR can occur in both gas zone and gas-water transition zone [17], but the sulfur fractionations were different in these two zones [14]. In the gas-water transition zone, the gas saturation was low so that some dissolved sulfate remains unreacted, and the TSR solid bitumen would be enriched in δ^{34} S and H₂S would be depleted in δ^{34} S. In the gas zone, the limited dissolved sulfate would be reduced completely, so that the δ^{34} S of TSR solid bitumen would be close to the δ^{34} S of sulfate. However, the solid bitumen samples from the Yuanba and Puguang gas fields were both collected from present gas zones, so we considered that the sulfur fractionation during the TSR was insignificant.

As shown in Figure 9, the δ^{34} S of the solid bitumen from the Puguang gas field in the eastern part of the K-L trough is generally less than 24.0‰ (except for one sample with 31.2‰), whereas the δ^{34} S of the solid bitumen from the Yuanba gas field in the western part of the K-L trough is generally greater than 24.0‰ (24.1-34.2‰). Thus, assuming minimal isotopic fractionation of sulfur during TSR, it is impossible for the sulfur in the solid bitumen from the Yuanba gas field to have come from the T_1f^1 and T_1f^2 sulfates because the δ^{34} S of these sulfates is lower than that of the solid bitumen. However, the δ^{34} S of the solid bitumen is similar to that of the $T_1 f^3$ and $T_1 f^4$ anhydrites. In addition, there is no anhydrite in T_1f^1 and T_1f^2 in the Yuanba gas field, and the anhydrite is only present in T_1f^4 (Figure 3). As a result, the sulfate of the TSR was most likely from T₁f⁴ in the Yuanba gas field. The primary source of the sulfate of the TSR in the Puguang gas field was the T_1f^1 and T_1f^2 anhydrite layers, which have slightly higher δ^{34} S than the solid bitumen, but part of the sulfate was from the T_1f^3 and T_1f^4 anhydrite layers because the δ^{34} S of some of the solid bitumen is higher than that of the T₁f¹ and T₁f² anhydrites (Figure 9). Based on the correlation between the δ^{34} S values of the solid bitumen and anhydrite samples, it can be concluded that the sulfate source of the TSR is different in the western and eastern parts of the K-L trough.

The H₂S-rich natural gases of the T₁f and P₂c are mainly located in the dolostone reservoirs in the northeastern part of the Sichuan Basin, whereas no H₂S was found in limestone reservoirs [6]. Thus, we suggest that the TSR sulfates (including anhydrites and dissolved sulfates) were enriched during the dolomitization process. Geological observations and numerical simulations have demonstrated that anhydrite can be a byproduct of the dolomitization process [35, 36], which indicates that sulfates can be enriched during the dolomitization process. According to the calculations by Li et al. [6], if all of the SO_4^{2-} in the formation water were converted to H₂S during TSR, the generated H₂S concentration would be 8.8-23.3%, which is similar to the range observed in the Puguang (5.09–19.22%, [4]) and Yuanba (1.20-12.16%, [6]) gas fields. In addition, the anhydrite produced in the dolomitization process can also supply additional sulfates. TSR calcites have been observed to replace anhydrite in the T₁f dolostone reservoir in the northeastern part of the Sichuan Basin [19]. Therefore, the sulfates enriched during the dolomitization process can supply sufficient sulfur to produce the observed H₂S concentration in the northeastern part of the Sichuan Basin.

