

Research Article

Permeability Experiment of Fractured Rock with Rough Surfaces under Different Stress Conditions

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To investigate the permeability changes and the mechanisms of fractured rock under dynamic and static stresses produced by earthquakes, permeability experiments on fractured rock with rough surfaces under axial dynamic and static stresses were conducted on the MTS815 Rock Mechanics Testing System. Surface asperity was investigated by scanning the specimen surfaces before and after testing. The results show that the roughness of fracture surface has a great influence on the permeability when the axial displacement is not enough to cause the fracture rock to slip. Moreover, the rougher fracture surface leads to severer surface damage as indicated by the more gouge productions. The accumulation of gouge materials on larger roughness fracture surfaces causes a slow drop in permeability. The fracture surfaces experience larger degradations, but it has small weights of gouge materials on fracture surface after testing under axial dynamic stress. The reason is that the gouge material transport and mobilization tend to occur in process of dynamic loading. Therefore, the permeability drops of axial dynamic stress are larger than those of axial static stress.

1. Introduction

Permeability of rock fracture is governing fluid filtration rates and particle mobilization [1] and the main parameter of the safety properties of fractured rock mass [2]. Permeability varies with the changes of roughness [3] and the production of gouge materials [4], when the deformation of fracture occurs under stress [5–7]. It is known that fractured rock is often subjected to significant dynamic and static stresses produced by earthquakes [8–19]. Therefore, the permeability changes and the mechanisms of fractured rock under dynamic stress and static stress are very important to predict seismic activity.

Permeability changes caused by various mechanisms, including the unclogging and clogging of fractures, variations in the fracture aperture, and the particle mobilization, under static stress have been observed in both the field and the laboratory. Some researchers found the permeability of fracture decreases with the increase effective static stress [20–26]. Vogler et al. found that gouge material production may have

caused clogging of the main fluid flow channels, resulting in reductions of permeability by up to one order of magnitude [4]. Zhao et al. investigated that the permeability decreases with the fracture apertures decrease in effective stress, and the effect of fracture roughness on the permeability is related to the magnitude of effective stress [27]. Wu et al. measured the roughness and the permeability drop with effective stress. The permeability of natural fracture only partially recovers after effective stress returns to initial value and decreases due to produced gouge blocking fluid flow pathways [3].

Some studies suggested that dynamic stress can lead to the severe fatigue damage or failure of rock even when the stress level is significantly lower than the static strength [28–31] and can change permeability. Brodsky [32] and Elkhoury et al. [33] found that distant earthquakes may even increase the permeability in faults by unclogging of fractures in the field. Xue et al. [34] and Shi and Wang [35] studied that after a large earthquake, the fault zone permeability transiently increases because of earthquakes generating fractures

in a damage fault in Wenchuan earthquake. Faoro et al. [36] showed transient increase in the effective permeability of the rock mass caused by dilates in the fracture aperture with fluid pressure at laboratory scale. Candela et al. [37] conducted that transient fluid pressure in fractured rock commonly involves permeability increases and has been attributed to mobilization of fine particles in laboratory experiments. In addition, a few studies have detected that earthquakes could decrease permeability [38–40]. For example, Shi et al. found the earthquake-induced permeability decrease in the fault zone reduced the recharge from deep hot water [39]. Given the azimuthal distribution of distant earthquakes, the observed permeability decrease could be attributed to the seismic wave-induced clogging of fractures that compose the flow paths in the shallow crust [40]. Shmonov et al. [8] applied oscillatory stresses to unfractured cores at high confining pressures and temperatures and found that permeability is more likely to increase. Liu and Manga [41] conducted similar experiments on already fractured sandstone cores saturated with deionized water and showed that the permeability in fractured sandstone can potentially decrease under the effect of dynamic stresses. Various mechanisms have been tried to explain the increases in permeability caused by dynamic stresses induced by the passage of seismic waves [40]. However, these mechanisms are poorly understood, and thus, it is difficult to predict permeability increase or decrease under dynamic stress and static stress.

In this study, the experiments of effect of axial dynamic stress and static stress on permeability were conducted on the MTS815 Rock Mechanics Testing System. The fracture roughness change was investigated by scanning the sample surfaces before and after testing. The permeability characteristics under axial dynamic stress and axial static stress were analyzed considering the asperity degradations and gouge materials of fracture surfaces. This provides the information on the impact of production, transport, and flow mobilization of gouge material on permeability changes under dynamic stress.

2. Test Preparation

2.1. Sample Preparation. The rock material used in this study is a fine-grained sandstone collected from the northwest of Kunming, Yunnan province of China. A series of preliminary tests have been conducted on standard rock samples to determine the crucial mechanical parameters of the sandstone, such as Young's modulus (34 GPa), Poisson's ratio (0.3), and uniaxial compressive strength (86 MPa).

All samples were extracted from one single sandstone slab to minimize the variations in properties [42]. Then, samples were manufactured into cylindrical geometry with 50 mm in diameter and 100 mm in length. Each sample was split at a 30-degree angle with respect to the axis to form a fracture surface, as shown in Figure 1(a). Subsequently, one borehole with 3 mm diameter was drilled parallel to the axial direction, at the corner of each half to facilitate fluid flow from core holders into the fracture (Figure 1(b)). After that, samples were saturated by soaking in water for more than 12 hours.

2.2. Surface Roughness Measurements. Fracture surface characteristics were scanned with an optical three-dimensional scanner manufactured by GOM (ATOS III TRIPLE SCAN). The ATOS Core sensor projects fringe patterns on the object surface, which are recorded by two cameras. The patterns form a phase shift that is based on a sinusoidal intensity distribution which enables one to calculate the three-dimensional (3D) surfaces. The photogrammetry scanner is calibrated with two tests. The diameter and shape of a sphere and the distance between two spheres that are mounted on a plate are measured with the photogrammetry scanner to derive calibration errors and accuracy. All equipment used for calibration are specifically developed by the company GOM, which manufactures the scanner. The overall scanning accuracy is less than 0.01 mm within the scanning range of $100 \times 75 \text{ mm}^2$, and length deviation errors are between 0.009 and 0.027 mm. The measurement resolution is 3692×2472 pixel with optimized calibration deviations of 0.014 ± 0.001 pixels. In addition, the tensile fractures were prepared carefully with a high degree of fracture matching; therefore, the joint matching coefficient of the fractures is close to 1.0. So we only use one fracture surface of each sample in the scanning contour before and after testing. After the 3D scanning, the digitized data were exported in xyz file format to estimate the fracture roughness. The surface roughness parameter, Z_2 , the root mean square of the slope of a 2D profile, is widely used to correlate with the JRC value [43–46]. For a 2D profile, Z_2 is given by

$$Z_2 = \left[\frac{1}{(n-1)(\Delta x)^2} \sum_{i=1}^{i=n-1} (Z_{i+1} - Z_i)^2 \right]^{0.5}, \quad (1)$$

$$\text{JRC} = 61.79Z_2 - 3.47, \quad (2)$$

where Z_2 is the root mean square of the slope of a given 2D profile, n is the number of data points along the 2D profile, Δx is the interval between the data points, Z_i is the value of the asperity height at point i , and JRC is the joint roughness coefficient.

As in many previous studies, an interval of 0.5 mm for sampling points was selected to estimate the roughness [43–46]. JRC of each profile on the fracture surface was calculated using equation (2), and the mean value of JRC was calculated to characterize the roughness of the fracture surface as listed in Table 1. They are a set of 10 typical roughness profiles as shown in Figure 2. The mean JRC values of samples Ss-1 and Sd-1 belong to the roughness profile type 4, that of samples Ss-2 and Sd-2 is the roughness profile type 5, and that of samples Ss-3 and Sd-3 is of the roughness profile type 10. The same process was repeated for the damaged surfaces after testing to analyze the surface changes that occurred during the experiments and compare these to the changes in permeability during the experiment.

2.3. Experimental Setup and Procedure. All experiments were conducted on a servo-controlled Rock Mechanics Testing System (MTS815) housed at the Advanced Research Center

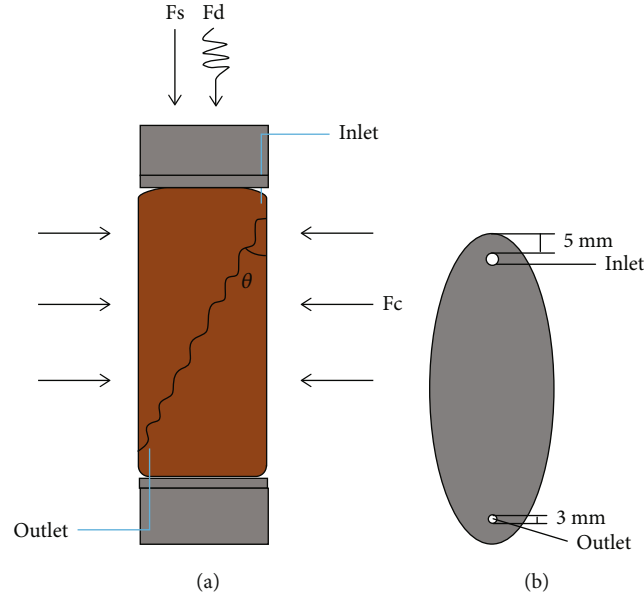


FIGURE 1: (a) Sketch of fractured rock sample and (b) a rough fracture with boreholes for fluid flow.

in Central South University. The system consists of five main units: a triaxial cell, a loading unit, a water supply unit, a deformation and pressure monitoring unit, and a data-acquisition unit. The maximum loading capacity of the system is up to 4600 kN. The maximum confining pressure and pore water pressure are 140 MPa. The axial deformation of the sample is measured by a pair of linear variable displacement transducers (LVDTs). The system is configured with a transient pulse apparatus for the permeability test of cylinder rock specimen. The permeability was measured with a pressure pulse-decay technique testing method as shown in Figure 3.

Two suites of permeability experiments on fractured rock were conducted. In the “axial dynamic test” suite of experiments, three groups (Sd-1, Sd-2, and Sd-3) of tests were designed. The experimental procedure includes the following steps:

- (1) Each sample was wrapped in a Teflon tape to avoid the slippage in the opening fracture when exerting the confining stress and preaxial static stress. Then, the sample was circumferentially sealed in a thermo-shrinking plastic membrane to separate the sample from the confining fluid. After that, the rock sample was placed at the triaxial cell filled with hydraulic oil
- (2) The confining stress (σ_3) and static axial stress (σ_s) were successively increased to the designed levels at a constant loading rate of 0.1 MPa/s. They were, respectively, set at 5 MPa and 10 MPa in all tests and kept constant during the whole test
- (3) Five minutes later, the sample and the loading system became stable. The successive axial sine wave was applied to the top of the sample by a

rigid loading bar. The actual axial stress is the superposition of the static prestress and the cyclic stress as

$$\sigma_{sd} = \sigma_s + \sigma_d \sin(2\pi ft), \quad (3)$$

where σ_{sd} is the superimposed axial stress, σ_s is static axial stress, σ_d is the amplitude of dynamic axial stress, f is the frequency, and t is the time. Preliminary compression tests on the fracture rock sample suggested that the critical axial stress for the slippage of the sample under the confining pressure of 5 MPa is 15 MPa. Hence, the superimposed axial stress should be limited below the critical stress. In this study, the axial stress was independent variables as listed in Table 1. Three groups of tests were designed according to the amplitude of dynamic stress, i.e., 0.25, 1.25, 2.5, 3.75, and 5 MPa. The loading paths with the frequency of 1 Hz were plotted in Figure 4

- (4) After the sample was subjected to 100 cycles of dynamic disturbance, the cyclic loading stopped, and then, the confining pressure and static axial stress were still maintained at 5 MPa and 10 MPa. Then, an initial water pressure was applied to both the upstream and downstream reservoirs, at a loading rate of 0.2 MPa/min. Next, the water pressure in the upstream reservoir suddenly increased to form a differential pressure (i.e., an initial pulse pressure) which makes the water flow from the top to the bottom through the fracture, as shown in Figure 3. The time pulse pressure (ΔP) decreased over time until equilibrium was attained. It was automatically monitored and recorded by two pressure gages in the water

TABLE 1: Test parameter for stress conditions.

Sample	Mean JCR value	Roughness profile type	Subset	σ_3 (MPa)	Fs (MPa)	Fd (MPa)	Cycles	f (Hz)
Sd-1	6.26	4	1	5	10	0.25	100	1
			2			1.25		
			3			2.5		
			4			3.75		
			5			5		
Sd-2	9.256	5	1	5	10	0.25	100	1
			2			1.25		
			3			2.5		
			4			3.75		
			5			5		
Sd-3	19.35	10	1	5	10	0.25	100	1
			2			1.25		
			3			2.5		
			4			3.75		
			5			5		
Ss-1	7.025	4	1	5	5	0	100	0
			2		7.5			
			3		10			
			4		12.5			
			5		15			
Ss-2	9.135	5	1	5	5	0	100	
			2		7.5			
			3		10			
			4		12.5			
			5		15			
Ss-3	20	10	1	5	5	0	100	0
			2		7.5			
			3		10			
			4		12.5			
			5		15			

tank. Accordingly, the permeability of the fractured sample can be obtained (see Section 2.4)

In the “axial static test” suite of experiments, three groups (Ss-1, Ss-2, and Ss-3) of tests were also designed. In each experiment, the axial stress was raised from 5 to 15 MPa in step of 2.5 MPa, and the confining pressure was kept at 5 MPa as listed in Table 1. The permeability of fractured rock was measured at each axial loading.

2.4. Permeability Measurements. Rock permeability can be measured using a steady state method, in which the flow rate of a fluid through a sample is measured for a known hydraulic pressure gradient. When permeability is low, however, a long time may be required to establish a steady state. This method is based on the analysis of the decay of a small-step change of the pressure imposed at one end of a specimen.