The sulfates of the TSR in the Puguang gas field in the eastern part of the K-L trough were most probably enriched during the reflux-seepage dolomitization process [37–39]. The coevally evaporating seawater of the T₁f flowed laterally into the porous T₁f oolitic limestone and the P₂c reef limestone, causing broad dolomitization, during which the TSR sulfates (anhydrites and dissolved sulfates) were enriched. The down migration model of dolomitization fluid from T_1f and P_2c had been supported by the evidence that the P_2c dolostones have ${}^{87}Sr/{}^{86}Sr$ ratios close to the ratios of T_1f seawater but heavier than P_2c seawater [38]. This is similar to the Permian Khuff Formation, offshore Dubai [40]. Finally, the sulfur from the sulfates was transferred into the solid bitumen and the H_2S . The SO_4^{2-} concentration of the formation water from the Puguang gas field (H₂S concentration of 10-20%) is less than 1.0 g/L, which is far lower than that of the formation water from the Jiannan gas field (H₂S concentration less than 5%), indicating that the dissolved sulfates in the formation water were consumed during the TSR process in the Puguang gas field [6].

As was previously discussed, the sulfates of the TSR in the Yuanba gas field were most likely from the evaporating seawater in T_1f^4 . However, the micritic limestone of T_1f^1 and T_1f^2 can act as a barrier to prevent evaporating seawater of T_1f^4 from flowing vertically into the P_2c reservoirs (Figure 3). However, there are porous dolostone reservoirs in T_1f^1 and T_1f^2 in the Longgang gas field (Figure 3), which is also in the western part of the K-L trough and adjacent to the Yuanba gas field, so we suggest that the evaporating seawater may have initially flowed vertically from T_1f^4 into T_1f^1 and T_1f^2 , then it flowed into the P_2c reservoir layers in the Longgang gas field, and finally, it flowed laterally into the P_2c reservoirs in the Yuanba gas field, which caused the dolomitization and enrichment of the sulfates of the TSR.

Cai et al. [30] proposed that the δ^{34} S and S/C ratio of the TSR solid bitumen samples can be used to show the TSR extent based on systematic analysis of TSR and non-TSR



FIGURE 9: Correlation between the sulfur isotopes of the solid bitumen and anhydrite samples in the northeastern part of the Sichuan Basin.

solid bitumen samples in the northeastern part of the Sichuan Basin. However, the δ^{34} S values of the P₂c solid bitumen in the western part of the K-L trough, such as those from the Yuanba and Longgang gas fields, are not included in that paper. As was previously discussed, the δ^{34} S of the solid bitumen from the Yuanba gas field is significantly higher than that from the Puguang gas field. However, the S/C of the solid bitumen from the Yuanba gas field (less than 0.06) is less than the S/C of the solid bitumen from the Puguang gas field (greater than 0.06). Also, the H₂S concentration of the Yuanba gas field (5-10%) is lower than that of the Puguang gas field (10-20%), and the corresponding GSI (gas sourcing index), which was used to reflect the extent of TSR [41], is lower in the Yuanba gas field (less than 0.1) than in the Puguang gas field (0.1-0.2). Thus, we conclude that the extent of TSR was greater in the Puguang gas field than in the Yuanba gas field. As a result, we suggest that the δ^{34} S of TSR solid bitumen can be used as an indicator of the TSR extent only when the sulfate source of the TSR is similar (δ^{34} S of sulfate source is similar).

6. Conclusions

The following conclusions can be drawn based on our study:

- Both the HNO₃ method and the CrCl₂ method successfully remove pyrite from solid bitumen, but the HNO₃ method is more effective than the method using CrCl₂ thrice
- (2) The δ^{34} S of the T₁f solid bitumen from the Puguang gas field (12.0-24.0‰) is lower than the δ^{34} S of the

 P_2c solid bitumen from the Yuanba gas field (24.1-34.2‰), and the sulfates of the TSR of the Puguang gas field were mainly formed by evaporating seawater in T_1f^{1-2} , whereas the sulfates of the TSR in the Yuanba gas field were mainly formed by evaporating seawater in T_1f^4

(3) When the δ^{34} S of the sulfates of the TSR is similar, the δ^{34} S of the TSR solid bitumen can be used to show the extent of TSR

Data Availability

The sulfur isotope composition of solid bitumen and anhydrite data used to support the findings of this study are included within the article.

Conflicts of Interest

The authors declare that they have no conflicts of interest.

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