In this study, the differential pressure along the fault plane after achieving a pressure step is measured; then, the permeability is calculated as [3]

$$k = \frac{cL\mu}{A_2 P_m (1/V_1 + 1/V_2)}, \quad (4)$$

where P_m is the average value of P_{V_1} and P_{V_2} , P_{V_1} and P_{V_2} are the upstream and downstream pressure, μ is the water viscosity (1.01×10^{-3} Pa · s), L is the distance between the two drilled boreholes along the fault plane, A_2 is the cross-sectional area of the fracture, V_1 and V_2 are the upstream and downstream reservoir volumes ($V_1 = V_2 = 3.32 \times 10^{-7}$ m³ for the MTS experimental setup as shown in Figure 3), and c is the rate of the variation rate of the time pulse pressure with time, which can be determined according to the evolution of the time pulse pressure (ΔP). Brace et al. [47] and Jang et al.

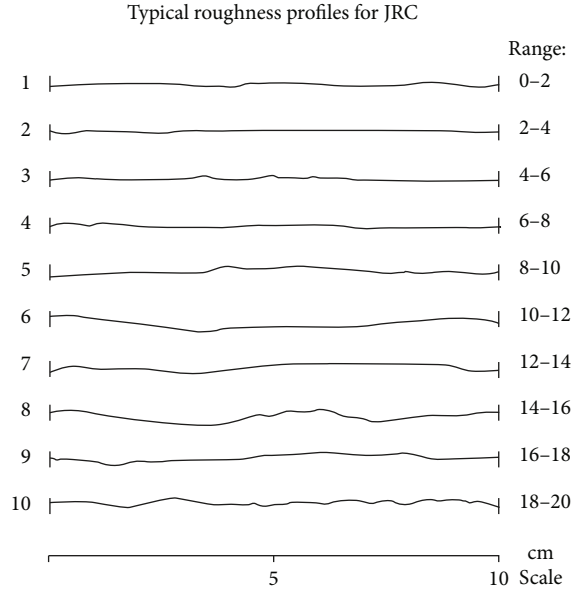


FIGURE 2: 10 typical roughness profiles showing the range of JRC values [44].

[48] suggested that the time pulse pressure exponentially attenuates with time as

$$\Delta P(t) = P_0(e^{-ct}) + c_0, \quad (5)$$

where P_0 is the initial pulse pressure at time $t = 0$ and c_0 is a fitting constant. Figure 5 presents the variation of the time pulse pressure versus time. It can be seen that the decay model of the time pulse pressure (equation (5)) matches the trend of test data well. Clearly, the coefficient c is 0.0021.

3. Results and Discussions

Our data shows that (1) the permeability changes are negatively correlated with the axial stress, (2) the permeability changes are affected by the roughness of fracture surfaces, and (3) the fractures of dynamic stress have slightly higher permeability drops compared to the fracture of static stress. We consider three influencing factors for transient permeability changes: (i) initial mean value of JRC, (ii) asperity degradation and gouge materials, and (iii) particle mobilization.

3.1. Permeability Changes. Figure 6 depicts the permeability evolution and axial displacements under axial dynamic stress. Figure 6 shows the permeability variation of Sd-1 and Sd-2 decreases at similar rates, while the axial displacements of Sd-1 and Sd-2 increase at similar rates. The permeability of Sd-3 drops slower and the axial displacement rises faster. Overall, three groups show an apparent reduction in the permeability with the rising amplitude under axial dynamic stress, an enhancement in axial displacement with that. This phenomenon does not match those observed in earlier studies that the decrease of permeability is sharp with increasing slip displacement [3, 4, 49, 50]. Figure 7 indicates the changes of permeability before and after testing under

axial dynamic stress and axial static stress. It is found that the permeability of Sd-1 and Sd-2 drops with 35% and 37% decline rate, and that of Sd-3 goes down 44% as presented in Figure 7. However, the mean JRC values of samples Sd-1 and Sd-2 belonging to the roughness profile types 4 and 5 are lower than those of sample Sd-3 in Table 1.

The similar phenomenon occurs under static stress. The changes of permeability and axial displacements of Ss-1 and Ss-2 are similar, while the permeability of Ss-3 decreases slowly and the axial displacements increase largely as presented in Figure 8. The roughness of fracture surfaces of Ss-1 and Ss-2 is smaller than that of Ss-3 under axial static stress (Table 1). The permeability values of Ss-1, Ss-2, and Ss-3, respectively, decrease with the decline rates 13%, 14%, and 20% in Figure 9. Figures 7 and 8 show that all the axial displacements are less than 1 mm and the fracture rocks are not damaged in two suites of permeability experiments. This indicates that when the axial displacement is not enough to cause the fracture rock to slip under the same stress conditions, the roughness of fracture surface has a great influence on the permeability.

3.2. Topographic Changes of Fracture Surfaces. Due to the large axial forces, significant surface deformation is expected in the fracture plane of the sample. For detailed analysis, the photogrammetric surface scans of the fracture sample and the mean values of JRC are generated before and after two suites of permeability experiments as shown in Figures 10–13. Photogrammetric scans produced profiles of the surfaces, which are oriented according to a best-fit plane in the $x-y$ coordinates. For visualization and comparison, an asperity height of 0 mm is assigned to the lowest point of the surface for Figures 10 and 12. Note that the same color scales are used. The changes of the mean values of JRC before and after testing under dynamic stress and that under static stress are, respectively, presented in Figure 13.

Figures 10 and 11 show the surface scans and the changes of the mean values of JRC of samples Sd-1, Sd-2, and Sd-3 before and after testing under axial dynamic stress. Comparing the fracture surfaces before and after testing, the maximum surface reliefs of Sd-1, Sd-2, and Sd-3 are, respectively, from 3.635 mm to 3.406 mm, 8.897 mm to 6.625 mm, and 5.448 mm to 4.835 mm in asperity height (Figure 10). Figure 11 indicates that the mean values of JRC of Sd-1 and Sd-2 are, respectively, from 6.26 to 3.45 and from 9.26 to 4.01 with the reductions of 2.81 and 5.25, while the mean values of JRC of Sd-3 drop from 19.35 to 6.42 with a higher reduction of 12.93. Therefore, the larger fracture roughness of Sd-3 is more decreased. This means that fracture surfaces with a high degree of relief and a large mean value of JRC variability experience more degradation during testing.

Under axial static stress, the initial maximum surface reliefs of Ss-1, Ss-2, and Ss-3 are 4.413 mm, 9.688 mm, and 5.443 mm and the initial mean values of JRC of Ss-1, Ss-2, and Ss-3 are 7.03, 9.14, and 20 as shown in Figures 12 and 13. After testing, the mean values of JRC are 4.75, 5.67, and 13.15. This means that the JRC degradations of 2.28, 3.47, and 6.85 are induced on the fracture surfaces under static stress.

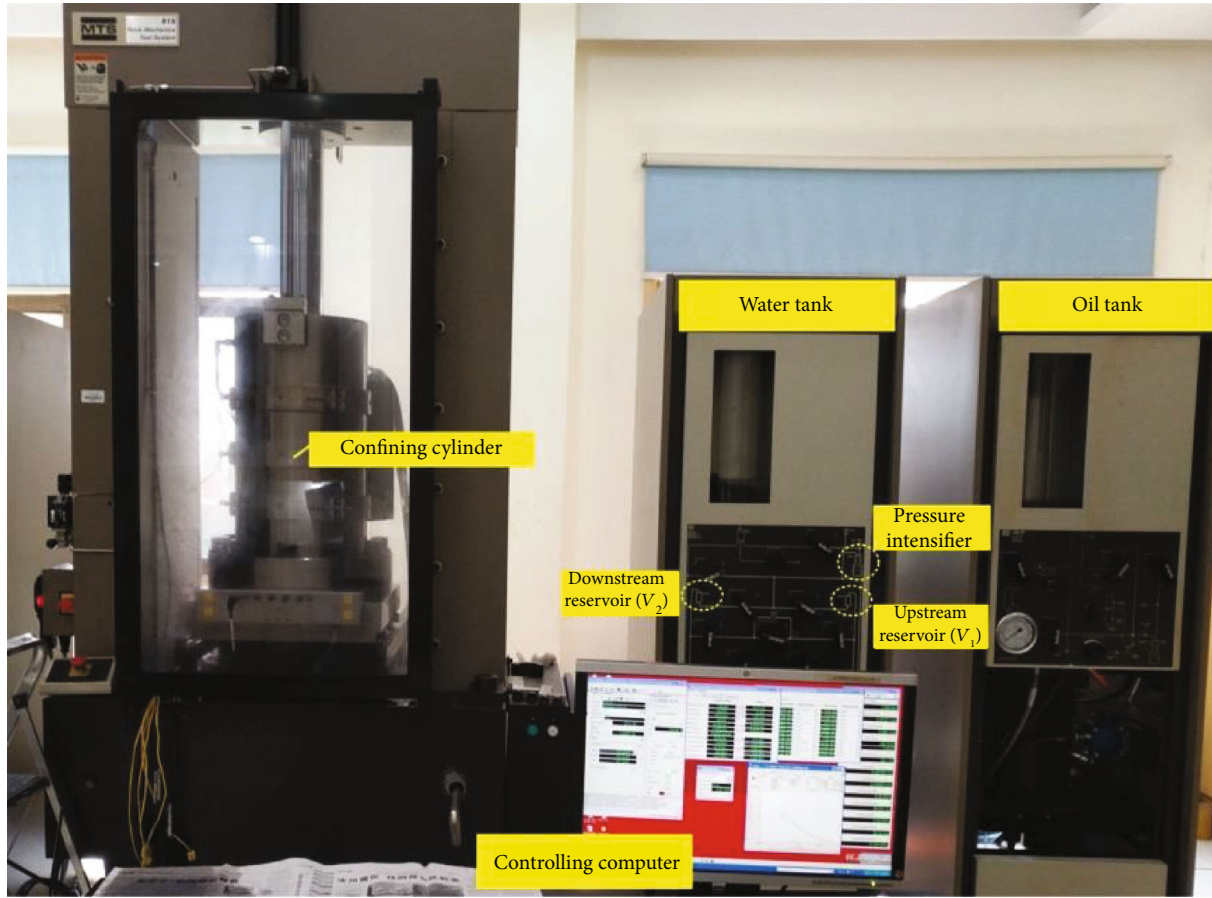


FIGURE 3: Schematic of transient permeability system.

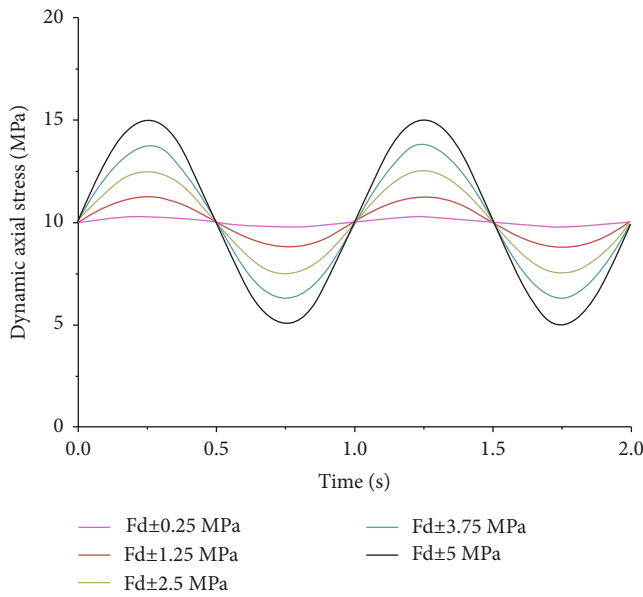


FIGURE 4: Schematic of loading paths.

3.3. Analysis of Gouge Material. Due to the experimental setup, the gouge material was not collected at the outflow end of the experiment, but that left on the fracture surface

after testing was collected. The gouge material was then brushed off the surface and weighed. The weight of gouge material collected after testing is shown in Figures 14 and 15. The weights of gouge material on fracture surface of samples Sd-1, Sd-2, and Sd-3 under dynamic stress are 1.20 g, 1.23 g, and 1.44 g, while those on fracture surface of samples Ss-1, Ss-2, and Ss-3 under static stress are 2.07 g, 2.11 g, and 2.24 g.

The mean value drop of JRC and the weight of the generated gouge materials under dynamic stress were compared as shown in Figure 14. The initial mean value of JRC of Sd-1 is smaller than that of Sd-2, and the initial mean value of JRC of Sd-1 and Sd-2 is lower than that of Sd-3 (Table 1), while the drop of mean value of JRC and the weights of fracture gouge produced by Sd-1 and Sd-2 are lower than those of Sd-3 during dynamic loading.

The same phenomenon occurs under static stress in Figure 15. The initial mean value of JRC of Ss-1 and Ss-2 is smaller than that of Ss-3 under axial static stress (Table 1). The drop of mean value of JRC and the weights of fracture gouge produced by Ss-1 and Ss-3 are lower than those of Ss-2 during static loading. This means that the larger mean value of JRC of fracture surface leads to larger surface damage as indicated by the more gouge productions. This observed gouge material production could potentially cause the hysteretic behavior by subsequently clogging flow paths, thus

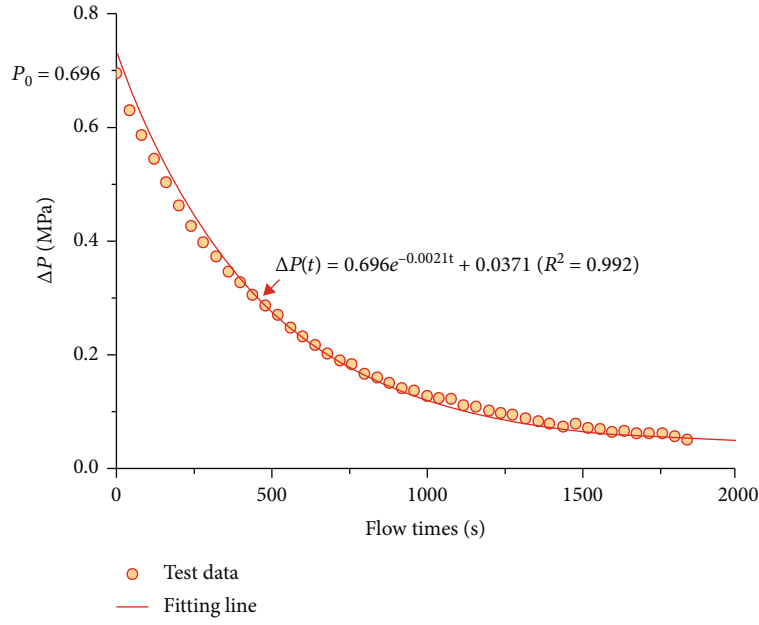


FIGURE 5: A representative differential pressure versus flow time curve.

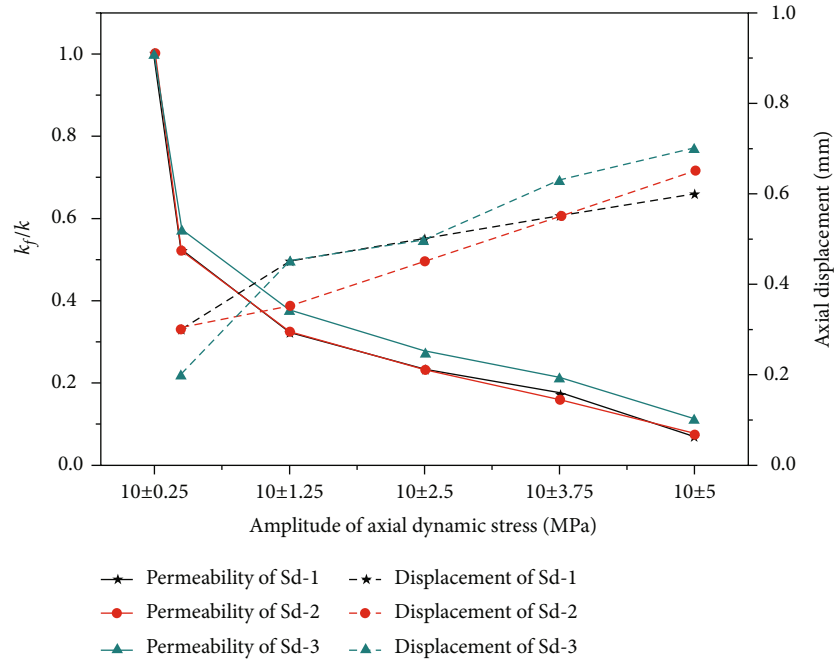


FIGURE 6: Permeability evolution and axial displacement of Sd-1, Sd-2, and Sd-3 under different amplitudes of axial stress of 10 MPa.

lowering the fracture permeability and bringing about the irregular permeability values [4]. This helps to explain that the permeability of Sd-2 drops slower than that of Sd-1 and Sd-2 under dynamic stress (Figure 6). The changes of permeability of Ss-1 and Ss-2 are similar, while the permeability of Ss-3 decreases slowly under static stress (Figure 8). Generally, under the same stress conditions, the accumulation of gouge materials on larger roughness fracture surfaces causes slower drops of permeability.

3.4. Comparing Dynamic Stress and Static Stress. Comparing the permeability changes before and after testing under dynamic and static stresses, the permeability drop, the mean value drop of JRC, and the weight of gouge material are indicated in Figure 16. Under similar roughness profile type for JRC, the permeability drops of fracture surfaces of Sd-1, Sd-2, and Sd-3 more than 35% under axial dynamic stress are larger than the drops of permeability less than 20% under axial static stress. The mean value drop of JRC of Sd-1, Sd-

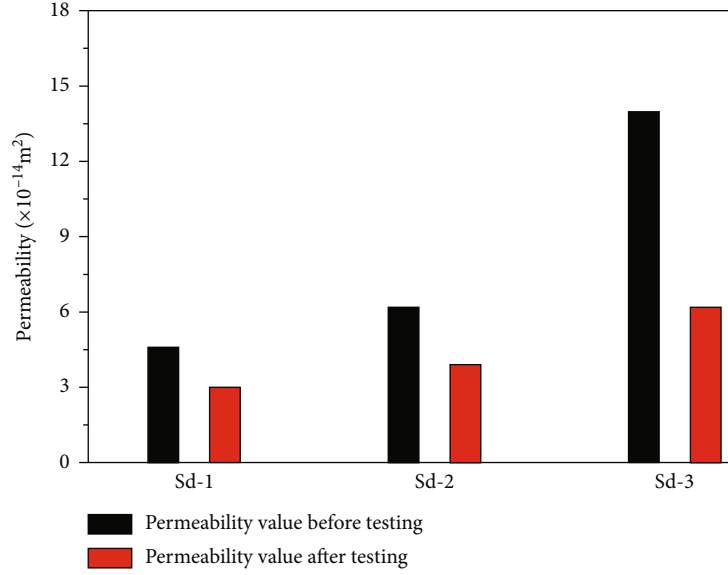


FIGURE 7: Permeability changes before and after testing under axial dynamic stress.

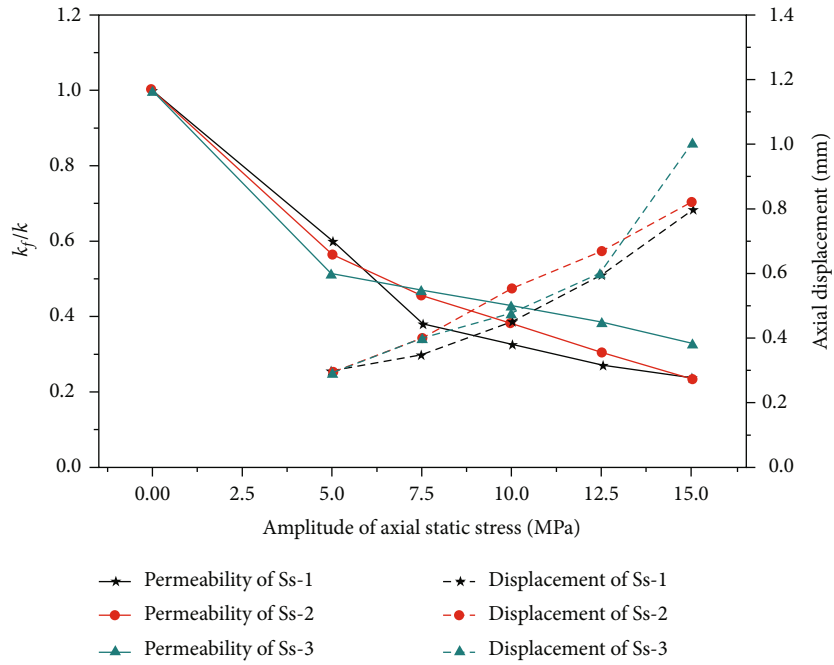


FIGURE 8: Permeability evolution and axial displacement of Ss-1, Ss-2, and Ss-3 under axial static stress.

2, and Sd-3 is larger than that of fracture surfaces of Ss-1, Ss-2, and Ss-3. This suggests that the permeability drops of fracture surfaces under dynamic stress are greater than those under static stress, and the asperity degradations under dynamic stress are larger than those under static stress. However, the weights of gouge materials on fracture surface after testing under axial dynamic stress are smaller than those under axial static stress, which is different from more gouge materials with larger asperity degradations. Overall, under the similar rough surface, the permeability decreases of fracture surfaces are easily caused by dynamic stress through

larger asperity degradations. But fewer gouge materials remain in fracture surfaces under dynamic stress conditions.

These observed responses can be summarized with a proposed conceptual model that considers permeability evolution modes of gouge material transport and flow mobilization, as shown in Figure 17. Figure 17(a) indicates that the gouge materials are more likely to stick to the fracture surface under static stress. When the forces increase, the compacted gouge materials will close the aperture, thus slowing down the rate of permeability reduction of the rock fracture (Figure 17(c)). However, the gouge materials migrate

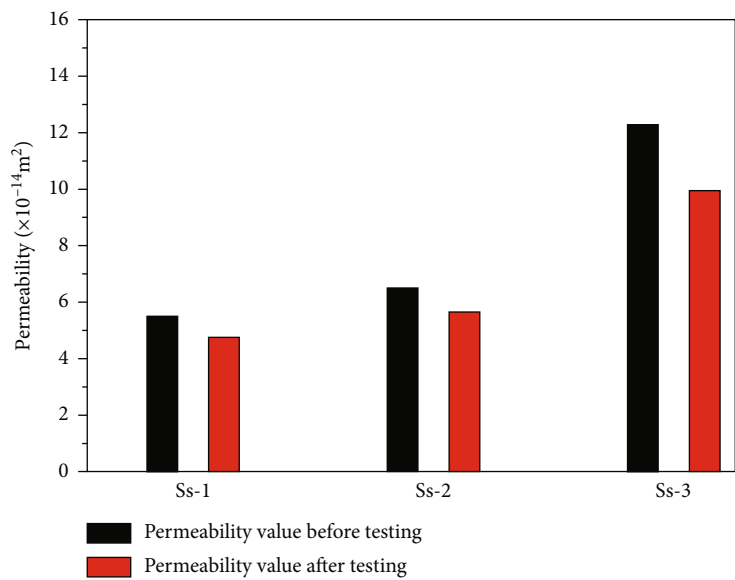


FIGURE 9: Permeability changes before and after testing under axial state stress.

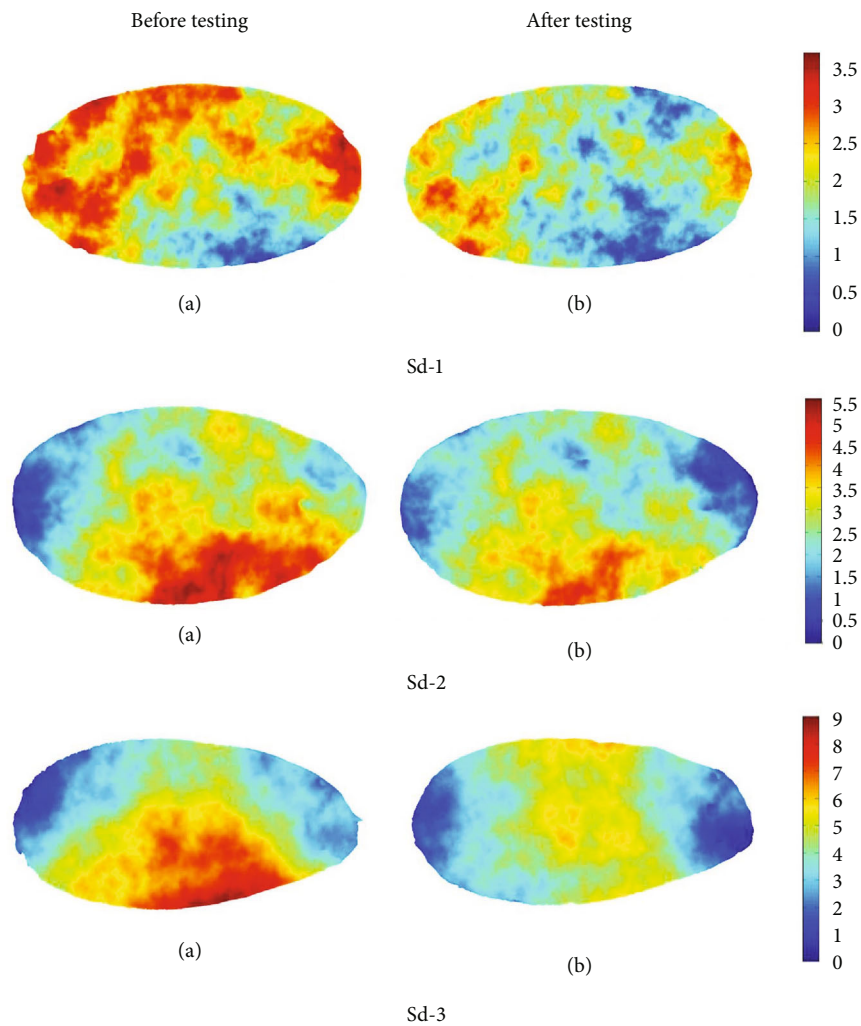


FIGURE 10: (a, b) Surface scans of samples Sd-1, Sd-2, and Sd-3 before and after testing under dynamic stress.

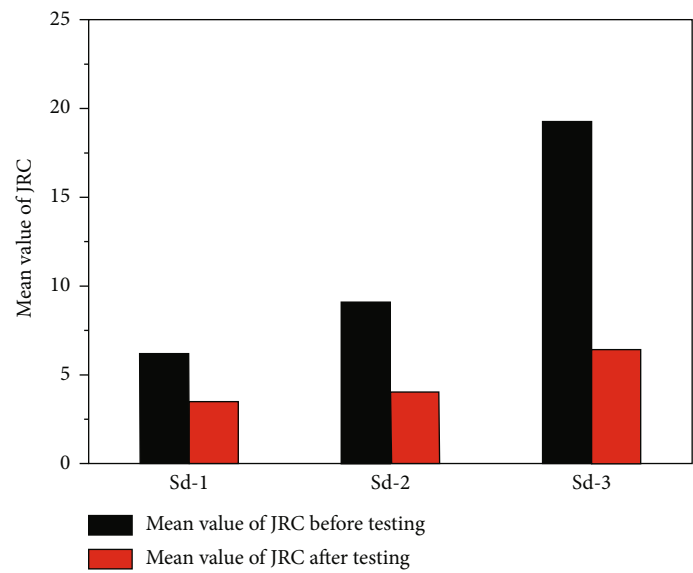


FIGURE 11: Mean value changes of JRC of samples Sd-1, Sd-2, and Sd-3 before and after testing under dynamic stress.

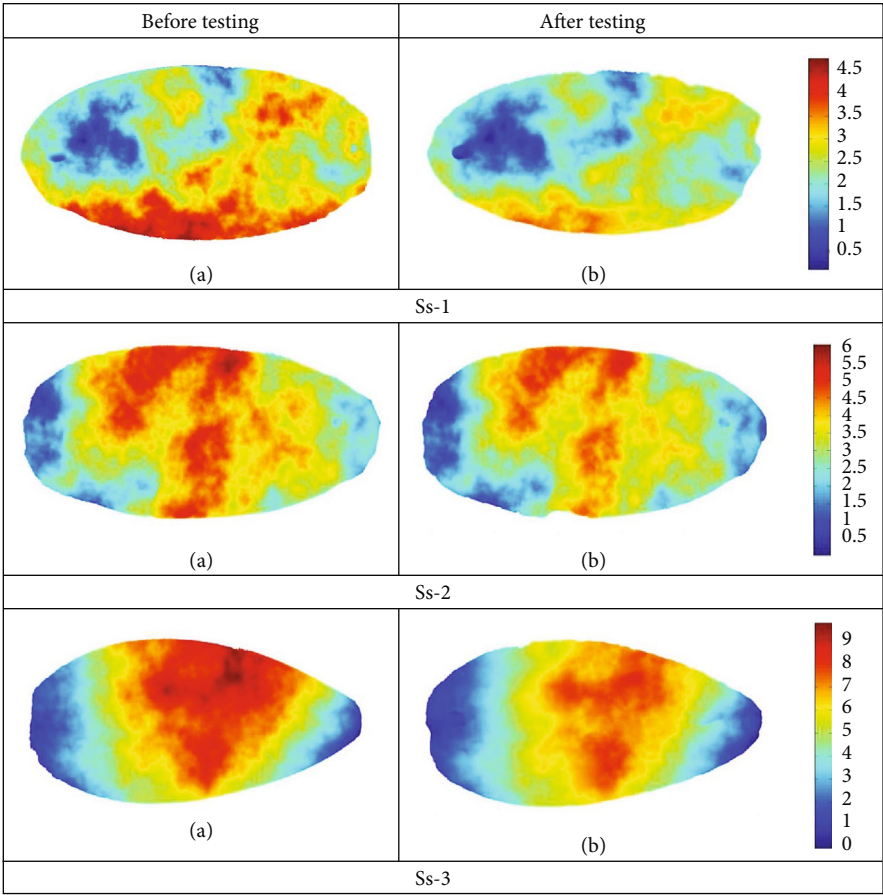


FIGURE 12: (a, b) Surface scans of samples Ss-1, Ss-2, and Ss-3 before and after testing under static stress.

gradually between fracture asperities (or the pore throats) under dynamic stress as described in Figure 17(b). During the amplitude of dynamic stress increasing, the gouge materials initially plugged between fracture asperities (or in the

pore throats) are flushed, producing the permeability drop enhancement under dynamic stress than static stress observed in our experiments (Figure 17(d)). Generally, when the rough fracture surfaces are crushed under more intensive

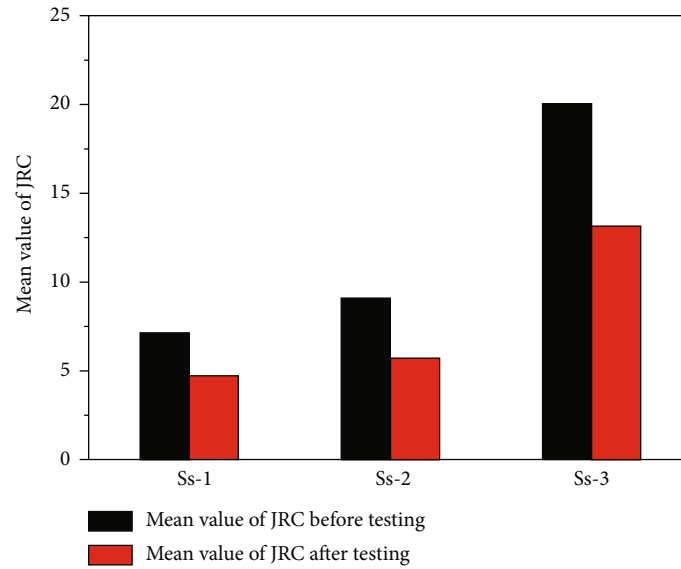


FIGURE 13: Mean value changes of JRC of samples Ss-1, Ss-2, and Ss-3 before and after testing under static stress.

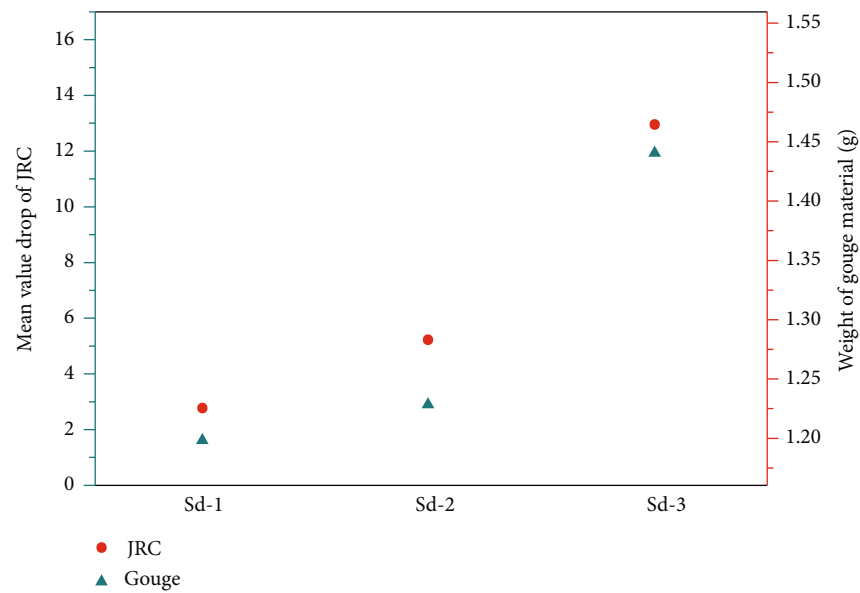


FIGURE 14: Mean value drop of JRC and weights of gouge material after testing under dynamic stress.

multiple dynamic impacts, the transport of gouge materials and the flow mobilization tend to occur during dynamic stress, causing the permeability drops of axial dynamic stress to be larger than those of axial static stress.

4. Conclusions

The permeability changes of nature rock fracture were investigated before and after testing under axial dynamic stress and static stress, so as to simulate the mechanisms of permeability changes of preexisting fractures via dynamic stress in fracture rock engineering. Six series of experiments under different amplitudes of cyclic axial forces and different axial static forces were conducted; the influence of surface rough-

ness was investigated by scanning the specimen surfaces before and after testing; and by which, gouge material transport and flow mobilization of rock fracture were analyzed. The main findings of this paper can be summarized as follows:

- (1) The axial displacements of Sd-1 and Sd-2 are lower than those of Sd-3 with dynamic stress, but the permeability of Sd-1 and Sd-2 is larger than that of Sd-3. In fact, the initial maximum and the mean value of JRC of fracture surfaces of Sd-1 and Sd-2 are smaller than those of Sd-3. This indicates that the roughness of fracture surface has a great influence on the permeability when the axial displacement is not enough to cause the fracture rock to slip

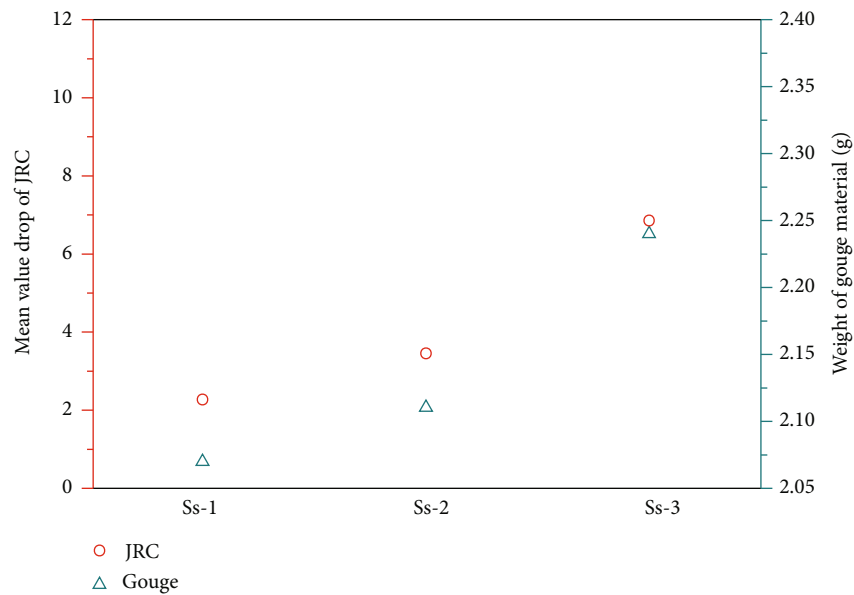


FIGURE 15: Mean value drop of JRC and weights of gouge material after testing under static stress.

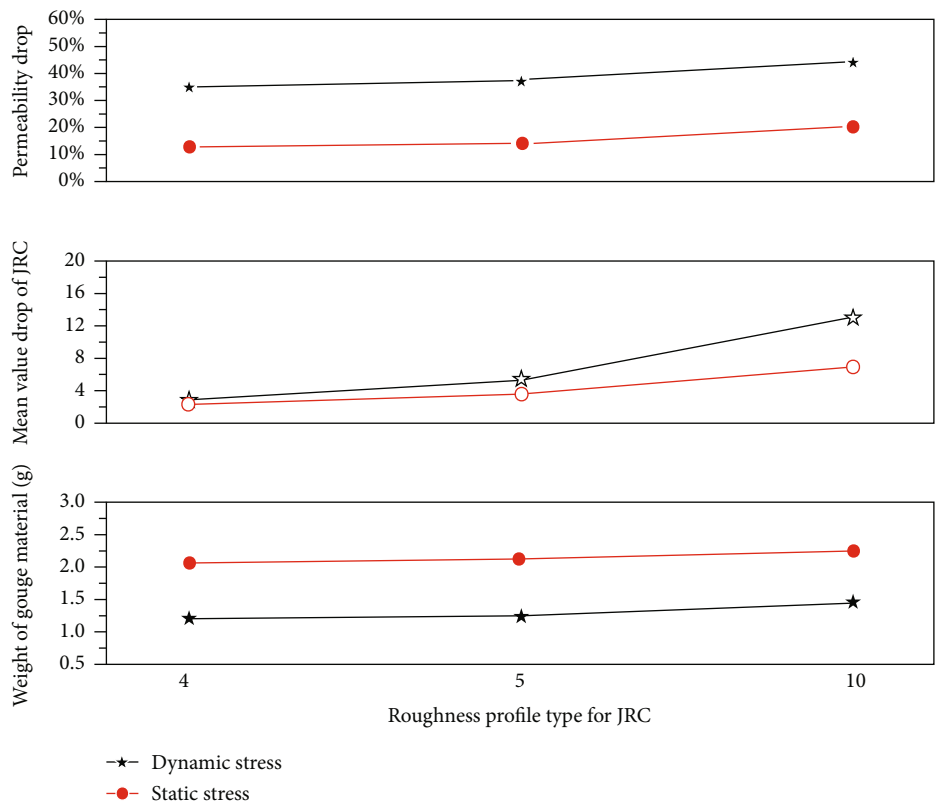


FIGURE 16: Comparing permeability drop, mean value drop of JRC, and weight of gouge material after testing under dynamic stress and static stress.

(2) The initial mean value drop of JRC of Sd-1 and Sd-2 is smaller than that of Sd-3 under dynamic stress, and the weights of fracture gouge produced of Sd-1 and Sd-2 are lower than those of Sd-3 during dynamic loading. This means that the larger

roughness fracture surface leads to larger surface damage as indicated by more gouge production. The accumulation of gouge materials on larger roughness fracture surfaces causes slow drops of permeability

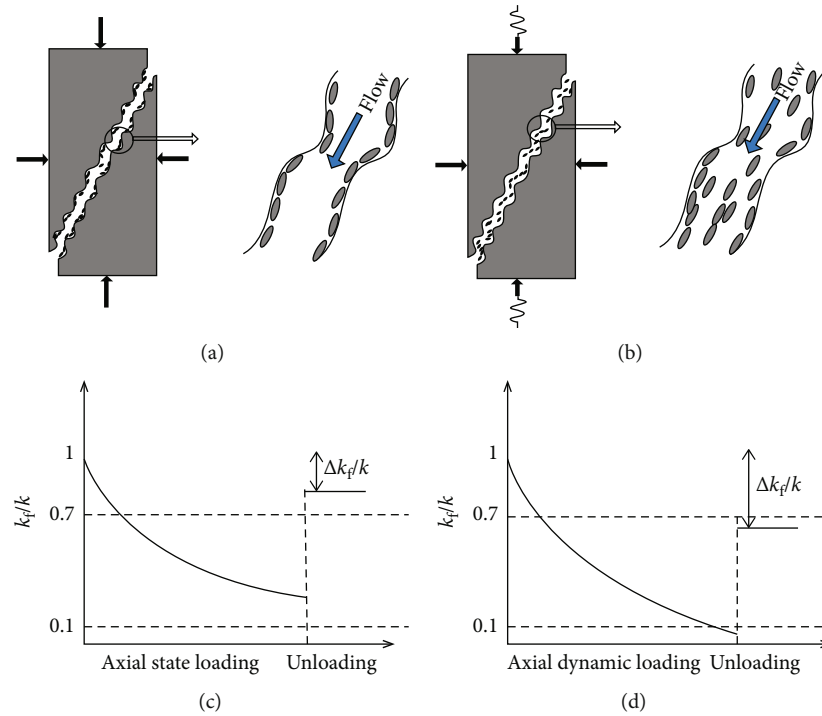


FIGURE 17: Permeability evolution modes of gouge material transport and flow mobilization. (a, b) Gouge material transport and flow mobilization under static stress and dynamic stress. (c, d) Permeability changes under static stress and dynamic stress.

- (3) The permeability evolution modes considering gouge material transport and flow mobilization indicate that the gouge material transport and the flow mobilization tend to occur during dynamic stress. When the rough fracture surfaces are crushed under more intensive multiple dynamic impacts, the permeability drops of axial dynamic stress are larger than those of axial static stress. This interprets that the asperity degradations on fracture surface after testing under axial dynamic stress are larger than those under axial static stress. However, the weights of gouge materials on fracture surface after testing under axial dynamic stress are smaller than those under axial static stress

Data Availability

Data is included in the manuscript.

Conflicts of Interest

The authors declare that there is no conflict of interest regarding the publication of this paper.

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